



Size Class Dependent Relationships between Temperature and Phytoplankton Photosynthesis-Irradiance Parameters in the Atlantic Ocean

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Over the past decade, a number of methods have been developed to estimate size-class primary production from either in situ phytoplankton pigment data or remotely-sensed data. In this context, the first objective of this study was to compare two methods of estimating size class specific (micro-, nano-, and pico-phytoplankton) photosynthesis-irradiance (PE) parameters from pigment data. The second objective was to analyse the relationship between environmental variables (temperature, nitrate and PAR) and PE parameters in the different size-classes. A large dataset was used of simultaneous measurements of the PE parameters (n = 1,260) and phytoplankton pigment markers (n = 2,326), from 3 different institutes. There were no significant differences in mean PE parameters of the different size classes between the chemotaxonomic method of Uitz et al. (2008) and the pigment markers and carbon-to-Chl a ratios method of Sathyendranath et al. (2009). For both methods, mean maximum photosynthetic rates (P_m^B) for micro-phytoplankton were significantly lower than those for pico-phytoplankton and nano-phytoplankton. The mean light limited slope (α^B) for nano-phytoplankton were significantly higher than for the other size taxa. For micro-phytoplankton dominated samples identified using the Sathyendranath et al. (2009) method, both P_m^B and α^B exhibited a significant, positive linear relationship with temperature, whereas for pico-phytoplankton the correlation with temperature was negative. Nano-phytoplankton dominated samples showed a positive correlation between P_m^B and temperature, whereas for α^B and the light saturation parameter (E_k) the correlations were not significant. For the Uitz et al. (2008) method, only micro-phytoplankton P_m^B , pico-phytoplankton α^B , nano- and pico-phytoplankton E_k exhibited significant relationships with temperature. The temperature ranges occupied by the size classes derived using these methods differed. The Uitz et al. (2008) method exhibited a wider temperature range compared to those derived from the Sathyendranath et al. (2009) method. The differences arise from the classification of mixed populations. Based on these patterns, we therefore recommend using the Sathyendranath et al. (2009) method to derive micro-phytoplankton PE

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parameters at sea water temperatures up to 8°C during monospecific blooms and the Uitz et al. (2008) method to derive *PE* parameters of mixed populations over the temperature range from 8 to 18°C. Both methods exhibited similar relationships between pico-phytoplankton *PE* parameters and temperatures >18°C.

Keywords: Atlantic Ocean, phytoplankton size, photosynthesis parameters, primary production, temperature

INTRODUCTION

Increased interest in predicting future effects of climate change on marine ecosystems has led to a concerted effort to understand how large-scale patterns in ocean productivity vary in response to changing environmental properties, especially to sea-surface temperature, which is routinely monitored from satellite. A range of models have been developed to estimate primary production from space (Platt and Sathyendranath, 1988; Platt et al., 1995; Antoine et al., 1996; Behrenfeld and Falkowski, 1997; Behrenfeld et al., 2005; Smyth et al., 2005; Uitz et al., 2008, 2012; Westberry et al., 2008) and several inter-comparisons of primary production algorithms have been conducted (Carr et al., 2006; Friedrichs et al., 2009; Tilstone et al., 2009; Saba et al., 2010). One principal conclusion from these studies was that a better understanding of the temperature regulation and parameterisation of photosynthesis is required to improve the accuracy of remotely-sensed models of marine primary production (Carr et al., 2006).

Relationship between Phytoplankton Photosynthesis and Temperature

To estimate primary productivity at basin or global scales, it is necessary to understand how the PE parameters; the light limited slope (α^B), the maximum photosynthetic rate (P_m^B), and the light saturation parameter (E_k) or associated parameters such as the optimum photosynthetic rate (P^b_{opt}) vary as a function of environmental variables that can be measured from remote sensing (Behrenfeld and Falkowski, 1997; Bouman et al., 2005; Platt et al., 2005). For the Scotian Shelf, Bouman et al. (2003) found that 65% of the variation in P_m^B could be described by temperature alone. The correlation between temperature and PE parameters is also likely to indicate relationships with other variables such as nitrate (Bouman et al., 2005; Marañon et al., 2012; Saux-Picart et al., 2014). It has already been documented that below 20°C, increasing temperature results in an increase in P_m^B (Raven and Geider, 1988). At temperatures above 20°C, commonly associated with stratified oligotrophic conditions and high surface irradiance, the maximum rate of carbon fixation is often reduced, which may be explained in part by an increased investment of resources in photo-protection, repair of photo-damage (Raven, 2011), and to endure nutrient limitation (Behrenfeld and Falkowski, 1997).

Phytoplankton Cell Size and Photosynthesis

All aspects of the life of phytoplankton are influenced by its size (Chisholm, 1992), which affects it's growth, metabolism and access to resources, and in turn shapes it's ecological

niche (Lichtman and Klausmeier, 2008). Total Chl a tends to increase with increasing cell size (Chisholm, 1992). In turn cell size is also an important source of variability in PE parameters. Historic evidence has shown that there are physiological differences between large and small phytoplankton cells in terms of Michaelis-Menten constants and sinking rates, which have implications for photosynthetic efficiency and carbon export from the pelagic zone (Paasche, 1967, 1968; Eppley et al., 1969; Smavda, 1970). Allometric theory states that metabolic rates scale with body size following Kleiber's law whereby the metabolic rate is a 3/4 function of body size (Peters, 1983; Brown et al., 2000), which applies to photosynthesis (Niklas and Enquist, 2001). A number of studies questioned the allometric 3/4 rule, since it does not account for the size dependence of light absorption and the size scaling of photosynthesis (Enquist et al., 1998; Finkel and Irwin, 2000; Niklas and Enquist, 2001). Eppley and Sloan (1965) similarly showed a dependency of respiration rate on cell size. The rate of respiration scales roughly as the -0.25 power of cell biomass (Laws, 1975; Banse, 1976), though significant deviations have been observed even within specific phytoplankton species (Raven, 1986). Similarly Chisholm (1992) found that this relationship is weak and that ecological differences between taxonomic groups mask the influence of size on growth. Recent studies showed that a single, universal scaling rule cannot predict the metabolism-size relationship in all photosynthetic organisms, since changes in taxonomic composition along the size spectrum have a significant effect on photosynthesis (Marañon et al., 2007).

The absorption cross-section also decreases with cell size because self-shading increases as cell size increases (Geider et al., 1986). Acclimation of phytoplankton to temperature and light conditions are important parameters in shaping the size structure of the phytoplankton community (MacIntyre et al., 2002; Thomas et al., 2012). Morel and Bricaud (1981) found that carbon-specific photosynthetic rates decrease with cell size in both light saturated and light limited cells. Photo-acclimation involves compensatory changes in the photosynthetic apparatus including cell volume, pigment composition and concentration, and the number and density of thylakoid membranes (Falkowski, 1980), which occurs over a cell generation of a few days (Falkowski and LaRoche, 1991). Typically there is a decrease in cellular photosynthetic pigment content across all taxa (MacIntyre et al., 2002), which is related to the amount of accessory pigment within the photosynthetic antenna or the amount of chlorophyll per reaction centre. Models of photoacclimation indicate that changes in cellular pigment content balance the energy captured by light absorption with the energy demand for growth (Geider et al., 1998), which can be related to size (Roy et al., 2011).

Phytoplankton Cell Size, Turbulence, and Nutrient Assimilation

Cell size has important implications for resource acquisition, such as nutrient uptake (Sournia, 1982) and light harvesting (Duysens, 1956; Morel and Bricaud, 1981). Large phytoplankton, especially diatoms, can fix carbon and assimilate nutrients rapidly (Pahlow et al., 1997), which are then removed from the photic zone as they sink to depth (Riegman et al., 1993). They tend to thrive in regions rich in nitrate. Smaller cells (pico-phytoplankton) fix less carbon (Chisholm, 1992) and sink slowly, resulting in potential remineralisation via the microbial loop (Legendre and Rassoulzadegan, 1995). They tend to inhabit nutrient poor, oligotrophic regions (Munk and Riley, 1974; Malone, 1977). Water column characteristics and pelagic food web structure are tightly coupled and mediated by the effect of turbulence, which in turn can affect phytoplankton cell size (Kiørboe, 1993). Micro-phytoplankton dominates in eutrophic, turbulent and partially mixed waters, which enhances nutrient uptake and sedimentation. Small phytoplankton is generally more abundant in vertically stratified oligotrophic waters (Margalef, 1978). Turbulence also influences the light environment either by increasing the light intensity variability or by decreasing the average light intensity, which microphytoplankton, particularly diatoms, are potentially well adapted to (Kiørboe, 1993; Lavaud et al., 2007).

Phytoplankton Cell Size, Growth, and Grazing

The dominance of larger phytoplankton in nutrient-rich areas is also attributed to the stronger grazing pressure suffered by smaller cells (Kiørboe, 1993; Irigoen et al., 2005). This may be related to the size scaling exponent of carbon fixation, whereby the larger micro-phytoplankton (>20 μ m) grows faster than medium size nano-phytoplankton (2–20 μ m) and the smaller pico-phytoplankton (0.2–2 μ m) under high-resource conditions (Niklas and Enquist, 2001). The relationship between cell size, phytoplankton abundance and carbon fixation suggests that the ecosystem size abundance distribution is the result of size-scaling metabolic rates, which when normalised to cell abundance, remains constant over the range in size (Huete-Ortega et al., 2012).

There seems to be an almost constant predator-prey size ratio in plankton food chains such that small predators graze small phytoplankton and large predators graze large phytoplankton (Berggreen et al., 1988). Some studies have suggested however, that predation pressure and growth rates decrease with size (Platt, 1985). Thus, despite the lower relative growth rate, nutrient uptake and light harvesting capabilities of large compared to small phytoplankton, their risk of being eaten is much lower (Kiørboe, 1993).

Phytoplankton Size from Space

Recently, there have been an increasing number of studies that have retrieved the size structure of phytoplankton from remotesensing (Brewin et al., 2014). There is also an impetus to develop new algorithms that divide total marine primary production

into different phytoplankton size classes (e.g., Brewin et al., 2017). Thorough validation and characterisation of some of these remote sensing models has been performed, and have an error of 0.27-0.63, depending on the size-class retrieved (Brewin et al., 2017). Our understanding of how the photophysiological parameters vary within size-classes or taxa over basin scales remains poor and is still limited by the availability of in situ observations. Coincident in situ measurements of PE parameters, phytoplankton taxonomic and/or functional groups are required for this and for accurate parameterisation of size-class models of primary production (Uitz et al., 2008). There are few direct measurements of size-fractionated PE parameters (e.g., Tilstone et al., 2017), but more coincident PE parameters and HPLC pigment data (e.g., Uitz et al., 2008). Since different phytoplankton taxa contain different diagnostic pigments, information on pigment composition can be used to estimate the taxonomic composition of a sample (Everitt et al., 1990; Letelier et al., 1993; Bidigare and Ondrusek, 1996; Mackey et al., 1996; Vidussi et al., 2001; Uitz et al., 2006, 2008; Van den Meersche et al., 2008; Sathyendranath et al., 2009). Two methods have been developed methods to derive PE parameters of different size classes from pigment markers. These are either based on empirical approaches between size class marker pigments, the associated photo-physiological parameters and its dependence on irradiance in the water column (Uitz et al., 2008), or on least-squares regression between particulate carbon and Chl a for different phytoplankton groups identified using pigment markers (Sathyendranath et al., 2009). These methods have not yet been compared to evaluate any potential differences between them.

The objectives of this paper are therefore to: (1) Compare two different pigment-based methods at estimating size class specific PE parameters, and then (2) Analyse the relationship between environmental variables and photosynthetic rates in different size-classes using these methods. The resulting relationships are discussed in the context of choosing appropriate models to estimate phytoplankton size-class primary production using remotely-sensed data.

METHODS

Study Area and Samples

Phytoplankton pigment and *PE* data were collected from 47 cruises in the Atlantic Ocean led by three principal institutes: the Bedford Institute of Oceanography (BIO) (Bouman et al., 2005; Sathyendranath et al., 2009), the Laboratoire d'Océanographie de Villefranche-sur-Mer (LOV) (Babin et al., 1994) and the Plymouth Marine Laboratory (PML) (Tilstone et al., 2003). Data from BIO were collected from 1997 to 2001 along the Scotian Shelf and from the NW sub-tropical Atlantic in March 1999; data from LOV were from the tropical NE Atlantic in June 1992 and data from PML were from AMT6 in May-June 1998 and the Irish Sea in May, June and July 2001. There were no significant differences in nano-phytoplankton biomass between laboratories [$t_{(2, 11)} = 2.65$, p = 0.022]. There was also no significant difference in pico-phytoplankton biomass between BIO and PML [$t_{(2, 6)} = 4.13$, p = 0.065]. There was, however a

significant difference in micro-phytoplankton between BIO and PML [$t_{(1, 63)} = 4.06$, p = 0.0001]. None of the LOV samples were dominated by micro- or pico-phytoplankton so a direct comparison with this laboratory was not possible.

Water samples were collected from 3 to 8 depths in the upper 200 m of the water column (Figure 1). Surface irradiance values were estimated using MODIS-Aqua monthly climatology of photosynthetically-active radiation (PAR) (NASA Ocean Colour Level 3 products; http://oceancolor.gsfc.nasa.gov). Mixed-layer depth (z_m) was calculated using climatological data from the World Ocean Atlas (Levitus et al., 2009). Temperature, salinity and potential density data were estimated using the algorithms given in Jackett et al. (2006) since profiles of Conductivity Temperature Depth (CTD) were not available at all stations. To estimate the euphotic depth (z_e) , which we operationally define as the depth at which the light intensity is 1% of surface irradiance, vertical attenuation coefficients of 0.016 m² (mg Chl a)⁻¹ (Schanz et al., 1997) for Chl *a* and 0.043 m⁻¹ for seawater were used (Platt et al., 2003). The mean mixed-layer PAR was also calculated using estimates of light attenuation coefficients from Chl *a* concentrations, z_m and surface PAR following the methods of Platt et al. (1988) and Babin et al. (1996).

For the analysis of nutrients, seawater samples were collected directly from the CTD rosette, stored frozen until analysis and then equilibrated to room temperature prior to analysis. BIO measured nitrate, phosphate and silicate using an autoanalyser (Technicon II or Alphkem RFA300) following the methods of Strain and Clement (1996). LOV measured nitrate, nitrite and phosphate using a Technicon Autoanalyzer using the methods of (Wood et al., 1967) and Murphy and Ryley (1962). For PML, dissolved inorganic nutrient concentrations were determined by gas segmented flow colorimetric analysis using a Bran and Luebbe AutoAnalyser (model AA3). Dissolved nitrate and nitrite were determined by the spectrophotometric methods described by Brewer and Riley (1965). The analysis of soluble reactive phosphorus (SRP) concentration was based on the method described by Zhang and Chi (2002). Dissolved silicate (DSi) was determined using standard colorimetric methods (Kirkwood, 1989).

Phytoplankton Pigments

Chl *a* was measured using a Turner Designs fluorometer before and after acidification. BIO used the method of Holm-Hansen et al. (1965) (N = 1,873), LOV used the method of Yentsch and Menzel (1963) adapted by Holm-Hansen and Riemann (1978) (N = 162), and PML used the non-acidification method of Welschmeyer (1994) (N = 227).

Total Chl *a* (monovinyl + divinyl Chl *a*) and accessory pigments were measured on a subset of samples (N = 2,361) using high performance liquid chromatography (HPLC) following the procedure described by Head and Horne (1993) for BIO, Claustre and Marty (1995) for LOV and Barlow et al. (1997) for PML. Water samples were filtered onto GF/F filters before being flash frozen in liquid nitrogen and stored at -80° C until analysis. Pigment standards from Sigma Chemical Company, were used for the identification of Chl *a*, Chl *b*, and β -carotene. Chl *c*, 19'-butanoyloxyfucoxanthin, fucoxanthin,





19'-hexanoyloxyfucoxanthin and diadinoxanthin were identified using standards from Dr R. Bidigare (Claustre and Marty, 1995; Stuart et al., 2000) for BIO and LOV, and from DHI Denmark (Gibb et al., 2000) for PML. Further details of the HPLC methods used by each laboratory are given in **Table 1**.

Photosynthesis-Irradiance (PE) Parameters

The protocols used for the determination of *PE* parameters are described in Platt et al. (1982) for the BIO data, Babin et al. (1994) for the LOV data, Tilstone et al. (2003) for the PML data. In brief, seawater samples were collected from 3 to 8 depths based on the vertical structure in the in vivo fluorescence profile. During the 14C incubations the temperature of the photosynthetron was closely matched with ambient *in situ* temperature. For surface samples, the ships surface underway seawater supply was used to maintain the incubations at sea surface temperature. For deeper samples, each photosynthetron was connected to a temperature control unit which was used to maintain the incubations at ambient temperature.

All labs added 185 to 370 kBq (5 to 10 μ Ci) of ¹⁴C-labelled bicarbonate to the PE bottles. Ambient PAR was monitored over a 2-3 hr period prior to collecting the water samples for 14C incorporation and the photosynthetron irradiance was matched to the average PAR over this period. Platt et al. (1982) irradiated the samples using 30 different light levels and incubated for 2-3 h. Tilstone et al. (2003) used 15 light levels, whereas Babin et al. (1994) used 12 and both methods incubated samples for 1-2 h. All laboratories used tungsten halogen lamps for incubation and ¹⁴C rates were spectrally corrected based on phytoplankton absorption coefficient and the spectrum of the lamp (Kyewalyanga et al., 1997). It has been shown that short incubation periods (1-3 hrs) minimise bottle effects and do not result in nutrient limitation (Falkowski, 1981). Following incubation, all labs filtered the samples through 25 mm glass fibre filters (Whatman GF/F) at a vacuum pressure of <20 cm Hg. The filters were then exposed to 37% fuming HCl fumes for <12 hr to remove any inorganic carbon present, immersed in scintillation cocktail and disintegration time per minute of ¹⁴C was counted on a liquid scintillation counter. The PE parameters, P_m^B [in mg C (mg Chl-a) ⁻¹] and α^B [in mg C (mg Chl-a) ⁻¹ h ⁻¹ (µmol photons m ⁻² s ⁻¹) ⁻¹] were estimated by fitting the data to the model of Platt et al. (1980) (N = 882) or Jassby and Platt (1976) (N = 194) before being normalised to Chl a. The choice of the model was determined by inspection of each PE curve to evaluate whether there was photo-inhibition or not. The maximum quantum yield (ϕ_m) (mol C mol quanta⁻¹) was calculated as:

$$\phi_m = 0.0231^* \text{Chl} a^* \alpha^B{}_{\text{PUR}} \tag{1}$$

Where PUR ($\mu E m^{-2} s^{-1}$) is the phytoplankton useable radiation and the constant 0.0231 in the numerator converts grams to moles and hours to seconds.

The total absorption coefficient of phytoplankton was determined by filtering between 0.25 and 2 L of seawater onto 25 mm GF/F filters and the absorbance of the particulate material was measured from 350 to 750 nm at a 1 nm bandwidth using dual beam Perkin Elmer Lambda spectrophotometers. Measurements were made of total particulate absorption $[a_{part}(\lambda)]$ and $a_{NAP}(\lambda)$ before and after pigment extraction with methanol following the methods of Kishino et al. (1985). The pathlength amplification correction of Kyewalyanga et al. (1997) was used and $a_{ph}(\lambda)$ was derived from the difference between $a_{part}(\lambda)$ and $a_{NAP}(\lambda)$.

Pigment-Based Classification of Phytoplankton Size-Classes

The Uitz et al. (2008) Method

The pigment-based method of Uitz et al. (2008) divides the total Chl *a* biomass into three size-classes (micro-phytoplankton, >20 μ m; nano-phytoplankton, 2–20 μ m; pico-phytoplankton, <2 μ m) using seven diagnostic pigments (fucoxanthin (fuc), peridinin (per), alloxanthin (allo) 19'-butanoyloxyfucoxanthin (19'-but), 19'-hexanoyloxyfucoxanthin (19'-hex), zeaxanthin (zea) and the monovinyl chlorophyll *b* and divinyl chlorophyll *b* [Chl *b*)]. A weighted sum of these diagnostic pigments (wDP) was calculated as follows:

wDP =
$$1.41[fuc] + 1.41[per] + 0.6[allo] + 0.35[19' - but]$$

+ $1.27[19' - hex] + 0.86[zea] + 1.011[Chl b]$ (2)

Using the ratio of the weighted sum of group-specific diagnostic pigments to wDP, the proportion of total biomass in the micro-phytoplankton (f_{micro}), nano-phytoplankton (f_{nano}) and pico-phytoplankton (f_{pico}) size classes were calculated using the following set of equations:

$$f_{micro} = (1.41[fuc] + 1.41[per])/wDP,$$
 (3)

$$f_{nano} = (0.6[allo] + 0.35[19' - but])$$

$$+1.27[19' - hex])/wDP$$
, and (4)

$$f_{pico} = (0.86[\text{zea}] + 1.01[\text{chl} - b])/\text{wDP}.$$
 (5)

In this study samples were defined as being dominated by a single size-class if more than 70% of the total biomass was assigned to the size-class.

The Sathyendranath et al. (2009) Method

Sathvendranath et al. (2009) identify samples dominated by any one of six phytoplankton groups (diatoms, dinoflagellates, prymnesiophytes, green algae, Prochlorococcus, other cyanobacteria) using the following diagnostic pigments: fuc, chl-c, diad, per, allo, 19'-but, 19'-hex, DV Chl a, monovinyl Chl b (MV Chl b), DV Chl b, and zea. Instead of trying to apportion Chl a to each size-class, ratios of diagnostic pigments with respect to Chl a were used to classify samples as either being dominated by a particular taxonomic group or containing a mixed assemblage of cells (Sathyendranath et al., 2009). In order to compare directly with the results obtained from the Uitz et al. (2008) method, these taxonomic classes were assigned to the three size-classes (diatoms and dinoflagellates were treated as micro-phytoplankton, prymnesiophytes and green algae as nano-phytoplankton, Prochlorococcus and other cyanobacteria as pico-phytoplankton), the distribution of each

TABLE 1 | Summary of HPLC methods.

Method details	BIO	LOV	PML Barlow et al., 2007		
Reference	Head and Horne, 1993; Hooker et al., 2005	Neveux and Lantoine, 1993; Hooker et al., 2005			
Equipment	Beckman System Gold HPLC system with a 126 solvent module with binary pump and a 168 UV-VIS diode array detector	Agilent Technologies ChemStation of LC software(A.09.03) with degasser, binary pump and DAD (Agilent Technologies 1100)	Thermo Electron UV6000 diode array absorbance detector		
Column	3-µm Ultrasphere-XL ODS column	$3-\mu m$ Hypersil-MOS C ₈ column	3 - μ m Hypersil MOS2 C ₈ column		
Auto-sampler	Manual injection	Thermoquest Autosampler (AS 3000)	Thermo-Electron S3000		
Extraction Solvent	100% acetone	100% methanol	100% acetone		
Extraction technique	Grinding with Teflon pestle	Ultra-sonification	Ultra-sonification		
Extraction time	1 min using glass tube and motorized grinder and centrifuged for 20 min at 12,000 g	0.5 h in freezer then macerated using ultrasonic probe for 10 s	Ultra-sonification		
Pigment Detection	430 nm using a Beckman 168 UV-VIS diode array detector and a Beckman 157 fluorescence detector fitted with a broad band excitation filter (585–650 nm) and a red-sensitive extended wavelength photomultiplier tube	Not reported	440 and 665 nm and identified by retention time and on-line diode array spectra		
Pigment standards	Chlorophylls <i>a</i> and <i>b</i> , and BB-carotene were purchased by Sigma-Aldrich. DHI for other pigment standards	Eight pigment standard provided by DHI	Chlorophyll a standard and trans-b apo-80 carotenal internal standard (Fluka) from Sigma-Aldrich Ltd. DHI for other pigment standards		
Filters	Whatman GF/F	Whatman GF/F	Whatman GF/F		

is given in **Figure 1**. Uitz et al. (2008) assigned all biomass associated with Chl *b* to the pico-phytoplankton size-class due to its presence in classes such as *Micromonas* (Not et al., 2004) and *Prochlorococcus* (Chisholm, 1992). For this study however, due to the importance of Chl *b* containing nano-flagellates in the North Atlantic (Wright et al., 2010; Eker-Develi et al., 2012), samples found to consist predominantly of green flagellates using the Sathyendranath et al. (2009) were categorised as nano-phytoplankton.

Statistical Analyses

All data were analysed using CRAN R statistical software, with package "reshape" 0.8.4 used to process the data (R Core Team, 2015). Principal component analyses (PCA) based on correlation matrix were employed to assess the variation in P_m^B and α^B for each phytoplankton size class in relation to other environmental variables to assess the forcing functions on size class photosynthetic parameters. PCA was performed using the prcomp command within "stats" package 2.5.1. Data were plotted using package "ggplot2" version 0.8.9. Significant differences between size classes and methods were analysed using the Wilcoxon Rank Sum test and ANOVA functions within the "stats" package 2.5.1. The ANOVA results are given as $W_{1,df} = x$, p = y where W is the Wilcoxon Rank Sum, the sub-script numbers denote the degrees of freedom and p is the ANOVA critical significance value.

RESULTS

Relationship between Phytoplankton Photosynthesis and Environmental Variables

Using the Uitz et al. (2008) method, stations dominated by micro-phytoplankton occurred in the NW Atlantic on the Scotian Shelf, in the Irish Sea and at some locations along the Benguela upwelling zone (**Figure 1A**). There were few stations entirely dominated by nano-phytoplankton, but samples with >50% nano-phytoplankton occurred throughout the data set (**Figure 1B**). Samples dominated by pico-phytoplankton occurred in the Gulf of Guinea and in the NW sub-tropical Atlantic (**Figure 1C**). A climatological frequency distribution of data over the seasons is given in **Figure 2**. More than 30 stations were sampled throughout the year except for the winter period from January to February and summer period from mid-July to early August. The highest number of stations sampled was during late summer, from September-October, and during spring, in April.

To validate the micro-phytoplankton biomass from the Uitz et al. (2008) method, values of f_{micro} were compared with the fractional contribution of micro-phytoplankton from microscopy. Total carbon biomass was calculated from cell counts using carbon conversion factors for the AMT6 cruise, which covered a range of oceanographic conditions across a latitudinal gradient. A significant correlation between f_{micro} using the two approaches was observed (p = 0.0005, $r^2 = 0.51$; Figure 3), which suggests that for a variety of oceanic regimes,





the Uitz et al. (2008) method serves as robust chemotaxonomic indicator of micro-phytoplankton biomass.

The possible dependence of phytoplankton photosynthesis upon biological environmental variables was examined using PCA (**Figure 4**), which allows us to visualise the similarities and differences between samples and correlations between variables. The data matrix consisted of 373 observations of 9 variables (environmental variables: temperature, depth and PAR; biological variables: total Chl *a*, the proportion of size-specific Chl *a*, P_m^B and α^B derived using the Sathyendranath et al. method). The data were mean centred and normalised to one standard deviation so that data with different dispersion and units could be compared. The first two eigenvalues of the data matrix account for over 65% of the variance; the x-axis accounts for 42.97% and the y-axis accounts for 22.68% of the variability in the data (**Figure 4**). We therefore limited our analysis to the first factorial plane. The individual samples are coloured by the temperature range occupied by micro-, nano- and pico-phytoplankton (**Figure 4**). Nutrients were not included in this analysis due to the lag phase between nutrient concentration



and uptake which can bias the relationship between cell size and nutrients.

The main variables contributing to the formation of the first principal component (PC1) are the fraction of microphytoplankton Chl *a* ($\%_{micro}$; $r^2 = 0.47$) and total Chl *a* ($r^2 = 0.34$; positive eigenvalues), temperature ($r^2 = 0.43$), P_m^B ($r^2 = 0.35$) and the fraction of pico- ($\%_{pico}$; $r^2 = 0.34$), nano- phytoplankton ($\%_{nano}$; $r^2 = 0.36$; negative eigenvalues; **Figure 4**). This indicates the association of high biomass regions with micro-phytoplankton dominance and that incrementing temperature is coupled with increases in pico-, nano-phytoplankton biomass and P_m^B . The main variables that form the second principle component are PAR ($r^2 = 0.55$), $\%_{pico}$ ($r^2 = 0.39$) and depth ($r^2 = 0.25$) with positive eigenvalues, and α^B ($r^2 = 0.45$), Chl *a* ($r^2 = 0.36$), P_m^B ($r^2 = 0.30$) and $\%_{nano}$ ($r^2 = 0.24$) with negative eigenvalues. The patterns in PC2 explain the vertical variability of light and $\%_{pico}$ with depth and the increase in photosynthetic parameters with increases in $\%_{nano}$ and Chl *a*.

Variability in Photosynthesis-Irradiance Parameters

From the in situ data, P_m^B varied from 0.33 to 20 mg C (mg Chl a)⁻¹ h⁻¹, with a mean of 3.5 mg C (mg Chl a)⁻¹ h⁻¹, α^B varied between 0.002 and 0.59 (mg chl)⁻¹ h⁻¹ (µmol quanta m⁻² s⁻¹)⁻¹ with a mean of 0.033 (mg chl)⁻¹ h⁻¹ (µmol quanta m⁻² s⁻¹)⁻¹, the range in E_k was 12 to 1,300 µmol photons m² s⁻¹, with a mean of 128 µmol photons m² s⁻¹, ϕ_m from 0.0004 to 0.123 mol C (mol photons)⁻¹ with a mean of 0.033 mol C (mol

photons)⁻¹ and a_{ph}^* from 0.059 to 0.074 m² mg⁻¹ with a mean of 0.0165 m² mg⁻¹ (**Figure 4**, **Table 2**).

There were no significant difference in mean P_m^B obtained using the methods of Uitz et al. (2008) and Sathyendranath et al. (2009) for nano- and pico-phytoplankton, however there were significant differences between methods for microphytoplankton (**Figure 5A**, **Table 2**). Using the Uitz et al. (2008) classification method, mean P_m^B values for micro-phytoplankton were significantly lower than those for pico- (W = 5597, p< 0.0001) and nano-phytoplankton (W = 4149, p < 0.0001). Micro-phytoplankton mean P_m^B estimates derived from the Sathyendranath et al. (2009) method were also significantly lower than that for both pico-phytoplankton (W = 851, p = 0.008) and nano-phytoplankton (W = 1061, p < 0.0001). There was no significant difference in mean P_m^B between pico- and nanophytoplankton using either classification method (**Figure 5A**, **Table 2**).

There were also no significant differences in mean α^B for nano- and pico-phytoplankton between the two methods, but for micro-phytoplankton there was a significant difference between the methods (**Figure 5B**, **Table 2**). The mean α^B was significantly higher for nano-phytoplankton compared to micro-phytoplankton, using both the Uitz et al. (2008) approach (W = 461, p < 0.0001) and the Sathyendranath et al. (2009) method (W = 93, p < 0.0001). The mean value of α^B was also significantly higher for nano-phytoplankton compared to pico-phytoplankton for both methods (Uitz et al., 2008: W = 21, p < 0.0001; Sathyendranath et al., 2009: W = 41, p = 0.003). There were no significant differences in mean α^B between pico-

TABLE 2 | Mean and standard deviation for the PE parameters; the maximum photosynthetic rate (P_m^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (mg Chl-a)⁻¹], the light-limited photosynthetic rate (α^B) [mg C (m

	Size-class	Micro		Nano		Pico	
		Mean	SD	Mean	SD	Mean	SD
P _m ^B	Uitz et al., 2008	2.3	1.5	4.34	2.4	3.25	1.5
	Sathyendranath et al., 2009	1.97	1.2	4.15	2.0	4.1	2.7
	Differences between	<i>t</i> _{1,138} =	2.17,	$t_{1,25} =$	0.23,	t _{1,20} =	-0.22
	methods.	<i>p</i> =	0.032	$\rho =$	0.815	$\rho =$	0.825
α ^B	Uitz et al., 2008	0.022	0.018	0.052	0.021	0.028	0.013
	Sathyendranath et al., 2009	0.017	0.01	0.043	0.018	0.021	0.026
	Differences between	t _{1,240} =	-4.18,	$t_{1,26} =$	1.16,	$t_{1,20} =$	-0.35
	methods.	<i>p</i> <	0.0001	$\rho =$	0.255	$\rho =$	0.726
E _k	Uitz et al., 2008	110.3	50.5	85.9	30.5	275.6	141.8
	Sathyendranath et al., 2009	125.8	47.8	126.3	98.4	291.9	162.0
	Differences between	$t_{1,126} =$	-2.45,	$t_{1,17} =$	-1.51,	$t_{1,28} =$	-0.31
	methods.	<i>p</i> =	0.015	$\rho =$	0.149	$\rho =$	0.761
фm	Uitz et al., 2008	0.033	0.025	0.041	0.016	0.008	0.011
	Sathyendranath et al., 2009	0.035	0.020	0.030	0.022	0.013	0.020
a [*] _{ph}	Uitz et al., 2008	0.017	0.004	0.018	0.012	0.020	0.008
	Sathyendranath et al., 2009	0.016	0.003	0.018	0.006	0.018	0.007
	Differences between .	$t_{1,252} =$	1.12,	$t_{1,38} =$	0.06,	$t_{1,69} =$	0.96,
	methods	$\rho =$	0.269	$\rho =$	0.948	$\rho =$	0.339

Results from Sathyendranath et al. (2009) are converted from taxonomic group to equivalent size-class for ease of comparison. Significant differences between methods for size-class and PE parameter are given in bold. For P_m^{β} , α^{β} , E_k for the Sathyendranath et al. (2009) method N = 356 for Micro-, N = 15 for nano-, N = 15 for pico-phytoplankton; for the Uitz et al. (2008) method for Micro- N = 1,212, for nano- N = 24, for pico-phytoplankton N = 19. For ϕm and a_{ph}^* for the Sathyendranath et al. (2009) method N = 21 for Micro-, N = 15 for nano-, N = 15 for pico-phytoplankton; for the Uitz et al. (2008) method for Micro- N = 97, for nano- N = 24, for pico-phytoplankton N = 29.

and micro-phytoplankton using either classification method (Figure 5B, Table 2).

Similarly there were no significant difference in mean E_k value for nano- and pico-phytoplankton between methods, but there was a significant difference for micro-phytoplankton (**Figure 5C, Table 2**). The mean E_k value for pico-phytoplankton determined using the Sathyendranath et al. (2009) approach was also significantly higher than that for nano-phytoplankton (W = 185, p = 0.003) and micro-phytoplankton (W = 964, p < 0.0001). There was no significant difference in mean E_k between nano- and micro-phytoplankton size classes between methods.

Mean ϕ_m for pico-phytoplankon derived from the Uitz et al. (2008) method was significantly lower than those for both nano-phytoplankton (W = 164, p > 0.0001) and micro-phytoplankton (W = 5214, p > 0.0001; **Figure 5D**, **Table 2**). Similarly, the mean ϕ_m for pico-phytoplankton using the Sathyendranath et al. (2009) approach was also significantly lower than that for nano-phytoplankton (W= 178, p = 0.007) and micro-phytoplankton (W = 995, p < 0.0001). There were no significant differences in ϕ_m between micro- and nano-phytoplankton using either method. Mean a_{ph}^* derived from both methods were almost identical for each size class and there was no significant difference between them (**Table 2**).

Taxa-Specific Relationships between Environmental Variables and Photosynthesis-Irradiance Parameters

Temperature varied between -1.6 and 26.3° C with a mean of 10.7°C. When P_m^B is plotted as a function of temperature (**Figures 6A,E**), the upper boundary of the points exhibited a linear increase with temperature until ~15°C, after which there is a decrease with increasing temperature. The α^B values also displayed a similar increase to 15°C, followed by a decrease to 28°C (**Figures 6B,F**), whereas E_k and a_{ph}^* remained relatively constant to 15°C, and then increased sharply from 15 to 28°C (**Figures 6C-H**), though the increase in a_{ph}^* was far less marked than for E_k (**Figures 6D,H**).

The shape of the relationships between the *PE* parameters and temperature were similar between methods, but there were differences in the temperature ranges associated with each of the size classes derived using the two methods. The temperature



range for micro-phytoplankton dominated samples was greater when using the Uitz et al. (2008) method (-1.6 to 18° C) compared to the Sathyendranath et al. (2009) method (-1.6and 10° C). Using the Uitz et al. (2008) approach, samples dominated by pico-phytoplankton occurred at temperatures as low as 3.7° C, whereas the coldest pico-phytoplankton dominated samples identified by the Sathyendranath et al. (2009) method were at 13° C (**Figure 6**).

For the Uitz et al. (2008) method, there were significant relationships between P_m^B and temperature for micro-phytoplankton (**Figure 6E**, **Table 3**), for α^B and temperature for micro- and pico-phytoplankton (**Figure 6B**, **Table 3**), and for E_k and temperature for micro- and pico-phytoplankton, though for micro- this relationship was also markedly weaker (**Figure 6C**,

Table 3). For the Uitz method, there was also no significant relationship between a_{ph}^* and temperature for each size class (**Figure 6D**, **Table 3**).

Using the Sathyendranath et al. (2009) method, there were significant relationships between P_m^B and temperature for each size-class (Figure 6A). For micro-phytoplankton and nanophytoplankton there were significant positive correlations with temperature (Table 3). Pico-phytoplankton, on the other hand, displayed a significant inverse correlation with temperature, though this was only just within the 5% limit. There were no significant relationships between α^B and temperature for samples dominated by micro- or nano-phytoplankton, whereas for pico-phytoplankton there was a significant inverse relationship with temperature (Figure 6F, Table 3). There was no significant correlation between E_k and temperature for microphytoplankton dominated samples. By contrast, there was a significant positive relationship between E_k and temperature for samples dominated by nano-phytoplankton and picophytoplankton, though for pico- these were markedly weaker (Figure 5G, Table 3). For this method, there was no significant relationship between a_{ph}^* and temperature for each size class (Figure 6H, Table 3).

We also examined the relationship between the photosynthetic parameters P_m^B and α^B derived from the Sathyendranath et al. (2009) and Uitz et al. (2008) methods with nitrate and PAR (**Figure 7**). Nitrate concentrations varied from below the limit of detection to 26.10 μ M, with a mean of 2.15 μ M and PAR varied from 2 to 55 E m⁻² d⁻¹. The distribution of micro-phytoplankton P_m^B and α^B against nitrate and PAR using the Uitz et al. (2008) method, since the latter classified fewer samples in each size class. There were no significant relationships for each parameter from either method for all data and for each size class (**Figure 7**).

DISCUSSION

Comparison of Size-Class Methods

The Uitz et al. (2008) classification method divides the total Chl a biomass of all samples into three size categories according to the relative abundances of specific indicator pigments. The reliability of this method for micro-phytoplankton was confirmed by the correlation between the pigment-based and microscopy-based estimates of the relative proportion of microphytoplankton biomass (Figure 3). The Sathyendranath et al. (2009) approach identifies samples where indicator pigment ratios suggest that they are comprised of a mixed population, and then omits them, leaving only samples dominated by a single taxonomic group. Samples dominated by individual taxonomic groups can then be combined to give a subset of samples dominated by a single size-class (N = 386), and are thus comparable with the output resulting from applying the threshold criteria to the size-class-based approach of Uitz et al. (2008). While the Uitz et al. (2008) method gives a proportion that can be plotted against the microscopy data,



FIGURE 6 | *PE* parameters as a function of temperature: (**A**) is Uitz P_m^B ; (**B**) is Uitz α^B ; (**C**) is Uitz E_k ; (**D**) is Uitz a_{ph}^* ; (**E**) is Sathyendranath P_m^B ; (**F**) is Sathyendranath α_{ph}^B ; (**F**) is Sathyendranath a_{ph}^* . Micro-phytoplankton (>20 µm) is blue circles, nano-phytoplankton (2–20 µm) is green circles, and pico-phytoplankton (<2 µm) is red circles. Black points represent samples classified as "mixed" populations; blue regression line is for micro-, green regression line is for nano- and red regression line is for pico-phytoplankton, black lines are the 10 and 90% quantiles fitted using a running average over all samples, the pink line is the 7th order polynomial relationship from Behrenfeld and Falkowski (1997).

the Sathyendranath et al. (2009) identifies samples dominated by a single size class which cannot be directly compared with the microscopy data, which is a limitation of this method. Full validation of nano- and pico-phytoplankton fractions by each method was not possible due to the lack of coincident flow cytometry data. Further work is therefore required to compare the nano- and pico-phytoplankton size classes against independent *in situ* data.

Another limitation of the Sathyendranath et al. (2009) approach is that the stringent criteria requires a sufficiently large dataset to derive group-specific properties, since many of the samples are excluded on the basis that they are classified as mixed

	Size-class	Micro		Nano		Pico	
		R ²	Р	R ²	Р	R ²	Р
P ^B _m	Uitz et al., 2008	0.15	0.0001				
	Regression Equation	y = 0.15x	+1.68				
	Sathyendranath et al., 2009	0.22	0.0001	0.45	0.003	0.24	0.037
	Regression Equation	y = 0.25x	+1.45	y = 0.22 x	+2	y = 11.7x	-0.33
α ^B	Uitz et al., 2008					0.17	0.0001
	Regression Equation					y = 0.04x	0.0001
	Sathyendranath et al., 2009					0.89	0.001
	Regression Equation					y = 0.15x	-0.005
E _k	Uitz et al., 2008	0.018	0.01			0.29	0.0001
	Regression Equation	y = 2.4x	+103			y = 10.7x	+25
	Sathyendranath et al., 2009			0.41	0.0001	0.3	0.02
	Regression Equation			y = 10x	+23	y = 21x	+200

TABLE 3 | Significant regression equations between size class photosynthetic parameters and temperature.

For the Sathyendranath et al. (2009) method N = 356 for Micro-, N = 15 for nano-, N = 15 for pico-phytoplankton; for the Uitz et al. (2008) method for Micro- N = 1,212, for nano-N = 24, for pico-phytoplankton N = 19.

populations. Many phytoplankton groups, especially in the nanophytoplankton size range, rarely form class-specific blooms (Hirata et al., 2008) and assessing only samples dominated by marker pigments of a single phytoplankton group could result in omission of the majority of phytoplankton communities from the analysis. The Uitz et al. (2008) method is much less restrictive, and when a cut-off point of 70% is used to define samples dominated by a single size-class, more samples are found to have mixed populations compared with the Sathyendranath et al. (2009) approach. This causes the differences in temperature range for each size-class using the two methods. For the Uitz et al. (2008) approach, biomass is assigned to the pico-phytoplankton size-taxa as a result of the presence of either zeaxanthin or Chl b. Some low-temperature samples in our dataset contained high Chl b concentrations, yet little zeaxanthin illustrating why some samples classified as picophytoplankton occur at 3.7°C with this method. While this may indicate the presence of pico-eukaryotes such as Micromonas, it could be also be a result of Chl b containing nano-flagellates being wrongly identified as pico-phytoplankton. The variability between the two methods reflects the way each method treats mixed populations.

Few independent studies have been used to verify the accuracy of these methods. Figueiras et al. (2014) tested the Uitz et al. (2008) method in the upwelling region of the NW Iberian peninsula and concluded that the method is accurate during stratified hydrographic conditions, but when the water column becomes mixed, the method over-estimates primary production. Our analysis suggests that the Sathyendranath et al. (2009) is applicable to conditions when the water column is completed dominated by a particular taxa or size-class and that the Uitz et al. (2008) method is more applicable to retrieving mixed populations. Further testing of the Sathyendranath et al. (2009) method in other shelf sea regions and the Uitz et al. (2008) method in other open ocean systems is required.

Phytoplankton Size Class and Photosynthesis-Irradiance Parameters

The size dependency of P_m^B that we observed (Figure 4A) is consistent with the trends reported in Geider et al. (1997) and Le Quéré et al. (2005), with smaller size-classes having a higher biomass-normalised rates of photosynthesis compared with the larger size-class. By contrast, the global study of Uitz et al. (2008) found that micro-phytoplankton had higher P_m^B . Similarly, a recent study of size fractionated PE parameters in the Atlantic Ocean, Tilstone et al. (2017) also reported that micro-phytoplankton had the highest P_m^B [4.54 mg C (mg Chl a)⁻¹], followed by nano- [4.15 mg C (mg Chl a)⁻¹)] and pico- [2.29 mg C (mg Chl a)⁻¹]. Cermeño et al. (2006) also observed higher biomass-normalised photosynthetic rates for larger cells. These differences may be due, in part, to the range of environmental conditions found in these contrasting studies. In this study, samples dominated by micro-phytoplankton were mostly located in the temperate North West Atlantic (Figure 1A) during the spring bloom (Figure 2), whereas the data used by Uitz et al. (2008) were from the sub-tropics and the Mediterranean. The study of Tilstone et al. (2017) covered both the Atlantic oligotrophic gyres and upwelling zones. The lower photosynthetic rates for micro-phytoplankton observed in this study could be the result of pigment packaging, which arises from self-shading by pigment molecules in the cell (Bidigare and Ondrusek, 1996). The package effect suggests that as the Chl a concentration per cell volume increases, the specific absorption coefficient decreases so that small cells require lower intracellular concentrations of Chl a compared to large cells to harvest equivalent quantities of light (Morel and Bricaud, 1981). Lower pigment packaging occurs due to a decrease in the pigment concentration per cell volume, as a result of either a decrease in cell size. For cells of a constant size, a decrease in internal pigment concentration can occur (Cleveland, 1995). Under this scenario, the carbon-to-chlorophyll ratio decreases as cell size increases.



FIGURE 7 | *PE* parameters as a function of nitrate an PAR: (**A**) is Sathyendranath P_m^B against nitrate; (**B**) is Uitz P_m^B against nitrate; (**C**) is Sathyendranath α^B against nitrate; (**D**) is Uitz α^B against nitrate; (**E**) is Sathyendranath P_m^B against PAR; (**F**) is Uitz P_m^B against PAR; (**G**) is Sathyendranath α^B against PAR; (**H**) is Uitz α^B against PAR; (**H**)

When measuring biomass in carbon units rather Chl a, the carbon-specific photosynthetic rates of small cells will decrease and those of large cells will increase.

The trends in Tilstone et al. (2017) are based on sizefractionated filtration whereas in this study we derived sizeclasses from phytoplankton pigment markers. Errors in the HPLC method may arise from the fact that taxonomic groups identified by specific diagnostic pigments have a large range in size and may therefore be classified across size classes (Brewin et al., 2014). Certain taxonomic groups can exhibit a high degree of variability in pigment concentration and composition. There can also be light-dependent variations in pigments filtration and HPLC methods (Vidussi et al., 2001). Differences in HPLC methods used across data sets may further compound these errors. In this study, differences in pigment concentrations arising from methodological differences in solvent, extraction procedures (see Table 1) could also contribute to differences between phytoplankton characterisation in the mesotrophic and eutrophic North West Atlantic (BIO), the sub-tropical Atlantic (LOV) and the oligotrophic Atlantic (PML).

A size-taxa dependent trend was also observed for α^B ; mean α^B for nano-phytoplankton dominated samples were significantly higher than for micro-phytoplankton (Figure 5B). This was observed in ϕ_m using the Uitz et al. (2008) method, and reflects an increase in light harvesting efficiency and possibly higher nutrient availability (Morel and Bricaud, 1981; Howarth et al., 1988). Nano-phytoplankton is known to form a large proportion of the biomass at the DCM where irradiance levels are low and there is adaptation to maximise the photosynthetic rate through increases in C: Chl a ratios or changes in pigment composition (Bouman et al., 2000; Gibb et al., 2001). From our data, nano-phytoplankton dominated samples in the surface layer have a similar capacity to maximise α^B at mid to high PAR (Figure 7D). Due to their small size and higher specific absorption coefficient, we may expect pico- to have higher α^B compared to nano-phytoplankton. We did not observe this however, though the sample size for pico-phytoplankton was small (Figures 7D,H).

 E_k is defined as P_m^B/α^B , and therefore size-class specific trends in E_k can be driven either by changes in P_m^B or α^B . Similar to Behrenfeld et al. (2004) and Huot et al. (2013), no significant difference was seen between E_k for micro- and nanophytoplankton due to a co-variance of P_m^B and α^B . Thus, whilst we observed nano-phytoplankton P_m^B to be higher than that of micro-phytoplankton, a corresponding increase in α^B resulted in no change in E_k below 15°C. Pico-phytoplankton however had a high mean P_m^B and low α^B resulting in higher mean E_k , reflecting their dominance in the high irradiance environments of the NW sub-tropics and Gulf of Guinea (**Figure 1C**). By contrast, we found no significant difference in a_{ph}^* between size classes (**Figure 5D**).

Environmental Variables, Phytoplankton Size Class and Photosynthesis-Irradiance Parameters

The objective of comparing these pigment based methods of classifying size-class *PE* parameters was to assess the range and

trend in these parameters against environmental variables. Of the environmental variables tested (temperature, nutrients, PAR), we found significant correlations between micro-, nano- and pico-phytoplankton photosynthetic parameters and temperature (**Figure 6**). For nitrate and PAR, correlations with size-class photosynthetic parameters were not significant (**Figure 7**).

Temperature is a major driver of successional changes in phytoplankton, both through direct effects on population growth and indirect effects through changes in water column stability and predator population growth (Lichtman et al., 2010). Many studies have shown that temperature is a strong predictor of phytoplankton cell size in marine ecosystems (Marañon et al., 2001; Bouman et al., 2003) such that the average cell size within phytoplankton communities tends to decreases with increasing temperature, which we also observed in this study (Figure 6). From our dataset, the relationships between temperature and the photosynthetic parameters P_m^B and α^B exhibited a clear envelope, with samples dominated by a specific phytoplankton size-class with a slightly different temperature range depending on the classification method used (Figures 6A,D). For micro- and nano-phytoplankton P_m^B derived from the Sathyendranath et al. (2009) method, the regressions with temperature were highly correlated (Figures 6A,E). Increasing temperature resulted in an increase in P_m^B up to 18°C, following standard reaction kinetics and possible enzymatic effects on the dark reactions of photosynthesis (Geider et al., 1998). For samples dominated by pico-phytoplankton there was a negative relationship between P_m^B and temperature, though this was based on few data. These relationships are similar to those given in Thomas et al. (2012), in which the optimum temperature for growth decreases above 25°C in sub-tropical regions. Under future warming scenarios, this could result in a reduction in phytoplankton diversity and productivity in sub-tropical regions. The low values of P_m^B at very high temperatures are likely to be the combined effects of extreme nutrient limitation (Geider et al., 1997), coupled with increased energetic costs of photo-protection under high irradiance in stratified oligotrophic waters (Raven, 2011). This relationship contrasts that observed in other studies where picophytoplankton and temperature are positively correlated (e.g. Mackey et al., 2013).

In our data, when temperature was high, α^B was low (**Figure 6B**). In this scenario irradiance would also potentially be high suggesting high light adaptation at high temperatures and/or photo-protection. At lower temperatures micro-phytoplankton dominated and α^B was also low, which may be due to the decreased efficiency of light absorption by phytoplankton caused by increased pigment packaging or the negative relationship between ϕ_m and nutrient availability, which can also occur at low temperature (Platt et al., 1992; Sosik, 1996).

 E_k exhibited temperature-independent variation to ~20°C, beyond which there was a linear increase in E_k with increasing temperature (**Figure 6C,F**) possibly as a consequence of thermal stratification. This pattern has been observed in previous studies (Marañon and Holligan, 1999; Behrenfeld et al., 2004), where the highest values of E_k were associated with warm, subtropical waters. An increase in E_k with temperature may represent a strategy of acclimation to high light (Geider et al., 1998; Behrenfeld et al., 2004) and was observed in samples where both P_m^B and α^B decrease (**Figures 6A–G**). The drop in α^B however, was greater than the drop in P_m^B (**Figures 6C,F**), possibly reflecting an energetic investment in photo-protection by nano- and pico-phytoplankton under these conditions (Babin et al., 1996; Bouman et al., 2000; Raven et al., 2005). We observed no significant relationship between a_{ph}^* and temperature in each size class (**Figures 6D,H**).

The majority of the data that we used were from the NW Atlantic (N = 1873) and dominated by micro- and nanophytoplankton (Figures 1A,B), compared to N = 389 in the oligotrophic Atlantic Ocean where pico-phytoplankton only dominated at stations in the Gulf of Guinea and NW subtropical Atlantic (Figure 1C). This skewed the relationships toward larger size classes that inhabit colder waters. For both methods over a temperature range of 0 to 8°C, the samples are dominated by micro-phytoplankton with some samples also dominated by nano-phytoplankton (Figures 6A,E). In the open ocean, nano-phytoplankton rarely comprises the entire phytoplankton community. Over the temperature range from 8 to 18°C, fewer samples were dominated by a single size class using the Sathyendranath et al. (2009) method (Figure 6E), and the phytoplankton community was represented more by mixed populations, as classified by the Uitz et al. method (Figure 6A). At temperatures $>18^{\circ}$ C, both methods showed a similar distribution in classifying samples dominated by picophytoplankton.

Using global in situ data Behrenfeld and Falkowski (1997) suggested a 7th-order polynomial relationship between temperature and the optimum photosynthetic rate for a given in situ productivity profile, P_{opt}^B , which incorporates a reduction in photosynthesis at high temperatures. P_{opt}^B is distinct from P_m^B since it is derived from longer *in situ* incubations (Behrenfeld and Falkowski, 1997) as it also describes the optimum value at a discrete bottle depth rather than an instantaneous rate at the optimum irradiance level. The shape of the relationship found in the study of Behrenfeld and Falkowski (1997) fit midway between the 10 and 90% quantiles shown in this study (Figures 6A,B). The 90% quantile fit for P_m^B can be regarded as the phytoplankton response under optimum conditions. In natural assemblages however, "optimal" conditions are rarely observed. The majority of production rates frequently fall below this upper bound due to the effect of changes in the community structure, nutrient, and/or light-limitation, which is illustrated by the scatter around the upper quantile (Figures 6A,B).

Temperature also co-varies with light (Finkel et al., 2010), stratification and nutrient supply (Sathyendranath et al., 1991; Kamykowski et al., 2002) suggesting that significant correlations between temperature and phytoplankton size-class are also likely to be due to these other co-varying parameters (Bouman et al., 2005). At higher temperatures for example, there is considerable spread in the data, which is likely a reflection of a myriad of other factors, such as nutrient availability and light, that govern the magnitude in P_m^B (Behrenfeld and Falkowski, 1997; Huot et al., 2007). The strong negative correlation between temperature and P_m^B at higher temperatures probably reflects the strong negative correlation between photosynthesis and

nutrient supply in the open ocean (Sathyendranath et al., 1991; Kamykowski et al., 2002). ϕ_m is also strongly dependent on nutrient status and the low nutrient regime associated with warm, stratified areas often results in very low quantum yields (Falkowski et al., 1992; Platt et al., 1992; Sosik and Mitchell, 1994; Sosik, 1996). Samples dominated by pico-phytoplankton are often found where nutrient concentrations are extremely low and surface irradiance levels are high, which can result in a negative relationship between α^B and temperature (Platt et al., 1992; Sosik, 1996), which corresponds with the patterns that we observed (Figures 6B,F). We did not however find significant correlations between size-fractionated photosynthetic parameters and nitrate (Figure 7). The lack of trend in these relationships is because ambient nutrient concentrations are generally uncoupled from changes in phytoplankton biomass and size-class such that nutrients can be high when phytoplankton biomass is low or at the peak of a bloom when phytoplankton biomass is high especially micro-phytoplankton, dissolved inorganic nitrogen (DIN), soluble reactive phosphorus (SRP) or dissolved silicate (DSi) may be used up rapidly and therefore be low. Phytoplankton biomass and nutrients remain out of phase resulting from the lag response between injection of nutrients into the euphotic zone and phytoplankton growth. For both methods the range in P_m^B and α^B was high at nitrate concentrations <1 μ molL⁻¹ and was low at higher concentrations (Figures 7A,C,E,F). For the Uitz et al. (2008) method, the highest α^B occurred at nitrate concentrations of between 5 and 10 μ molL⁻¹. This pattern was not evident using the Sathyendranath et al. (2009) method for classifying size-class PE parameters.

Similarly we found no significant trend between P_m^B and α^B and PAR (Figure 7). This may be due to a combination of photo-acclimation and photo-inhibition. Phytoplankton can become photo-acclimated to their optimum photosynthetic rate over a large range in PAR (Sosik, 1996), which is evident by the narrow range in P_m^B and α^B and similar values for micro-, nano- and pico-phytoplankton. Sosik and Mitchell (1994) found that at high irradiance there is no significant difference in ϕm between micro- and nano-phytoplankton. In addition, high surface irradiance can result in the synthesis of photoprotective pigments to prevent damage to the photosystems (Dimier et al., 2007), which can potentially reduce α^B and ϕm (Babin et al., 1996; Bouman et al., 2000), which is also evident in our data (Figures 7D,H). We found the xanthophyll cycle transformation product Diatoxanthin was present in ~25% of our samples, suggesting that a quarter of the data set was affected by photo-protection. For the Uitz et al. (2008) method, micro-phytoplankton had higher P_m^B and α^B at lower PAR (~15 μ m⁻²d⁻¹), nano- at ~20 μ m⁻²d⁻¹, and pico- at higher irradiance (>40 μ m⁻²d⁻¹). For the Sathyendranath et al. method these patterns changed, with micro- P_m^B and α^B being generally lower than values for nano- and pico-phytoplankton (Figure 7F), which had the highest values at 15-20 and >40 μ m⁻²d⁻¹. These differences arise from the treatment of mixed populations by each method.

In summary, we found that micro- and nano-phytoplankton *PE* parameters are highly correlated with temperature. We found

no significant relationship between size-class *PE* parameters and nitrate and PAR. We recommend the Sathyendranath et al. method for classifying micro-phytoplankton over the temperature range of $0-8^{\circ}$ C and pico-phytoplankton at temperatures >18°C during blooms. The Uitz et al. method is recommended for mixed micro- and nano-phytoplankton populations over the temperature range of $8-18^{\circ}$ C. The *PE* parameters of these size classes could be derived from satellite sea surface temperature using these methods from the regression equations given in **Table 3**.

As remote-sensing models for estimating the biomass of phytoplankton functional types or size classes have developed (Sathyendranath et al., 2001; Nair et al., 2008; Brewin et al., 2010; Uitz et al., 2012), global estimates of primary production for particular groups of phytoplankton are improving. Models such as that of Uitz et al. (2012) have demonstrated the importance of size structure in estimates of primary production, where inclusion of different size classes of phytoplankton lowered the primary production estimates for the Mediterranean by a factor of 2. By linking the size-specific relationships for phytoplankton photosynthetic rates with temperature using the equations given in Table 3, class-specific primary production can be obtained. Such models will enable higher spatial and temporal frequency estimates of primary production in different size classes, which could be applied, through data assimilation, to improve the accuracy of biogeochemical models of the pelagic ecosystem.

SUMMARY

Two chemotaxonomic methods of classifying phytoplankton size taxa showed similar temperature-photosynthesis size-class dependence. There were no significant differences between these methods for mean photosynthetic parameters of the different size classes. Mean P_m^B and α^B were significantly higher for nanophytoplankton compared to micro- and pico-phytoplankton. The temperature ranges occupied by the size classes derived using the Uitz et al. (2008) method however exhibited a wider

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temperature range than those derived from the Sathyendranath et al. (2009) method. These differences arise from way that mixed populations are classified. Based on the relationship between *PE* parameters in different size classes and temperature, we recommend using the Sathyendranath et al. (2009) method to derive micro-phytoplankton *PE* parameters in sea water temperatures up to 8°C during monospecific blooms and the Uitz et al. (2008) method to derive *PE* parameters of mixed populations over the temperature range from 8 to 18°C.

AUTHOR CONTRIBUTIONS

AR processed and analysed the data, prepared the figures and drafted the manuscript as part of his PhD thesis. HB, GT, and SS conceived and directed the research, provided the *in situ* data, contributed to the data analysis and writing of the paper. HB obtained the funding to carry out the research. GT distilled the PhD chapter into a paper and provided Case support funding. All authors have contributed substantially to the study's conception, data acquisition and analysis and to writing the manuscript.

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Conflict of Interest Statement: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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