



Quantitative Evaluation of Tight Sandstone Reservoir Based on Diagenetic Facies—A Case of Lower Silurian Kepingtage Formation in Shuntuoguole Low Uplift, Tarim Basin, China

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The tight sandstone reservoirs of the Lower Silurian Kepingtage Formation are important exploratory targets for tight gas resources in the Shuntuoguole Low Uplift of Tarim Basin. How to evaluate tight sandstone reservoir is an urgent problem to be solved. In this study, we investigated the effects of diagenesis on the heterogeneity of tight sandstone deposits in similar sedimentary facies and established the relationship between the diagenetic facies and reservoir guality. Cores of the tight sandstone reservoirs of Lower Silurian Kepingtage Formation in Shuntuoguole Low Uplift are studied with thin section observation, SEM, XRD, and mercury injection. Quantification of diagenesis influencing porosity suggests that sandstone densification is mainly controlled by compaction, cementation, and hydrocarbon charging (bitumen charging), and the reservoir properties are effectively improved by dissolution, based on which 6 types of diagenetic facies are classified. Interpretation of the log data from individual wells with "K nearest neighbor" algorithm concludes that top and base of the upper member of Kepingtage Formation are believed to have favorably diagenetic reservoirs mainly falling in Type V; favorably diagenetic facies develop best in the lower member of Kepingtage Formation predominated by Types V and VI which mainly distribute in its top. Composite analysis of diagenetic facies, sedimentary facies, and porosity distribution shows that the favorable area of further exploration and development is east of Well SH903 and north of Well SH10. The quantitative identification of diagenetic facies based on logging information can provide reasonable results for the evolution of the tight sandstone reservoirs for a similar area in the Tarim Basin.

Keywords: tight sandstone, diagenesis, quantitative evaluation, Kepingtage formation, Tarim basin

INTRODUCTION

The evaluation and prediction of the tight sandstone reservoirs have been hotspots for the petroleum industry (Guo et al., 2019; Ren et al., 2019; Xiao et al., 2019; Li et al., 2020). The actual exploration effect of the traditional method for the evaluation of tight sandstone reservoirs is poor, which is mainly because the fact that the reservoir physical parameters can not reflect the strong heterogeneity of tight sandstone reservoir (Nian et al., 2016; Zhu et al., 2016; Zhou et al., 2019; Qiao et al., 2020). Continuous researches on diagenesis have made significant achievements in which multi-technological quantitative evaluation is used (Hao et al., 2010; Guo et al., 2018; Qin et al., 2018; Xi et al., 2019). Rapid quantitative evaluation of diagenesis based on diagenetic facies study is the most widely used now and will be the major methodology (Jiang et al., 2014; Zhu et al., 2016; Cui Y et al., 2017). Diagenetic facies including mineral constituents, cement, fabric, pores, and fractures, which reflects the current mineral composition and fabric features and is the genetic mark representing reservoir nature, type, and quality, is the result of sediments which underwent diagenesis and a period of diagenetic evolution (Zou et al., 2008; Li et al., 2019; Li et al., 2020). Diagenetic facies study focusing on the effects of diagenesis process on reservoir properties and pore configuration (Ochoa, 2010; Cui et al., 2017; Lai et al., 2018) is to effectively help hydrocarbon exploration and development by determining the spatial distribution of high-quality reservoir to provide reliable evidence for reservoir prediction (Maast et al., 2011; Meng et al., 2015; Li et al., 2019).

The previous study of diagenetic facies mainly relied on relative core analysis, especially on data of thin section, SEM, and cathode luminescence reflecting microscopic features (Lai et al., 2013; Qin et al., 2018). Limitation of core data results in the inability to show the vertical variation of the diagenetic facies whereas log data continuously record all the physical properties of the penetrated layers (Aysen et al., 2011; Cui et al., 2017; Lai et al., 2019; Zhou et al., 2019). Thus, core-based diagenetic facies identification combined with related well log responses to different diagenetic facies leads to the identification of vertically continuous diagenetic facies in individual wells, so that diagenesis can be evaluated quantitatively and rapidly.

Lower Silurian Kepingtage Formation in Shuntuoguole low uplift is typical tight sandstone with very high heterogeneity (Zhao et al., 2015; Yin et al., 2017; Wang et al., 2020). In relatively stable tectonic settings, sedimentation is the basis of reservoir development while diagenesis which plays a vital role in the intensification of sandstones determines the reservoir quality and distribution. Therefore, systematical quantitative studies of diagenesis and diagenetic facies result in the identification of the high-quality reservoir distribution pattern and are of great significance to the prediction of Silurian tight reservoirs in the study area.

GEOLOGIC SETTING

Located in the Xinjiang Uygur Autonomous Region, northwest of China, Tarim Basin is a large composite sedimentary basin developed on the Presinian metabase. Shuntuoguole low uplift, lying in the north of central uplift zone of Tarim Basin with Awati Sag to the west, Manjiaer Sag to the east, Tabei Uplift to the north, and Tazhong Uplift to the south, is a "saddle" hinge structure between two positive south-north structural belts and two negative east-west structural belts, which is wide in the northern and narrow in the southern (Xiong et al., 2013). It is a sub-low-uplift on the northern slope of central Tarim Basin (Figure 1A). The Paleozoic, Mesozoic, and Cenozoic strata are completely distributed, and the Ordovician, Silurian, Devonian, and Jurassic Cretaceous are missing in some parts, so the overall tectonic activity is relatively stable (Li et al., 2012). The Silurian in the study area is mainly a set of clastic rocks of shoreside-shallow shelf facies, lithologically composed of gray-green argillaceous siltstone, gray mudstone, and purple-red siltstone (Zhao et al., 2015). The Silurian is composed of the Kepingtage Formation, the Tataaiertage Formation, and the Imutantawu Formation. The research horizon is the Kepingtage Formation mainly divided into a lower member (S_1k^1) , middle member (S_1k^2) , and upper member (S_1k^3) . The lower member (S1k1) is composed of Barrier bar and Shoreface interbedded deposits, and sandstone reservoir developed (Sun et al., 2013). The upper member (S_1k^3) developed the delta front environment, mainly developing underwater distributary channels, mouth bar, and underwater distributary interchannel microfacies (Zeng et al., 2018; Yin et al., 2017). Gray mudstone developed in the middle member (S_1k^2) and upper part of upper member (S_1k^3) , constitutes a good cap layer in the region, thus forming several sets of favorable reservoir-cap assemblages in the study area (Figure 1B). They are the major exploration targets for Paleozoic clastic rock hydrocarbon exploration in the Shuntuoguole region (Xiong et al., 2013; Cai et al., 2014; Peng et al., 2018; Wang et al., 2020).

SAMPLES AND METHODS

This study takes 300 cores (sampling from 5,164–5,712 m) from 8 wells in the low uplift in Shuntuoguole. 308 thin sections with each one having 200 points are made for a series of tests, including petrology, petrophysical property, porous structure, and diagenesis. Polarizing microscope (LV100POL) is used to observe the mineral composition, shape, texture, optical feature, and type and size of the pores. Microscopic shape and texture are observed for 55 samples using a Scanning electronic microscope (XL30). Clay mineralogical content is analyzed for 53 samples through X-ray diffraction (D/max2,500 PC). Physical properties and pore configuration of the sandstone are analyzed for 122 samples using mercury intrusion porosimeter (AutoPore9500). The original porosity of the sandstone is restored through particle size analyzer (SFY-D Acoustic Screener and CG-1 Centrifugal Settling Particle Size Distribution Analyzer).

RESULTS

Reservoir Lithologies

Kepingtage Formation mainly consists of fine sandstone (grain sizes ranging from 0.09 to 0.25 mm) which contains litharenite



FIGURE 1 | Structural location and well location distribution in the study area. (A) Location of Shuntuoguole low uplift and surrounding structural features. (B) Lithological profile and reservoir-caprock assemblage of Kepingtage Formation (S_1k^1 , totally 90 ~ 150m thick shoreside (non-barrier coast)-shallow marine shelf facies deposit that is different from common high-energy shoreside due to low-energy water body, mainly consists of fine-grain sandstone with medium-grain sandstone and siltstone in the upper part and mudstone in the lower part. S_1k^2 , 43 ~ 59m thick, deposited in the shallow marine shelf. S_1k^3 , 287 ~ 342m thick, depositing in delta (tidal action) depositional system, mainly consists of silt-fine sandstone interbedded with silty mudstones, with a 15 ~ 30m mudstone in its middle part. The three sandstone members, S_1k^{3-3} , S_1k^{3-1} , and S_1k^1 , containing reservoirs with abundant oil and gas, form several favorable reservoir-cap assemblages in combination with S_1k^{3-2} , S_1k^2 , and S_1t^1).

(61.86%), sublitharenite (17.63%), and feldspathic litharenite (11.22%) (**Figure 2**).

The clastic constituents predominantly comprise quartz (62.0%) that mainly contain single-crystal quartz with some chert. They comprise less feldspar (7.0%), which is commonly sericitized, and K-spar content is higher than plagioclase. The proportion of debris in the rock fragments which primarily consist of metamorphic rock debris and sedimentary rock debris. Besides, heavy minerals like zircon,

tourmaline, garnet, apatite, and pyrite can be seen despite their very low content. Grains sorting is good and poorly rounded.

The cement account for 66.7% of interstitial materials, and the first biggest proportion is calcareous cement (45.7%) with calcite that is inhomogeneous as the main composition. The second is siliceous cement (16.7%) which is dominated by quartz overgrowth and intergranular authigenic quartz. There is very little clay (0.4%) which is mainly illite-smectite hybrid, illite, and



diagram from Zhao and Zhu, 2006).

chlorite. Moreover, argillaceous matrix (27%) and slight bitumen (6.3%) are also seen in the interstitial materials.

Reservoir Properties

Kepingtage Formation sandstone mainly contains secondary dissolved pores, primary intergranular pores, microfractures, and composition matrix pores. The secondary dissolved pores, accounting for about 52.3% of plane porosity, include intergranular dissolved pores that are the most dominant, intragranular dissolved pores, intercrystal pores, and moldic pores (**Figure 3A**). Primary intergranular pores (Figure 3b) take 39.2% while microfractures that are commonly filled by bitumen or argillaceous material only take about 2.2% of the plane porosity. For the three sandstone members, S_1k^{3-3} is dominated by primary intergranular pores, whereas S_1k^{3-1} and S_1k^1 dominantly contain secondary intergranular pores (**Figure 4**).

Core-based data from 1,400 samples show that Kepingtage Formation has generally poor physical properties. Porosity

ranges from 0.5% to 16.4% and averages 7.0%, while air permeability, which is mainly 0.1–1 mD, ranges from 0.01 to 15.4 mD and averages 0.54 mD (**Figure 5**). Physical properties of the three sandstone members worsen increasingly with deeper burial: S_1k^{3-3} porosity is 9.1% with an air permeability of 1.34 mD; S_1k^{3-1} porosity is 7.50% with an air permeability of 0.57mD; S_1k^1 porosity is 6.1% with an air permeability of 0.32 mD.

Diagenesis

Kepingtage Formation underwent compaction, cementation, dissolution, metasomasis, hydrocarbon charging, and structural fracturing. It is the reworking of the multiple diageneses that leads to the property variations of different sandstone zones (Carvalho et al., 2014; Zhao et al., 2015; Peng, et al., 2018; Zeng et al., 2018).

Compaction which worked throughout the burial history of the sandstone reservoir (Zhou et al., 2008) reduced the primary pores sharply at the early stage of diagenesis (Paxton et al., 2002). Kepingtage Formation experienced strong compaction which mainly accounts for the reduction of the Petrophysical



FIGURE 3 | Microphotographs of Kepingtage Formation reservoir rocks (A) Intergranular dissolved pores and intragranular dissolved pores, fine to medium grained litharenite, Well Sh904H, 5576.58m, 10 × 10, (-) (B) Residual primary intergranular pores and intergranular dissolved pores, fine-medium grained quartz litharenite, Well Sh901, 5503.62m, 10 × 10, (-) (C) Intergrowth calcite cement, fine to medium grained litharenite, Well Sh904H, 5584.40m, 10 × 10, (+) (D) Quartz overgrowth, fine to medium grained litharenite, Well Sh904H, 5511.08m, 10 × 20, (+) (E) Particle surface with filamentous and laminated lilite-smectite mixed layer, fine-medium grained quartz litharenite, Well Sh904H, 5367.66m, SEM, 300×. (F) Particle surface with laminated chlorite, fine to medium grained litharenite, Well Sh2, 5678.26m, SEM 3000×.



property. Deep burial and high content of rock fragments which caused Kepingtage Formation to undergo strong compaction are the predominant causes of reduced porosity. Mechanical compaction was dominant and pressolution functioned little. Microscope-based analysis of compaction finds 1) that intergranular contact changes from point contact to line contact, even to concave-convex contact (**Figure 3b**), 2) that

plastic grains like mica and soft debris became deformed and even fractured, and 3) that rigid grains such as quartz and feldspar misplaced or broken under compression.

Cementation is common in Kepingtage Formation, including calcareous cementation, siliceous cementation, and clay cementation. Cement generally reduces the porosity due to their filling up the pore space, but they can also resist



compaction by enhancing the frame strength (Zaid, 2013; Yuan et al., 2015; Peng, 2018).

- (1) Calcareous cementation. Calcareous cement primarily consists of calcite. Microscope-based observation and thermometry of primary inclusion captured by the calcite-cement show that early- and late-stage calcite cementations occurred. The early-stage cement has intergrowth texture with alternation of rock fragments of which quartz has no overgrowth at their edges (Figure 3C). The homogenization temperature of the inclusion ranges from 80°C to 100°C. The late-stage calcite cement is commonly clot-shaped or granular with quartz overgrowth. The homogenization temperature of the inclusion ranges from 140 to 150°C.
- (2) Siliceous cementation. In sandstones with intense calcareous cementation, siliceous cement, developing to a relatively low degree, have quartz overgrowth around the surface in an ear-like or irregular ring-like shape, with the enlarged edge of 0.01–0.08 mm thick (Figure 3D).
- (3) Clay cementation. Clay cement that mainly contains illite, illite-smectite mixed layer, and chlorite, with very little kaolinite and chlorite-smectite mixed layer. SEM-based observation shows that illite occurs in sheet-like and cellular shapes on the clay particle surface, or extends outward in filament shape along the surface. The illite-smectite mixed layer resembles the irregular sheet of illite in morphology (Figure 3E) while chlorite whose individual morphology resembles niddles commonly coats the particle (Figure 3F).

Dissolution plays a vital role in the improvement of the physical properties of the reservoir. According to the SEM observation, dissolution occurs at the feldspar particle edge or cleavage where soluble compositions are dissolved, and residual part is sieve-like or wreckage-like with intergranular pores and intragranular pores (**Figure 3A**); feldspar is entirely eroded where mold pores are formed.

Metasomatism indicates rock fragments were alternated with calcite, and with feldspar and quartz, or with pyrite, and the alternated fragment particle has sawtooth-like or embayed edges (**Figure 3C**).

Hydrocarbon charging slows down the cementation rate and accelerates dissolution (Li et al., 2003; Molenaar et al., 2008; Guo et al., 2018; Wang et al., 2020), but residual bitumen filling the pores reduces porosity. Two stages of bitumen charging are identified. Residual bitumen during the first charging stage filled the primary intergranular pores along the eroded edge of the grains (**Figure 6A**), and the oil inclusion displays light grayyellow or light yellow fluorescence (**Figure 6B**). Residual bitumen during the second charging filled the space beyond the quartz overgrowth edge (**Figure 6C**) and the dissolved pores of calcite cement, and the oil inclusion displays light bluish-green, light green, or green fluorescence (**Figure 6D**).

QUANTITATIVE EVALUATION OF DIAGENESIS

Diagenesis and Porosity

Diagenesis controls the pore evolution of the sandstone reservoir (Pan et al., 2011; Zhao, 2015; Peng, 2018; Li et al., 2019; Xiao et al., 2019). Based on the above diagenesis study, a quantitative evaluate method of the effect of diagenetic events on sandstone porosity was performed using a screen test of particle size and thin section identification. First of all, the relationship between unconsolidated sandstone porosity and sorting coefficient under Earth surface condition (Beard and Weyl, 1973) is used to calculate initial porosity:



FIGURE 6 | Microphotography of reservoir bitumen and fluorescent photography of inclusion (A) Bitumen filling along the corroded edge of the grains, Well Sh9, 5335.79m, 10 × 10 (-) (B) The hydrocarbon inclusions developed in quartz grains with a beaded distribution in the cracks, displaying yellow fluorescence, fine grayish sandstone with oily spots, Well Sh9, 5581.89m. (C) Bitumen filling into the dissolved pores of calcite cement later than quartz overgrowth, Well Sh9, 5476.48m, 10 × 10 (-). (D) The hydrocarbon inclusions develop in quartz grain cracks, which are scattered and uneven in size, displaying bluish-green fluorescence, fine grayish sandstone with oily spots, Well Sh9, 5580.4m.

Initial porosity =
$$20.91 + 22.90/S_0$$
 (1)

where Po is the Trask sorting coefficient.

And then the contents of different cement, different pore types, and bitumen are calculated. Effects of different diagenetic events on porosity are calculated quantitatively using the relationship between the plane porosity and measured porosity (**Table 1**):

 $Compaction-Reduced\ porosity-initial\ porosity-(total\ cement content$

+ bitumen content + primary intergranular pore volume

 \times measured porosity/plane porosity)

(2) Cementation – Reduced porosity = cement content (3)

Bitumen - charging - reduced porosity = bitumen content (4)

$$\times$$
 measured porosity/plane porosity (5)

The quantitative calculation shows that the Kepingtage Formation is densified due to compaction, cementation, and hydrocarbon charging (bitumen filling) whereas its reservoir porosity is enhanced due to dissolution. Comparison of the evolution of the three sandstone members of Kepingtage Formation suggests that compaction influenced S_1K^1 most due to its biggest burial depth and its highest content of plastic grains of rock fragments. The physical properties of S_1K^{3-3} are best because of secondary pores generated from relatively good dissolution though a lot of pore space was filled by calcareous cement and sandstone porosity was affected by cementation. None of the three sandstone members of Kepingtage Formation has been influenced significantly by clay cementation. Besides, bitumen filling has also reduced porosity, which is the most obvious in the S_1K^{3-1} sandstone reservoir.

Classification and Characteristics of Diagenetic Facies

Since there has been no unified scenario for classification and naming of diagenetic facies so far (Zou et al., 2008; Cui et al., 2017; Lai et al., 2018; Li et al., 2019), this study proposes "cause (type and intensity of diagenesis) + effect (influence on the reservoir)" scenario on the basis of practicability and quantification. Based on the characteristics and differences

TABLE 1 | Relationship between diagenesis and porosity of Kepingtage Formation.

Sub- member	Well	P _i (%)	Compa- ction RP due to compaction (%)	Cementation			Hydro- Carbon charging	Dissolution				Current porosity (%)	Number of data gathering	
				RP due to carbonate cementation (%)	RP due to siliceous cementation (%)	RP due to clay cementation (%)	Total RP due to cementation (%)	RP due to bitumen filling (%)	IP due to feldspar dissolution (%)	IP due to rock fragment dissolution (%)	IP due to quartz dissolution (%)	Total IP due to dissolution (%)	p	point
S1K3-3	SH9	36.18	19.74	5.63	1.40	0.13	7.16	1.95	2.12	0.76	0.15	3.03	10.36	8
	SH902H	38.53	23.53	8.10	1.45	0.00	9.55	2.10	3.40	1.21	0.24	4.85	8.20	10
	SX2	37.63	16.68	11.50	2.30	0.00	13.80	3.75	4.38	1.56	0.31	6.25	9.65	4
Average		37.45	19.98	8.41	1.72	0.04	10.17	2.60	3.30	1.18	0.24	4.72	9.42	
S1K3-1	SH9	37.04	21.19	9	0.58	0.18	9.76	3.52	3.47	1.73	0.58	5.78	8.35	40
	SH902H	37.27	24.13	5.55	1.56	0.47	7.58	3.61	2.53	1.27	0.42	4.22	6.17	33
	SH904H	34.54	18.98	4.1	2.05	1.07	7.22	4.26	0.53	0.26	0.09	0.88	4.96	28
Average		36.28	21.43	6.22	1.40	0.57	8.19	3.80	2.18	1.09	0.36	3.63	6.49	
S_1K^1	SH9	34.15	24.22	4.18	1.24	0.08	5.50	0.75	1.46	1.12	0.08	2.66	6.34	31
	SH901	31.18	21.87	3.31	2.31	0.66	6.28	0.86	1.66	1.27	0.09	3.02	5.19	19
	SH902H	36.18	28.04	1.54	3.55	0.02	5.11	1.62	2.23	1.71	0.12	4.06	5.47	39
	SH904H	33.03	22.11	5.15	2.00	0.67	7.82	1.55	2.07	1.58	0.11	3.76	5.31	22
Average		33.64	24.06	3.55	2.28	0.36	6.18	1.20	1.86	1.42	0.10	3.38	5.58	

RP, reduced porosity; IP, Incremental porosity.

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TABLE 2 | Diagenetic facies types and characteristics of kepingtage formation.

DF	Types of sand-stone	CC (%)	CR	CCC (%)	BC (%)	PIPV (%)	SDV (%)	Por (%)	Perm (mD)	DP (MPa)
I	Litharenite, sublitharenite	60.3–100, 75.2	Line, and line-concave/convex	0–8, 1.82	0.5–5, 2.65	0–3.1, 1.08	0–5.4, 2.49	2.5–5.4, 4.38	0.01–3.4, 0.34	0.2–3, 1.54
П	Litharenite, sublitharenite	3.8-61.4, 30.9	Point-line	9–38, 19.14	1-6, 3.67	0-2.9, 0.98	0-4.5, 1.71	1.7-5.4, 3.44	0.01–0.88, 0.17	0.46-3.97, 1.35
Ш	Sublitharenite	41.0–77.3, 61.7	Point-line	0–7, 2.53	7–17, 10.9	0.51–3.77, 2.51	0–1.53, 0.52	0.6-7, 4.49	0.06-0.55, 0.34	_
IV	Quartzarenite, sublitharenite	17.0-69.1, 51.7	Point, and point-line	0-5, 2.51	0-4, 0.52	4.97-12.13, 7.29	0-3, 1.16	5.8–13.4, 8.65	0.1-6.02, 0.93	0.15–2.58, 0.91
V	Litharenite	30.9-82.2, 58.7	Point-line, and line	0-9, 2.55	0.2–5, 1.6	1.11-8.56, 3.75	2.13–7.3, 3.88	5.5-13.7, 7.66	0.03-5.59, 0.47	0.15-2.09, 1.03
VI	Litharenite	<u>22.1–71.9,</u> 59.5	Point-line, and line	0–9, 2.67	0	<u>0–4.32,</u> 1.2	<u>5.52–13.26,</u> 7.74	<u>6.1–15.6.</u> 9.35	<u>0.02</u> – <u>6.54</u> , 0.75	0.1-2.96, 1.1

of diagenetic facies, this scenario classifies Kepingtage Formation sandstones into 6 types of diagenetic facies with parameters like compaction coefficient, calcareous cement content, bitumen content, primary intergranular pore volume, and secondary pore volume (from thin section analysis), and measured physical properties, and pore texture (from mercury intrusion): tight facies from strong compaction (Type I), tight facies from calcareous cementation (Type II), tight facies from bitumen filling (Type III), porosity preservation facies from weak diagenesis (Type IV), porosity enhancement facies from medium dissolution (Type V), and porosity generation facies from strong dissolution (Type VI). According to the clastic rock reservoir evaluation criteria of the Kepingtage Formation (Cai et al., 2014), Types I, II, and III are unfavorable diagenetic facies while Types IV, V, and VI are favorable. The characteristics of each diagenetic facies are shown in Table 2.

Automatic Identification of Diagenetic Facies

The effective identification of diagenetic facies is a difficult problem in reservoir evaluation. Due to the limitation of the core data, it is a common method to characterize sandstone reservoir by logging curve (Zhou et al., 2008; Hao et al., 2010; Qiao et al., 2020; Zhu et al., 2020). Scholars adopt different methods, including multi-parameter comprehensive identification (Paxton et al., 2002; Jiang et al., 2014; Lai et al., 2019; Li et al., 2019), optimization of Logging Mathematical Model for Identification (Cui et al., 2017; Qin et al., 2018; Guo et al., 2019). The well logs including GR that reflects the lithology and sedimentary environment of sandstone, AC, DEN and CNL that directly indicate porosity differences, and the resistivity logs like RILM and RILD are used to characterize the distribution of diagenetic facies continuously and quantitatively. Picking different logs for specific diagenetic facies as sampled data, this study actualizes the automatic identification of the vertical diagenetic facies using "K nearest neighbor (KNN)" discriminance (He et al., 2014; Lai et al., 2019; Zhou et al., 2019).

The KNN algorithm is to express n term well log value in the training sample set of different diagenetic facies samples as n-dimension characteristic vector, and then to calculate similarity between n-dimension vector point that is formed by well log data to be classified and each n-dimension sector in training sample set. If the very point and the majority of K samples of the nearest neighbor belong to the same diagenetic facies type, the sample belongs to the same type accordingly. In this way, obvious overlap of the well logs for different diagenetic facies types can be solved easily (**Table 3**).

The Algorithm goes as follows

- (1) A training sample set is generated based on the known diagenetic facies classification.
- (2) Similarity between the unknown sample and each sample in the training sample set is calculated using the following equation:

$$S(d_{i}, d_{j}) = \frac{\sum_{k=1}^{M} S_{ik} \times S_{jk}}{\sqrt{\sum_{k=1}^{M} S_{ik}^{2} \sqrt{\sum_{k=1}^{M} S_{jk}^{2}}}}$$
(6)

where d_i is the vector derived from well log value of the unknown sample; d_j is the characteristic vector of well log response for the *j*-th training sample; *M* is the number of dimensions of the characteristic vector; S_{ik} and S_{jk} are the k-th dimension of correspondent vectors, respectively.

(3) Weight for each category of the unknown sample in the known categorized sample set is calculated with the following equation:

$$p(\overline{x}, C_j) = \sum_{d_1 \in D} S\left(\overline{x}, \overline{d_i}\right) y\left(\overline{d_i}, C_j\right)$$
(7)

where \overline{x} is the characteristic vector of the unknown sample; $S(\overline{x}, \overline{d_i})$ is the similarity; $y(\overline{d_i}, C_j)$ is the category property function. If $\overline{d_i}$ belongs to Category C_j , the value of the function is 1, otherwise, the value is 0.

(4) The Unknown sample is categorized into the known category with the biggest weight.

The diagenetic facies of individual wells in the study area is identified with the above algorithm on Matlab software platform (**Figure 7**). To verify the accuracy of the identification result, this study correlates the log-based identified facies with core-based ones. The correlation shows that matching rate is up to 78.6%, and the favorable diagenetic facies corresponds well to the show of oil and gas from the mud log data (**Figure 7**), suggesting the prediction of the diagenetic facies of the non-coring intervals with the above method is reliable. The logging information of tight sandstone can effectively reflect the lithologic and property changes of reservoir, and can be used as a reliable basis for diagenetic facies identification.

Distribution Regularity of Diagenetic Facies

The spatial distribution of different types of diagenetic facies is a key problem in the evaluation of the tight sandstone reservoirs. Traditional diagenetic facies prediction is based on reservoir parameters such as thickness, porosity, and permeability (Zou et al., 2008; Hao et al., 2010; Lai et al., 2013; Guo et al., 2019; Li et al., 2020; Qiao et al., 2020). However, there is a problem of low accuracy under the influence of reservoir heterogeneity. Identification of diagenetic facies logging effectively solves the problem of fewer sample points (He et al., 2014; Cui et al., 2017; Lai et al., 2019). At the same time, the prediction of reservoir main control factors is used for reference and has high accuracy and good practical effect.

The vertical distribution pattern was discussed by calculating proportions of different diagenetic facies in each sandstone member of the Kepingtage Formation (**Figure 8**). Unfavorable

TABLE 3	Bore log	responses f	or	samples	of	different	diagenetic	facies.	

Diagenetic facies	DEN (g cm ⁻³)	CNL (%)	AC (μm m ⁻¹)	GR (API)	RILM (Ω·m)	RILD (Ω⋅m)
	1.99–2.61	6.38–36.03	60.46-107.98	37.48–114.16	3.98–54.89	4.13-41.91
	2.03-2.62	6.57-38.26	60.96-91.88	53.70-99.21	3.91-24.70	4.29-32.05
11	2.40-2.58	8.29-16.81	62.69-73.32	60.96-76.28	5.58-10.04	6.03-9.31
IV	2.13-2.59	6.45-23.87	62.80-88.91	43.98-90.42	4.50-11.54	4.00-13.55
V	2.12-2.62	8.11-21.52	60.73-93.89	54.48-116.01	3.93-13.62	3.24-18.14
VI	1.99-2.62	8.55-22.50	62.88-96.72	54.49-108.46	4.19-51.04	3.45–21.37





diagenetic facies occupies 50% of S_1k^{3-3} , of which Type I takes up 26.9%, Type II 21.6%, and Type III a smaller proportion. All the three types of favorable diagenetic facies develop in this zone, primarily Type V (25.1%), secondarily Type IV (15.9%), and then Type VI. The favorable diagenetic facies predominantly occur in the top and bottom sandstone zones although mud log does not show oil and gas.

The 60% diagenetic facies of S_1k^{3-1} is unfavorable, where Types I, II, and III develop well of which Type III occupies the biggest proportion (13.2%). Favorable diagenetic facies, which do not develop regularly, are mainly Type V and Type VI, and then Type IV. Evidence of oil and gas is commonly discovered in the bottom sandbody.

Favorable diagenetic facies develop best in S_1k^1 , occupying over 75% of which Types V and VI are predominant. They primarily occur in the top part with a great quantity of evidence of oil and gas. Unfavorable diagenetic facies are mainly Type I (22.2%) while the other two are seldom seen. The above analysis is consistent with the result of the quantitative calculation of pore evolution.

Due to the lack of effective prediction methods, the study of diagenetic facies distribution is less (Hao et al., 2010; Pan et al., 2011; Cai et al., 2014; Li et al., 2019). Most scholars believe that sedimentary differentiation, which results in the spatial distribution variation in debris mineral composition, interstitial material content, granule size, and pore fluid property, provides different prerequisites for the later diagenesis, so the diagenetic facies distribution is



significantly controlled by sedimentary facies (Zhang et al., 2011; Zeng et al., 2018; Qiao et al., 2020). In this study, the face-controlling method is used to study the favorable diagenetic facies distribution. Analysis of diagenetic facies and sedimentary facies of core samples combined shows that underwater distributary channels and mouth bar sand

bodies primarily develop in S_1k^{3-3} and S_1k^{3-1} . All types of diagenetic facies occur in the major part of the channel and its peripheries. The sandbody in the main part of the channel is thick and rich in quartz. The mid-part of the sandbody is favorable for primary intergranular pore preservation and the development of Type IV diagenetic facies whereas the upper and lower parts predominantly contain Types V and VI diagenetic facies. On the channel peripheries where the sandbodies are thinner with lower quartz content and stronger calcareous cementation, Type I and Type II diagenetic facies predominantly occur. In the mouth bar where the sandbody rich in quartz is well rounded and well-graded, types of favorable diagenetic facies occur. The Shoreface barrier bars has a high content of quartz and high sorting degree which make it easier for acid fluid to enter commonly occur in S_1k^1 , and mainly develop Type V and Type VI diagenetic facies.

The purpose of diagenetic facies evaluation is to predict the range of favorable diagenetic facies distribution. This study evaluates the favorable diagenetic facies in the upper and lower members of Kepingtage Formation, identifies their distribution pattern, and draws a diagenetic facies distribution map (**Figure 9**) based on the lateral distribution pattern of the sedimentary facies in combination with porosity contours.

According to the diagenetic facies map, Type I facies occurs in S_1k^3 extensively. Type II facies occur in the peripheries of the underwater distributary channel and eastern fluvial channel (inferred) where Well SH903H is located. Type III facies



FIGURE 9 | Distribution of diagenetic facies in Kepingtage Formation (a. S_1k^3 . **B**. $S_1k^{1.}$). (Diagenetic facies boundaries are dotted lines representing the approximate range of distribution, and distribution is based on dominant diagenetic facies in single Well identification and the distribution refers to the sedimentary microfacies. Plane map of sedimentary microfacies are modified after Sun et al., 2013. Porosity contour is mapped according to the plane map of seismic porosity).

typically occurs in Wells SH1 and SH901 area while Type V facies distributes in the area where Wells SH902H and SH904H lie. Porosity distribution shows that the area to the northeast of Well SH903H is the areas with high porosity values, which lie in the fluvial channel fairway. They are supposed to be the Type V facies distribution areas where both primary intergranular pores and secondary dissolved pores develop. Type VI diagenetic facies mainly distributes in the Well SH9 area.

Favorable Type V and Type VI diagenetic facies commonly occur in barrier bar sandbodies within S₁k¹. Type V facies distribute in Wells SH10, SH902H, and SH903H areas. It is inferred that the area of high porosity value to the west of Well SH903H contains Type V facies. Type VI facies dominantly distributes where Wells SH904H and SH901 are located because this area is a fracturedeveloped belt which is favorable for fluid dissolution. The northwest and east parts of the study area which are also fracture-developed have high porosity values and, therefore, are thought to have Type VI diagenetic facies dominantly. Type I diagenetic facies occurs extensively in the areas with alternating beds of sandstone and mudstone. Analysis of the lateral distribution of diagenetic facies comes to the conclusion that favorable diagenetic facies develop extensively to the east of Well SH903H and to the north of Well SH10, which are supposed to be the targets of future exploration and development.

CONCLUSIONS

- (1) The quantitative evaluation of porosity in the Kepingtag Formation diagenetic event supposed that the tightness of sandstone results from compaction, cementation, and hydrocarbon charging (bitumen filling) whereas the improvement of the reservoir properties is controlled by dissolution.
- (2) 6 types of diagenetic facies of Kepingtage sandstone are identified based on this study proposes "cause (type and intensity of diagenesis) + effect (influence on the reservoir)" scenario: tight facies from strong compaction (Type I), tight facies from carbonate cementation (Type II), tight facies from bitumen filling (Type III), porosity preservation facies from weak diagenesis (Type IV), porosity enhancement from medium dissolution (Type V), and porosity generation facies from strong dissolution (Type VI), of which Types I, II, and III are unfavorable whereas Types IV, V, and VI are favorable for effective reservoir development.
- (3) The results of the quantitative evaluation of diagenetic facies show that the distribution of diagenetic facies is mainly

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controlled by sedimentary microfacies, and the favorable diagenetic facies are mainly developed in the middle of underwater distribution channel and barrier bar thick sandbodies. Favorable diagenetic facies are best developed in S_1k^1 sandstone, of which Type V and Type VI, occurring primarily in its top, are dominant. The analysis of the lateral distribution of diagenetic facies in combination with the sedimentary facies and porosity concludes that the areas to the east of Well Sh903H and the north of Well SH10 are favorable for the future exploration and development.

(4) A method for the quantitative identification of diagenetic facies was established, and the well logging parameters reflecting the tight sandstone rocks, sedimentary environment, and physical characteristics were selected for setting the similarity with the corresponding sample. The study adopted "KNN" discriminance to identify the vertical diagenetic facies, and the accuracy reached 78.6%. It shows that the logging information of tight sandstone can effectively reflect the lithology and physical property changes of the rock. Therefore, the quantitative identification of diagenetic facies based on logging parameters can provide reasonable results for the evolution of the tight sandstone reservoirs.

DATA AVAILABILITY STATEMENT

The original contributions presented in the study are included in the article/Supplementary Material, further inquiries can be directed to the corresponding author.

AUTHOR CONTRIBUTIONS

All authors listed have made a substantial, direct, and intellectual contribution to the work and approved it for publication.

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Conflict of Interest: Author ZH was employed by PetroChina Zhejiang Oilfield Company.

The remaining authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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