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Vertical heterogeneity of hydrocarbon-degrading bacteria in a core sediment sample from the Central Indian Ridge

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Hydrocarbons are ubiquitous in marine environments and might fuel hydrocarbon-metabolizing microbes in the ocean. Numerous studies have documented microbial hydrocarbon degradation in water columns and deepsea surface sediment. However, the degradation potential and biogeochemical cycling of hydrocarbons in subsurface sediments remain largely unknown. In this study, we used two different hydrocarbons, n-hexadecane (HEX) and methylcyclohexane (MCH), to investigate the distribution and diversity of hydrocarbon-consuming bacteria in a core sediment sample from the Central Indian Ridge (CIR), which is adjacent to mid-ridge hydrothermal vents in the Indian Ocean. We observed different vertical profiles of HEX- and MCHdegrading bacteria in the core sediments. Specifically, HEX-degrading bacteria were universally distributed, while MCH-degrading bacteria were found only in the intermediate layers of the core sediments. Changing factors including dissolved oxygen might affect the natural distribution of different hydrocarbon consumers. We found that a novel species of the genus C1-B045 might play a pivotal role in metabolizing MCH in the CIR deep biosphere. Through amino acid identity comparison with published sequences, we determined that C1-B045 harbors two novel classes of cyclohexanone monooxygenases involved in MCH metabolism. This study sheds light on the structure and function of hydrocarbon-consuming microbes in deep biospheres.

KEYWORDS

deep biosphere, subsurface sediment, biodegradation, cycloalkane, cyclohexanone monooxygenases, metagenomics



1 Introduction

Hydrocarbons are constantly introduced into the marine environment via both natural oil seeps and anthropogenic oil spills and are therefore universally distributed in the ocean. A diverse assemblage of hydrocarbon-containing crude oils regularly escapes into the marine environment via natural seepage from underground reservoirs (Atlas and Hazen, 2011). In addition, the hydrocarbon biosynthesis by cyanobacteria also contributes alkanes and alkenes (C13-C17) to the ocean (Schirmer et al., 2010). Love et al. (2021) estimated that the global flux of pentadecane (C15) produced by cyanobacteria exceeds the total oil input in the ocean by at least 100-fold (Love et al., 2021). Alkaline hydrothermal systems also contribute to the abiotic synthesis of hydrocarbons (even long-chain alkanes up to C24) and other organic compounds under hydrothermal conditions (Konn et al., 2009; Colman et al., 2017; He et al., 2021). Therefore, the environmental impact, biogeochemical cycling, and fate of hydrocarbons in the ocean have drawn a considerable amount of attention. Microorganisms play a pivotal role in driving the ocean hydrocarbon cycle (Lea-Smith et al., 2015; LaRowe et al., 2020; Zhou et al., 2020). To date, numerous studies have documented microbial hydrocarbon degradation in the marine environment, including in the water column and surface sediment (Yuan et al., 2015; Zhang et al., 2018). However, insufficient attention has been paid to studies on the biodegradation of hydrocarbons in subsurface sediment.

Previous studies have described the abiotic source of hydrocarbons in subsurface environments. McCollom and Seewald (2001) reported that the reduction of CO_2 during

serpentinization produces methane and other hydrocarbons abiotically (McCollom and Seewald, 2001). Moreover, Foustoukos and Seyfried (2004) suggested that chromium-bearing minerals catalyzed a Fischer-Tropsch type of synthesis of hydrocarbons (methane, ethane, and propane) through water–rock interactions in mid-ocean ridge hydrothermal systems. The abiotically produced hydrocarbons maybe sustain the growth of microbes in the subsurface environment at hydrothermal zones (Foustoukos and Seyfried, 2004). Bains et al. (2015) also indicated that hydrocarbons maybe fuel microorganisms in the deep biosphere to a depth of 10 km, except when the temperature was higher than a hypothetical life-maximum of 150°C (Bains et al., 2015).

The oceanographic context of this work was the Central Indian Ridge (CIR), which is proximate to the Edmond and Kairei hydrothermal vents; the distance from our sampling site to these two hydrothermal vents is within 200 km. Therefore, we speculated that the geochemistry of the CIR subsurface sediments might be impacted by matter (e.g., hydrocarbons and metals) released from nearby hydrothermal systems (Resing et al., 2015). Of note, Peng et al. (2011) analyzed the aliphatic hydrocarbons in the Kairei sediments and found that C16 and C17 alkanes exhibited the highest concentrations in the CIR hydrothermal field (Peng et al., 2011). Therefore, the CIR subsurface sediment might house distinct microorganisms fueled by abiotic hydrocarbons related to hydrothermal activity. This study could provide new insight into the geochemical cycle and fate of hydrocarbon in deep biospheres.

The purpose of this study was to 1) investigate the biodiversity, vertical distribution, and degradation potential of hydrocarbon consumers in a 4.8 m long gravity core from the CIR by adopting

two different substrates as sole carbon sources (n-hexadecane as labile hydrocarbon and methylcyclohexane as recalcitrant hydrocarbon); 2) identify the key uncultured cycloalkane consumers in deep-sea subsurface sediment through metagenomics; and 3) explore the key functional genes and metabolic pathways in cycloalkane-consuming bacteria derived from deep subsurface habitats.

2 Materials and methods

2.1 Sediment sampling

A sediment sample (CIR-GC01) was collected from the CIR in 2019 using a gravity column sampler on board the R/V Da-Yang-Yi-Hao. The sampling site ($24^{\circ}37'S$, $68^{\circ}46'E$, 4034 m depth) was proximate to the Edmond ($23^{\circ}53'S$, $69^{\circ}36'E$, 3300 m depth) and Kairei ($25^{\circ}19'S$, $70^{\circ}02'E$, 2415-2460 m depth) hydrothermal vents. The depth of this core sediment was approximately 4.8 m below the seafloor (mbsf). Core sediments were subsampled at various depths, including at 0 m, 1.0 mbsf, 2.0 mbsf, 3.0 mbsf, and 4.8 mbsf. Then, these subsamples were divided into three groups for the enrichment experiment and downstream data analysis, namely the shallow layer (0 m and 1.0 mbsf), the intermediate layer (2.0 mbsf and 3.0 mbsf), and the deep layer (4.8 mbsf). The samples were kept at -20° C on board prior to enriching hydrocarbon-consuming bacteria.

2.2 Measurement of hydrocarbons in the core sediment

We measured the hydrocarbons in the core sediment as described by Zheng et al. (2022). More specifically, sediment samples were prepared by freeze-drying for 24 h. Next, 5.0 g (dw) of each sample was placed into a 50 mL quartz colorimetric tube (with stopper) and 10 mL of hexane (\geq 97%)/dichloromethane (\geq 99.5%) solvent mixture (1:1, v/v) was added for ultrasonic extraction. Then, the supernatant was transferred into a 60 mL flask. The extraction procedures were repeated and the supernatant was combined. After that, the extract was concentrated to 2 mL at 40°C using rotary evaporation. Finally, the eluent was purged to 1 mL using N₂ at a flow rate of 30–50 mL·min⁻¹.

Hydrocarbon concentrations in core sediments were analysed using an Agilent 7890A gas chromatograph (GC) coupled with a microelectron capture detector. The GC is equipped with an HP-5MS fused silica capillary column. The temperature of the injector and detector was adjusted to 300°C. The oven temperature of the GC was set to start at 70°C, rise to 150°C at 20°C·min⁻¹, then to 300°C at 7°C·min⁻¹, and finally hold at 300°C for 7 min. High-purity helium was set at a flow rate of 1 mL·min⁻¹. The concentrations of standard solutions for HEX were 0, 1, 2, 5, and 10 mg·L⁻¹. The concentrations of standard solutions for MCH were 200, 400, 500, 800, and 1000 µg·L⁻¹. Each standard was measured in triplicates. The correlation coefficients of the calibration regression curves were 0.9965 for HEX and 0.9942 for MCH. The hydrocarbons in the abovementioned samples were identified according to the retention time of the peak area. Finally, the data were processed to determine the concentrations of hydrocarbons in the core sediment using Chemstation software (F.01.03.2357).

2.3 Enrichment of hydrocarbon-degrading bacteria from the core sediment

The different hydrocarbon-degrading bacteria from the core sediment sampled at the CIR were enriched in accordance with the approach described by Cui et al., Cui et al., 2022). Two species of hydrocarbons, namely n-hexadecane (HEX; Chemic, purity > 99%) and methylcyclohexane (MCH; Aladdin, purity 99%), were used as the sole sources of carbon and energy, respectively. 0.5 g of sediment was inoculated into the serum bottle containing 150 mL of ONR7a medium (Wang et al., 2008) amended with HEX/MCH (final concentration 0.23 mM/0.52 mM). Two duplicates were prepared for each treatment. The dissolved oxygen (DO) in the closed serum bottles was monitored daily via a highly sensitive fiber-optic oxygen meter (Pyroscience, FireStingO2). The enrichment cultures were incubated in the dark at 10°C without shaking until the microbial hydrocarbon respiration drew the DO concentration down to $0 \, \mu M$ in the closed serum bottles (bloomed hydrocarbon-consuming cultures). After incubation, microbial biomass was collected by filtration of the enrichment cultures using a 0.22-µm polycarbonate membrane (Whatman) and kept these samples at -80°C prior to DNA extraction.

The bloomed enrichment cultures were inoculated into the abovementioned media at an inoculum size of 7% (v/v). Then, the subcultures were incubated under the abovementioned conditions until they bloomed. The microbial biomass of the hydrocarbon-consuming subcultures was collected and stored as described above.

2.4 16S rRNA gene amplicon sequencing of different hydrocarbondegrading bacterial assemblages

Microbial DNA was extracted from the core sediment subsamples described in 2.1 (two duplicates for the treatment of original bacterial communities) and enrichment cultures/ subcultures described in 2.3 (two duplicates for each hydrocarbon-spiked treatment) using the OMEGA soil DNA kit (Omega Bio-Tek, Norcross, GA, USA) according to the manufacturer's protocols. The concentration of extracted DNA was measured using a NanoDrop NC2000 spectrophotometer (Thermo Scientific, Waltham, MA, USA) and the DNA integrity was assessed using electrophoresis on a 0.8% (w/v) agarose gel.

Next, the V3-V4 region of the 16S rRNA gene was amplified using PCR with primer pair 338F (5'-ACTCCTACGGGAGGCAGCA-3') and 806R (5'-GGACTACHVGGGTWTCTAAT-3'). Specifically, PCR reactions were performed in triplicate using 25 μ L of the mixture containing 0.25 μ L of Q5 high-fidelity DNA polymerase, 5 μ L of 5× reaction buffer, 5 μ L of 5× high GC buffer, 2 μ L of dNTPs (10 mM), 1 μ L of each primer (10 μ M), and 2 μ L of DNA template. The PCR reagents were purchased from NEB and Invitrogen, USA. Thermal cycling consisted of initial denaturation at 98°C for 5 min, followed by 25 cycles consisting of denaturation at 98°C for 30 s, annealing at 53°C for 30 s, and extension at 72°C for 45 s, with a final extension of 5 min at 72°C, then hold at 12°C.

The amplicons were extracted and purified from 2% (w/v) agarose gels using the Axygen DNA gel recovery kit (Axygen

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Biosciences, Union City, CA, USA) according to the manufacturer's protocols. Then, purified amplicons were sequenced with pairedend (PE) 250 strategy on an Illumina NovaSeq platform using the NovaSeq 6000 SP reagent kit according to the manufacturer's protocols at Personal Biotechnology Co., Ltd. (Shanghai, China). The raw reads were deposited into the NCBI Sequence Read Archive (SRA) database under accession number PRINA876038.

2.5 Metagenomic sequencing and analysis of the MCH-degrading bacterial communities

Microbial DNA of the MCH-degrading bacterial communities was extracted from the membrane samples (eight samples derived from the MCH-spiked enrichment cultures and subcultures) described in 2.3 using the OMEGA Mag-Bind soil DNA kit (Omega Bio-Tek, Norcross, GA, USA) according to the manufacturer's protocols. The extracted DNA was measured (optical concentration > 2.5 ng· μ L⁻¹ and total content > 200 ng) using a NanoDrop NC2000 spectrophotometer and agarose gel electrophoresis. Next, libraries with insert sizes of 400 bp were constructed for metagenome shotgun sequencing using the Illumina TruSeq Nano DNA LT library preparation kit. Each library was sequenced with the PE150 strategy on an Illumina NovaSeq platform at Personal Biotechnology Co., Ltd. The raw reads of metagenomics were deposited in the NCBI SRA database under accession number PRJNA876059.

Non-target fragments were removed from the raw sequences using Cutadapt (v1.17) and FASTP (v0.20.0), and the clean reads were obtained using SortMeRNA and bmtagger. Next, we assembled each sample, removed the unmatched sequences, and obtained the contigs set without redundancy using MEGAHIT (Li et al., 2015). MetaWRAP was used to performed the metagenomic binning (https://github.com/ bxlab/metaWRAP). The alleles were sorted to obtain the original bins using MaxBin2, metaBAT2, and CONCOCT with default settings. An improved bin set was produced using the bin_refine module, which combines two or three bin sets obtained by different software (or the same software with different parameters). Then, CheckM was used to assess the completeness and incompleteness of the bins. The genomic features (protein-coding genes and RNA) and their functions were annotated with rapid annotation using the subsystem technology (RAST) annotation engine. The phylogenetic information of the bins were annotated using the classify_wf model in GTDB-tk, which classifies genomes using the Genome Classification Database (Chaumeil et al., 2020). In addition, phylogenetic trees were constructed directly from the genomes using the de_novo_wf model in GTDB-tk. All parameters in de_novo_wf were set to the default values. The assembled bins of metagenomic datasets were deposited into the NCBI GenBank database (accession number PRJNA892034).

2.6 Phylogenetic analysis of the bins and key genes involved in MCH metabolism

The average nucleotide identity (ANI) was analysed to illustrate the phylogenetic relationship between the bins using ANI webbased software provided by Kostas lab (http://enveomics.ce.gatech.edu/ani/index). The software calculator estimated the ANI between two bins using both best hits (one-way ANI) and reciprocal best hits (two-way ANI). In addition, the amino acid identity (AAI) was analysed between the key genes involved in MCH metabolism using Clustal Omega (https://www.ebi.ac.uk/ Tools/msa/clustalo/).

2.7 Statistical analysis

Statistical analyses were performed using R (v3.6), QIIME2 (2021.8), and Vegan package (v2.5-6). Specifically, stacked histograms of species composition was plotted using QIIME2. The alpha diversity index with a normalized amplicon sequence variant table was calculated using QIIME2. Non-metric multidimensional scaling (NMDS) analysis was performed to present beta diversity using the Vegan package. The statistical significance was determined using the Dunn test and the Kruskal– Wallis test. Finally, pheatmap package was used for visualization.

3 Results

3.1 Geochemical features of the core sediment from the CIR

The core sediment collected from the CIR was pale yellow with a water content of 46–58% (shallow layer to deep layer). We detected the presence of HEX and MCH at various depths of the core sediment. Specifically, the concentration of HEX in the core sediment ranged from 2.33 mg·kg⁻¹ to 3.05 mg·kg⁻¹. The concentration of MCH in the core sediment ranged from 111.53 μ g·kg⁻¹ to 591.22 μ g·kg⁻¹ (Figure S1). The average concentration of HEX was 6.65 times higher than that of MCH. The source of these hydrocarbons is unknown. However, we speculated that they might come from the process of long-term sedimentation affected by proximate hydrothermal activities.

3.2 Hydrocarbon consumers derived from the CIR core sediment

We determined the microbial respiration of hydrocarbons in the enrichment cultures by monitoring the remaining DO level throughout incubation. The initial DO concentration was $223 \pm 14 \mu$ M (n = 10) for the enrichment cultures with HEX as the sole carbon source (Figure S2A). All the HEX-spiked enrichment cultures completely consumed the oxygen in the medium within 25 days of incubation (bloomed cultures; 0 μ M, n = 10), indicating microbial growth and respiration on alkane as the sole carbon source. Therefore, HEX-degrading bacteria derived from various depths of the CIR core sediment proliferated in the enrichment cultures after incubation.

The initial DO concentration was $223 \pm 12 \,\mu$ M (n = 10) for the MCH-spiked enrichment cultures (Figure S2B). The DO in the

MCH-spiked enrichment cultures derived from two shallow sediment layers (0 m and 1.0 mbsf) and one deep sediment layer (4.8 mbsf) did not fall below 92 \pm 24 μ M (n = 6) after incubation (> 70 days). These enrichment cultures failed to achieve substantial respiration of the hydrocarbons. In contrast, the MCH-spiked enrichment cultures derived from two intermediate sediment layers (2.0 mbsf and 3.0 mbsf) bloomed within 50 days of incubation (0 μ M, n = 4). Therefore, MCH-degrading bacteria derived from the intermediate sediment layer proliferated preferentially in the enrichment cultures after incubation. In addition, we observed that the average incubation time for bloomed cultures of the MCH treatment was two times longer than that of the HEX treatment. In other words, MCH is more recalcitrant than HEX for microbes.

The initial DO concentration for the HEX/MCH-spiked subcultures was comparable to that of the enrichment cultures (Figures S2C, D). After incubation, all the HEX-spiked subcultures bloomed within 12 days of incubation (0 μ M, n = 10). Similarly, the MCH-spiked subcultures derived from two intermediate sediment layers bloomed within 15 days of incubation (0 μ M, n = 4). In addition, we observed that the incubation time for the bloomed MCH-spiked subcultures was at least 30 days less than that of the corresponding enrichment cultures.

3.3 Biodiversity of different hydrocarbondegrading communities derived from the CIR core sediment

The alpha diversity of the hydrocarbon-degrading bacterial communities derived from the CIR core sediment was significantly lower (Kruskal– Wallis test, P < 0.01) than that of the blank treatment of the original bacterial communities in the subsurface sediment (Table S1). The supplement of HEX/MCH as the sole source of carbon and energy probably fuelled the growth of specific microbes derived from the subsurface sediment and decreased the alpha diversity of hydrocarbon-spiked enrichment cultures and subcultures compared with the original bacterial communities. Furthermore, the result of the NMDS analysis showed the divergence of bacterial community structure between hydrocarbon-spiked treatments and blank treatment (Figure 1, stress value < 0.2). Specifically, both bacterial community structures of the hydrocarbon-spiked cultures exhibited a noticeable shift compared with the original bacterial communities and formed two separate clusters on the plot. In addition, the MCH-spiked communities exhibited different bacterial community structure from that of the HEX-spiked communities suggested by the result of ANOSIM (P < 0.01, Table S2), indicating that distinct bacterial taxa were involved in the consumption of labile/refractory hydrocarbons in the subsurface sediment.

3.4 Effect of different hydrocarbons on bacterial community dynamics

The bacterial community dynamics (phylum level) results showed that Proteobacteria and Firmicutes were predominant in the hydrocarbon-degrading enrichment cultures and subcultures



Divergence of different hydrocarbon-degrading bacterial communities derived from the CIR core sediment. The colored solid symbols denote the bacterial community structures of different treatments. MCH, methylcyclohexane-spiked enrichment cultures, and subcultures; HEX, n-hexadecane-spiked enrichment cultures, and subcultures; BLK, the original bacterial communities of the CIR core sediment without the addition of any carbon source.

derived from the CIR core sediment (Table S3). Furthermore, the bacterial community dynamics (genus level) results showed that each group of hydrocarbon-spiked bacterial communities derived from the CIR core sediment harbored distinct dominant species (Figure 2A; e.g., top three most abundant genera). More specifically, the dominant genera in the treatments of HEX-degrading communities were Alcanivorax, Marinobacter, and Pseudomonas, the total relative abundance of which accounted for 77.61% in enrichment cultures and 85.60% in subcultures, respectively. In contrast, the dominant genera in the treatment of MCH-degrading communities were the novel genus C1-B045, Halomonas, and Paenibacillus, the total relative abundance of which accounted for 55.41% in enrichment cultures and 83.34% in subcultures, respectively. Therefore, the degradation of labile/recalcitrant hydrocarbons was ascribed to distinct hydrocarbon consumers derived from the same core sediment. It is worth noting that the relative abundance of C1-B045 increased dramatically from extremely low relative abundance (< 0.001%) in the original bacterial community to 71.99% in the MCH-degrading subcultures, implying it plays a key role in cycloalkane decomposition.

The differentiated genera in the hydrocarbon-spiked treatments were illustrated using a standardized heatmap (Figure 2B). Specifically, strains of *Alcanivorax*, *Marinobacter*, and *Pseudomonas* in the HEX-degrading enrichment cultures and subcultures exhibited higher relative abundance than the rest of the treatments, which reinforced that they were cosmopolitan alkane consumers in the marine environment. Remarkably, strains of *C1-B045*, *Cobetia*, *Halomonas*, *Paenibacillus*, and *Pseudoalteromonas* in the MCH-degrading enrichment cultures and subcultures exhibited higher relative abundance than in the rest of the treatments, suggesting the pivotal role of these bacteria to metabolize MCH in the core sediment.



3.5 Functional genes involved in MCH metabolism and putative degradation pathways in representative bins

We reconstructed 24 high-quality metagenome-assembled genomes (MAGs) (completeness > 51.72% and contamination < 8.62%) by binning from seven metagenomic datasets derived from the MCH-consuming enrichment cultures and subcultures (Figure 3; Table S4). One out of the 24 bins was classified using GTDB-Tk as novel genus *C1-B045* (completeness > 55.17% and contamination < 1.72%), which has been previously reported as one of the pivotal players in MCH decomposition (Cui et al., 2022). Moreover, five out of the 24 bins were classified as *Pseudomonas* (completeness > 53.44% and contamination < 4.83%) and *Alcanivorax* (completeness > 51.72% and contamination <



Maximum likelihood phylogenomic tree of the 24 bins reconstructed from seven metagenomic datasets derived from the MCH-consuming cultures. The bins of the genus *Alcanivorax* are highlighted in light purple; the bins of the genus *Marinobacter* are highlighted in light blue; the bin of genus *C1-B045* is highlighted in light red.

8.62%), respectively. In addition, another three bins were classified as genus *Marinobacter* (completeness > 60.34% and contamination < 1.30%). All these bacteria have been previously reported as hydrocarbon consumers, though their capacity to degrade cycloalkanes has not yet been characterized.

The genus C1-B045 exhibited the highest abundance in MCHconsuming subcultures, implying that it might be a key MCH consumer in the intermediate layer of the CIR core sediment (Figure 2A). The genome size of the C1-B045 MAG was 3,294,068 bp, and the GC content was 50.2%. A total of 3,216 coding sequences were predicted in this MAG based on the RASTtk annotation scheme (Figure 4A). Specifically, we found two copies of alkane monooxygenase-like genes that activate the initial oxidation of MCH in the MAG of C1-B045. Moreover, we annotated eight copies of genes encoding cyclohexanone monooxygenase and one copy of the gene encoding 6hexanolactone hydrolase in the C1-B045 genome. We proposed an MCH degradation pathway based on the key enzymes annotated in the MAG of C1-B045 and the intermediates experimentally detected in MCH metabolism, including 4-methyl-2-oxepanone, 4-methylcyclohexanone, and 3-methyladipic acid (Donoghue and Trudgill, 1975; Koma et al., 2004) (Figure 4B).

4 Discussion

The source of HEX and MCH measured in the core sediment in this study might derive from either biogenetic hydrocarbons or abiogenetic hydrocarbons. Given the proximate hydrothermal activities from the Kairei and Edmond hydrothermal zones, we speculate that they probably originated from abiogenesis of hydrocarbon. In addition, Yamashita et al. (2023) also proposed that hydrothermal-derived black carbon was a source of recalcitrant dissolved organic carbon in the ocean, and could migrate long distances (4,440km) in the deep ocean with hydrothermal plumes. Hao et al.



Key genes encoding degradation pathway for MCH in the MAG of *C1-B045* derived from the CIR core sediment. (A) circular representation of the MAG of *C1-B045*. CGView map represents the coding sequence (annotated by RAST), contigs, GC skew, GC content, and genome backbone from outside to inside, respectively. *alkB*, alkane 1-monooxygenase; *chnB*, cyclohexanone monooxygenase; *chnC*, 6-hexanolactone hydrolase; *adh*, alcohol dehydrogenase; *aldh*, aldehyde dehydrogenase. (B) putative MCH degradation pathway with key enzymes predicted in the MAG of *C1-B045*. * indicates the transporter ABC is not annotated in this MAG.

However, source analysis is needed to further validate the origin of these hydrocarbons in the CIR core sediment in future studies.

In this study, a HEX/MCH-spiked enrichment experiment was used to investigate the vertical distribution of different hydrocarbondegrading bacteria in a CIR core sediment. We found that the labile hydrocarbon (HEX) consumers were widely distributed in the core sediment (0-4.8 mbsf), whereas the refractory hydrocarbon (MCH) consumers were stratified only in the intermediate layer of the core sediment (2.0 and 3.0 mbsf) (Figure S2). The vertical heterogeneity of the different hydrocarbon consumers might be ascribed to the gradient of DO along the vertical core sediment. Durbin and Teske (2011) reported that the concentration of DO in subsurface sediments decreased with increasing depth. They found that the DO dropped from 115 µM at 5 cm below the sediment (cmbsf) to 2-10 µM (an apparent detection limit) at 100-150 cmbsf (Durbin and Teske, 2011). This suggests that the absence of oxygen in the deep layers of the CIR core sediment (4.8 mbsf) was unfavorable to the aerobic degradation of the refractory hydrocarbons and the growth of refractory hydrocarbon consumers. MCH-consuming bacteria could not be enriched in the shallow layer of the core sediment. Therefore, we speculated that the labile hydrocarbon consumers outcompeted their counterparts of refractory hydrocarbon consumers in the shallow layer of the original core sediment where the DO was sufficient. In contrast, decreased DO in the intermediate layer of the original core sediment might have limited the growth of fast-growing labile hydrocarbon consumers (Massmig et al., 2019). Therefore, the slow-growing refractory hydrocarbon consumers coexisted with their counterparts of labile hydrocarbon consumers in the intermediate layer of the original core sediment. Among these refractory hydrocarbon consumers, we found the occurrence of cycloalkane-degrading bacteria C1-B045 in the deep-sea subsurface environment for the first time (Figure 2A). Importantly, the abundance of C1-B045 surged to 71.99% in MCH-consuming subcultures compared to < 0.001% in the original bacterial community, implying that it might be a potential key cycloalkane consumer in the deep biosphere of the CIR.

To the best of our knowledge, the genus *C1-B045* has not yet been cultured or characterized in detail since the first report of their

near-complete 16S rRNA gene sequences in the clone library derived from the Guaymas Basin sediment (Teske et al., 2002). Several of the most recent studies have demonstrated that it might also be involved in the microbial degradation of light crude oil (Uribe-Flores et al., 2019), polycyclic aromatic hydrocarbons (Peng et al., 2020), and short-chain n-alkylcyclohexanes (Arrington et al., 2022; Cui et al., 2022). Due to the poor culturability of genus C1-B045, we obtained only MAGs rather than pure cultures by binning from metagenomic datasets in this study. The ANI value between the MAG of C1-B045 derived from the CIR deep biosphere and its counterpart derived from Bohai Sea (BS) (unpublished data) was 85.84%, implying that they represent two different species of genus C1-B045 (Figure S3; cutoff value of ANI for species delineation < 95%)(Parks et al., 2020). Therefore, strains of the genus C1-B045 are phylogenetically diverse and might represent cosmopolitan cycloalkane consumers in different marine habitats. In addition to C1-B045, almost all the other differentiated bacteria in MCH-spiked cultures (Figure 2B) have been characterized to be capable of metabolizing hydrocarbons, implying they might take part in MCH metabolism. For instance, Govarthanan et al. (2020) reported that strains of Halomonas were capable of utilizing both low and high molecular weight polycyclic aromatic hydrocarbons for growth. Hedlund and Staley (2006) found that members of Pseudoalteromonas contain dioxygenase genes and can effectively decompose naphthalene.

Hitherto, two MCH degradation pathways have been proposed in previous studies, including lactone formation and aromatization. The lactone formation pathway, characterized by Koma et al. (2004) in *Rhodococcus* sp.NDKK48, might represent a primary mechanism for cycloalkane degradation. Moreover, the aromatization pathway was found in *Rhodococcus* sp. EC1, which might be relatively rare in aerobic bacteria (Yi et al., 2011). In our study, we proposed a lactone formation pathway in the MAG of *C1-B045* (Figure 4B), which is similar to that of *Rhodococcus* sp.NDKK48 (Koma et al., 2004). More recently, we discovered two nearly complete MCH degradation pathways, lactone formation and aromatization, concurrently in the MAG of *Roseovarius* derived from Antarctic surface water (unpublished data). Therefore, future research should be devoted to the metabolic diversity of cycloalkanes by marine microorganisms.



FIGURE 5

Phylogeny of CHMOs in MCH-consuming bacteria. CIR, CHMOs annotated in genus *C1-B045* derived from the Central Indian Ridge; BS, CHMOs annotated in genus *C1-B045* derived from the Bohai Sea; AcCHMO CA16, CHMO of *Acinetobacter calcoaceticus* CA16; RhCHMO HI-31, CHMO of *Rhodococcus* sp. HI-31; AcCHMO NCIMB9871, CHMO of *Acinetobacter calcoaceticus* NCIMB9871; XfCHMO, Of *Acinetobacter flavus*; BpCHMO, CHMO of *Brachymonas petroleovorans*; ArCHMO L661, CHMO of *Arthrobacter* sp. L661. Class I of CHMOs is highlighted in light purple; The class II of CHMOs is highlighted in light red.

Remarkably, we annotated abundant copies of genes encoding cyclohexanone monooxygenase (CHMO/ChnB) in the MAGs of C1-B045 derived from the CIR and BS. CHMOs belong to Baeyer-Villiger monooxygenases (BVMOs) and can catalyze a variety of oxidation reactions, including Baeyer-Villiger oxidations. This reaction involves the conversion of cyclic ketones into lactones or ketones into esters by introducing an oxygen atom adjacent to the carbonyl group (Yachnin et al., 2012). In this study, we discovered the presence of two classes of CHMOs based on phylogenetic analysis with the available amino acid sequences of CHMOs (Figure 5; Table S5). Specifically, class I CHMOs derived from the MAGs of C1-B045 exhibited low AAI values with that of Rhodococcus sp. HI-31(29.04-41.71%). Similarly, class II CHMOs derived from the MAGs of C1-B045 exhibited low AAI values with that of Acinetobacter calcoaceticus CA16 (28.60-31.87%). This indicates that CHMOs derived from genus C1-B045 represent two novel classes of enzymes (Kim et al., 2013; Walock et al., 2014), which exhibited different primary and advanced protein structures from the published reference CHMOs. In addition, eight pairwise CHMOs (one from the CIR and the other from the BS) were seen in the phylogenetic tree and exhibited high similarity to one another (90.53-99.36%).

5 Conclusion

In this study, we found that HEX-consuming bacteria were ubiquitous in the CIR core sediment. The vertical profile of the MCH-consuming bacteria was sandwich-like in the CIR core sediment. Specifically, the MCH-consuming bacteria were stratified only in the intermediate layer of the core sediment. We also discovered that a novel species of genus C1-B045 might be involved in the degradation of cycloalkane in the deep biosphere of the CIR. The bacteria of genus C1-B045 harbored two novel classes of CHMOs to break down the metabolites of MCH.

Data availability statement

The datasets presented in this study can be found in online repositories. The names of the repository/repositories and accession number(s) can be found in the article/Supplementary Material.

Author contributions

TH and ZC contributed to conception and design of the study. ZC and ZW led funding acquisition and project management. JC managed the resources, and guaranteed their availability to perform the experiments and analysis. TH and XL performed the experiments and analyzed the data. TH, ZC and XL wrote the first draft of the manuscript. TH, ZC and GZ are responsible for the review and editing of article. ZC and ZW supervised the writing of the entire manuscript. All authors contributed to the article and approved the submitted version.

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Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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Supplementary material

The Supplementary Material for this article can be found online at: https://www.frontiersin.org/articles/10.3389/fmars.2023.1170238/ full#supplementary-material

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