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The geomagnetic and geoelectric response to the May 2024 geomagnetic storm in the United Kingdom

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The “Gannon” geomagnetic storm of 10–12 May 2024 was the first extreme storm of the solar cycle 25 and the largest storm in more than 20 years. The auroral electrojet, driven by a strong negative interplanetary magnetic field exceeding -50nT , moved towards the equator in the evening of the 10th May reaching the latitudes of central and southern England (below 54N) for several hours. Widespread sightings of the aurora were observed across the country, and rapid variations of the magnetic field were recorded in the United Kingdom (UK). Here we present the geomagnetic and geoelectric data recorded during the storm in the UK together with models of ground effects and images of auroral displays around the country. We compare the May 2024 storm with geomagnetic data from the September 2017, October 2003, March 1989 as well as September 1859 Carrington event to demonstrate the differences in magnitude, timings and latitudinal extent between these events. We use the geomagnetic observations, and a ground electric field model based on magnetotelluric data combined with the high-voltage power grid network information to estimate geomagnetically induced currents (GICs) at substation level during the storm. The highest modelled GICs exceeded 60 A in substations in southwest and east-central England as well as northern Wales. Substation GICs modelled in higher latitude stations in Scotland exhibited lower values because the leading edge of the auroral oval rapidly moved to lower latitudes. The “Gannon” storm compared to historical storms on a global scale in terms of the aa* index, ranks third since 1868, after March 1989 and September 1941. However locally, the maximum magnetic field rate of change suggests it is closer to a 1-in-30 years event. Hence, there was relatively little impact on grounded technology in the United Kingdom.

KEYWORDS

space weather, geomagnetic storm, geomagnetic field, geoelectric field, geomagnetically induced currents

1 Introduction

Solar activity in the current solar cycle 25 exceeded prior forecasts, with monthly average sunspot number and F10.7 cm radio flux reaching much higher values than initially predicted

(Nandy, 2021). This could indicate the end of the longer-term trend of declining activity observed throughout the last four cycles (McIntosh et al., 2020). In the first 6 months of 2024 the Sun produced more solar flares than observed in a full year during the previous solar maximum in 2014 (according to NOAA SWPC solar flares count¹). On 15 October 2024, NOAA and the Solar Cycle Prediction Panel announced the official start of the solar maximum. Activity had started to reach its peak in early May 2024, when a large active region (AR13664) located at the centre of the solar disc just below the equator produced multiple M-class and X-class solar flares, most of which had an associated Earth-directed coronal mass ejection (CME). Spogli et al. (2024) provides a detailed description of the solar activity of May 2024. The combination of up to seven fast CMEs arrived just after 17:00 UT on the 10th May, leading to a significant increase in geomagnetic activity, lasting over 30 h (Hayakawa et al., 2025). The “Gannon” storm (named after Dr. Jennifer Gannon, a leading figure in space weather research, who passed away a week before the event) was the most extreme space weather event since the Halloween storm of October 2003, based on its duration, prolonged periods of high Kp index (Bartels, 1949) and minimum Dst index observed (Gonzalez et al., 1994). The duration and intensity of the storm rapidly pushed the auroral oval to lower latitudes, resulting in beautiful displays of the ‘northern lights’ (aurora borealis) across all of the United Kingdom which popularized the topic of space weather with the general public.

However, space weather poses a potentially serious threat to our technology-dependent society. Geomagnetically induced currents (GICs) flowing in ground-based systems during periods of enhanced geomagnetic activity can cause damage to power grids (Boteler et al., 1989; Pulkkinen et al., 2017; Rodger et al., 2017), railways (Liu et al., 2016; Love et al., 2019; Patterson et al., 2023), oil and gas pipelines (Khanal et al., 2019; Ingham et al., 2022; Oliveira et al., 2024) and potentially submarine cables (Chakraborty et al., 2022; Boteler et al., 2024). Extreme space weather can also disrupt satellite-based navigation by distorting or delaying signal transmission through the ionosphere (Roy and Paul, 2013; Hapgood, 2017; Demyanov and Yasyukevich, 2021). Various impacts of May 2024 storm have been summarized in several studies: Parker and Linares (2024) describes how the perturbations during the event affected satellite trajectories in low Earth orbit; Schennetten et al. (2024) outlines the impacts of the event on radiation fields at aviation altitudes; Aa et al. (2024) and Themens et al. (2024) investigate ionospheric disturbances and storm-enhanced density during the storm at high and mid latitudes; Diaz (2024) shows how the May storm disturbances affected seismic measurements.

After 20 years of predominantly mild geomagnetic conditions, solar cycle 25 has produced an increased frequency of eruptive events on the Sun and will likely cause further geomagnetic storms and, potentially, severe space weather. Therefore, recording information from extreme events, such as the ‘Gannon’ storm, is important for understanding their true impact on both a global and local scale. Continuous observations and further research in solar activity, propagation of CMEs in the solar wind and the resulting space and ground effects are essential in improving forecasting and nowcasting capabilities in order to reduce the impacts of extreme geomagnetic storms in the future.

In Section 2, 3.1, we present a detailed description of the local geomagnetic and geoelectric field during the May 2024 event using available measurements in the United Kingdom (UK). The geoelectric field across the United Kingdom was modelled using magnetotelluric data (Section 3.2) and then used to estimate the GICs in the high-voltage power grid to highlight potential vulnerabilities in the network (Section 4). Section 5 shows the aurora observations across the United Kingdom, indicating the wide spread of the auroral oval during the May storm in fortuitously clear skies. In Section 6 we compare the ‘Gannon’ storm with storm levels dating back to 1868 as well as with selected events in September 2017, October 2003, March 1989 and the Carrington event from September 1859, to examine their differences in severity and ground impacts. Final remarks and discussion points can be found in Section 7.

2 Geomagnetic field

The British Geological Survey (BGS) Geomagnetism group operates three INTERMAGNET-standard magnetic observatories in the United Kingdom - Lerwick (LER) in the Shetland Islands, Eskdalemuir (ESK) in the Southern Uplands of Scotland, and Hartland (HAD) in North Devon. The data are recorded by a number of suspended or non-suspended fluxgate magnetometers, with a sampling rate of 1-s. The observatories also measure absolute strength using Overhauser proton precession magnetometers (GEM Systems GSM-90), with a sampling rate of either 1, 5 or 10 s. Weekly absolute observations ensure baseline control and that high accuracy is maintained (Clarke et al., 2013).

In addition, three new variometer sites equipped with tri-axial fluxgate magnetometers only were set up during the United Kingdom Space Weather Instrumentation, Measurement, Modelling and Risk (SWIMMR) programme. This was a £20 million, 4-year programme which aimed to improve the UK’s capabilities for space weather monitoring and prediction. In 2022, as part of the SWIMMR Activities in Ground Effects (SAGE) project, the BGS installed variometers at sites in Florence Court, Enniskillen [FLO], Market Harborough, Leicestershire [LEI], and Herstmonceux Castle, Sussex [HTX], to measure magnetic field variation with a 1-s cadence. This filled in a spatial ‘gap’ between existing observatories and decreased the distance between observation points to below 200 km. The new sites predominantly improve the East-West extension compared to the North-South coverage offered by the existing United Kingdom observatories. The variometer systems installed are based on the design of the Space Weather Impact on Ground-based Systems (SWIGS) differential magnetometer systems (Hübert et al., 2020). The locations of the current United Kingdom observatories and variometer stations are shown in Figure 1.

The northward (Bx), eastward (By) and downward (Bz) ground magnetic field variations measured during the ‘Gannon’ storm are shown in Figures 2–4, respectively. The traces correspond to the six BGS magnetic sites located across the United Kingdom, arranged by descending latitude. To analyze space weather effects, only the external field components are needed and their values at 1-min resolution are plotted. The core and crustal field contributions were removed from each time series to isolate variations caused by magnetospheric and ionospheric currents. The arrival of the CMEs from early May resulted in compression of the magnetopause which led to a rapid

1 <https://www.spaceweatherlive.com/en/solar-activity/solar-cycle.html>

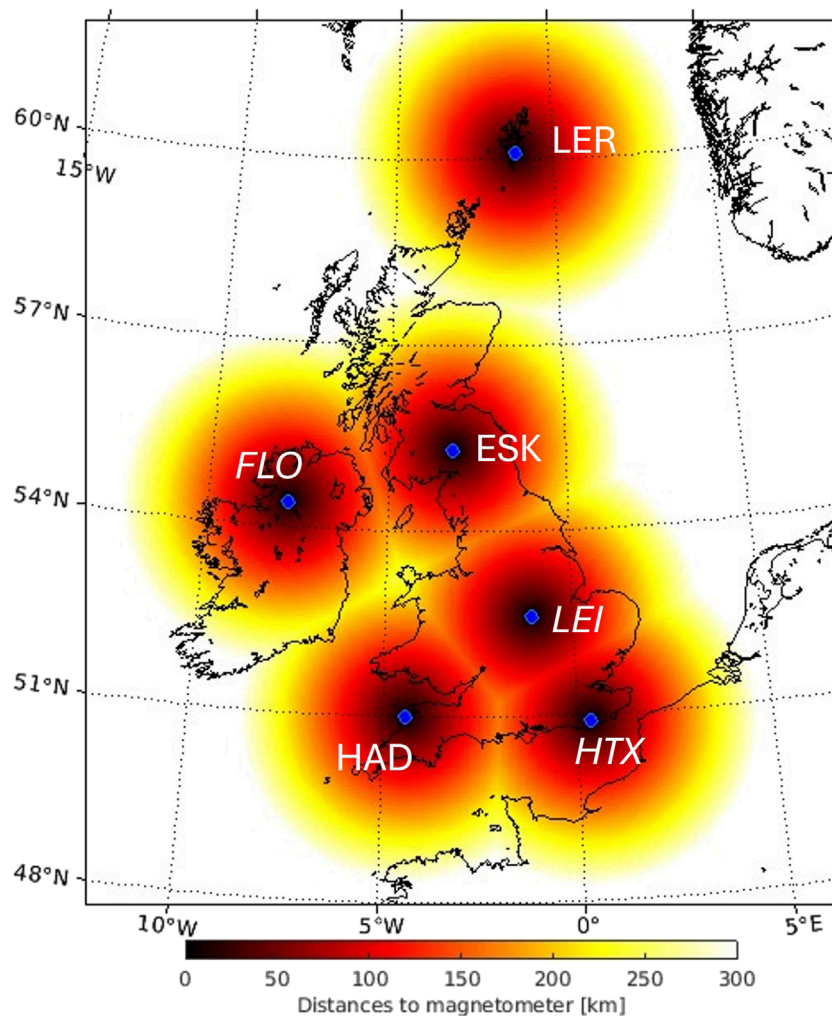


FIGURE 1 Locations of BGS magnetic observatories (LER, ESK, HAD) and variometer stations (FLO, HTX, LEI). Colors indicate the distance to the closest site.

increase in horizontal, mainly northward ground magnetic field. This process is often referred to as a sudden storm commencement (SSC), if followed by enhanced geomagnetic activity, or as a sudden impulse (SI) if no further disturbance occurs (Matsushita, 1962; Araki, 1994). The ‘Gannon’ storm SSC occurred at 17:08 UT on the 10th May, with the strongest ground response of $B_x = 320$ nT observed at ESK. All remaining United Kingdom magnetic stations measured variations in B_x between 200 nT and 300 nT. The lowest variation associated with the SSC ($B_x = 217$ nT) was recorded at LER, the highest latitude observatory. The eastward component overall reached much lower values during the SSC, with the highest variation of $B_y = 122$ nT also measured at ESK and the smallest $B_y = 48$ nT observed at HAD. The vertical component response to the SSC was not as significant, with maximum values reaching between around $B_z = 20$ nT at ESK and LEI, and up to 90 nT at LER.

A significant decrease in the horizontal magnetic field is typically observed during the main phase of a geomagnetic storm, usually shortly after the SSC, when the disturbance storm time (Dst) index reaches values below -50 nT (Gonzalez et al., 1994). At United Kingdom latitudes the minimum B_x values during the ‘Gannon’

event main phase (after 10 May 19:00 UT) varied between -770 nT and $-1,340$ nT, with the largest decrease measured at LER ($B_x = -1,339$ nT) and smallest at HAD ($B_x = -775$ nT), showing a gradual trend towards less negative values with decreasing latitude. The geomagnetic field remained highly disturbed throughout the 11th May and early morning hours of the 12th May before returning to predominantly quiet levels around 12:00 UT. Another noticeable sharp increase in the horizontal components, up to 65 nT in B_x and 25 nT in B_y measured at LER, can be observed at 09:20 UT on the 12th which is likely to be an SI response to one of many additional CMEs that arrived during the May 2024 storm.

Rapid fluctuations of the ground magnetic field during severe space weather events can be observed not only in the horizontal components but also in declination - the angle between True North and the field vector on the horizontal plane, measured positive eastward. Figure 5 shows how the ‘Gannon’ storm affected the compass variations at the three United Kingdom observatories. The largest deviation from True North, reaching almost 4° overall, was observed at LER on the 11th May during the main phase of the event. Although both ESK and HAD measured lower

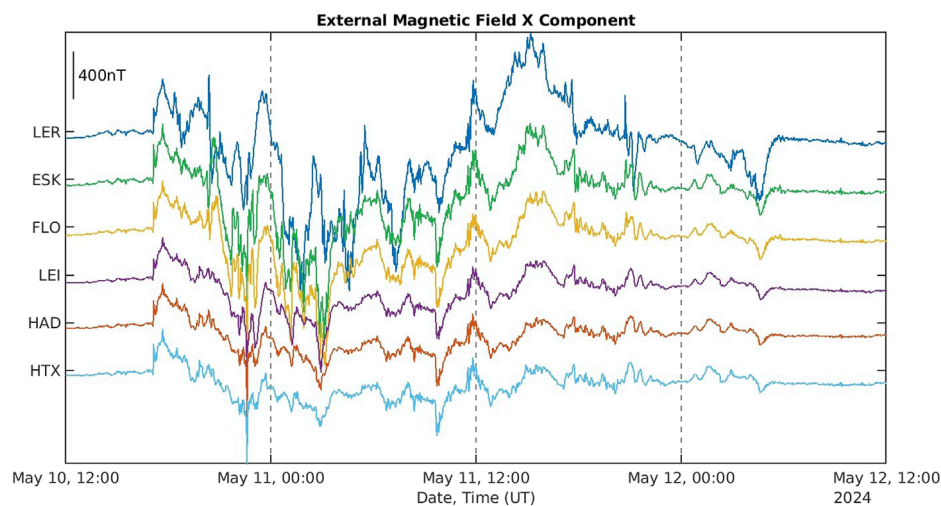


FIGURE 2
External geomagnetic field measurements (X/B_x component) from six United Kingdom magnetic monitors for 10–12 May 2024.

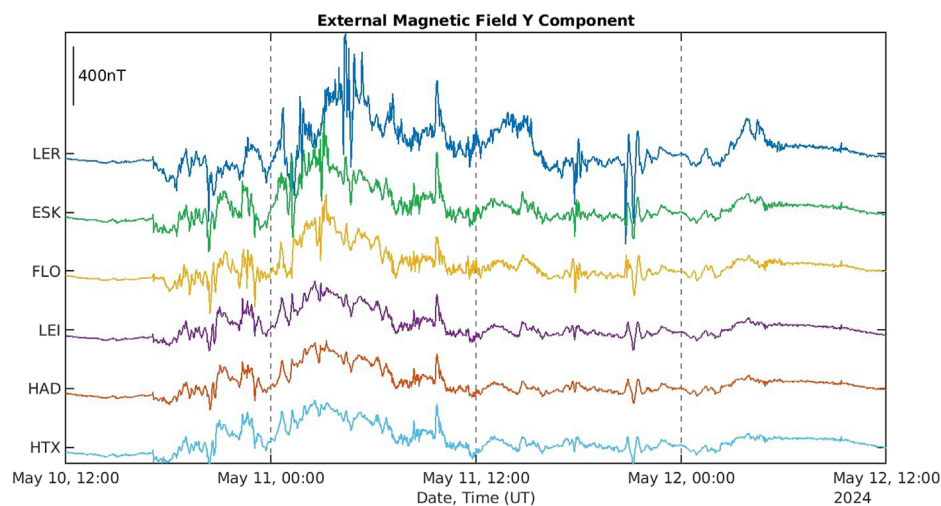


FIGURE 3
External geomagnetic field measurements (Y/B_y component) from six United Kingdom magnetic monitors for 10–12 May 2024.

variations in declination, the largest changes (around 2.4° and 1.2° , respectively) were observed around an hour earlier than at LER, which indicates the latitudinal progression of the storm related to the expansion of the auroral oval. Another noticeable rapid change in compass variation was observed at all observatories at 10:00 UT on the 11th May.

3 Goelectric field

3.1 Measurements

BGS operates ground electric field monitoring systems at the three magnetic observatories in the United Kingdom. These measure the changing electric field generated at the Earth's surface

(Beggan et al., 2021). Electric field data are recorded at 10 Hz along East-West and North-South electrode lines. The measurements for the May 2024 storm are shown in Figure 6. Unfortunately, the ESK site was affected by lightning the week before the May storm, damaging the electronics on the recording equipment of the electric field sensors. Therefore, ESK electric field data were not recorded during the event. The temporary loss of the electric field monitoring during such an extreme space weather event was disappointing and may have been a valuable data point in extreme value analysis and worst-case scenario studies, for example.

Both northward and eastward goelectric fields measured at the other sites (LER and HAD) show a distinct response to the SSC at 17:08 UT on the 10th May, reaching 0.24 V/km and 0.44 V/km at LER, and 0.06 V/km and 0.13 V/km at HAD, respectively. The electric field remained disturbed during the evening of the

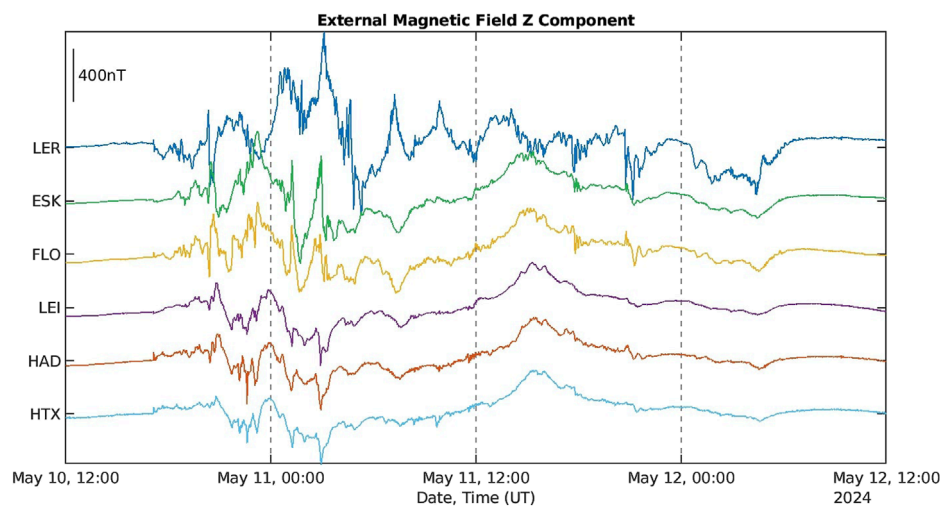


FIGURE 4
External geomagnetic field measurements (Z/Bz component) from six United Kingdom magnetic monitors for 10–12 May 2024.

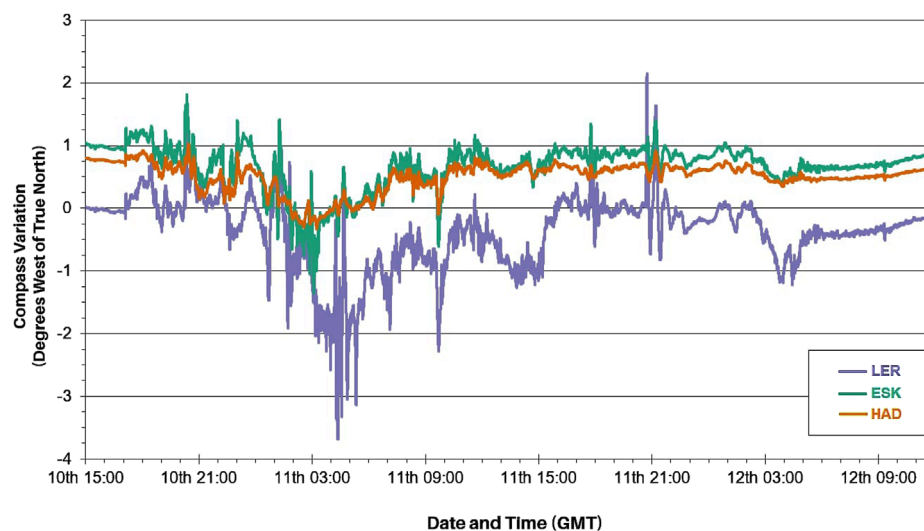


FIGURE 5
Compass variation for LER (blue), ESK (green) and HAD (orange) during 10–12 May 2024, given in degrees West of True North.

10th and all day on the 11th May before gradually declining to ambient levels in the morning on the 12th, clearly following the storm variations seen in the magnetic field. Measurements from HAD show noticeable peaks of 0.21 V/km in the northward, and 0.28 V/km in the eastward field at 22:36 UT on the 10th May. Disturbances at LER were more regular; however there is one distinctive peak of over 1 V/km in the North-South component at 20:20 UT on the 10th. The East-West field reached over 1 V/km on multiple occasions throughout the main phase of the storm. Slightly higher values in the East-West orientation correspond to the overall larger variations in the horizontal magnetic field observed during the storm. The brief and sharp SI variation seen in the magnetic field at 09:20 UT on the 12th was also observed in the electric field measurements, particularly at LER, reaching over 0.20 V/km.

3.2 Modelling

Electric field variations during geomagnetic storm times differ greatly across the country due to, not only the magnetic field variations, but importantly, also the electrical conductivity of the subsurface. The electrical conductivity of rocks at depth can be inferred with the magnetotelluric (MT) method. MT measurements comprise of simultaneous recordings of the horizontal magnetic and electric field variations. In the United Kingdom, long-period MT data have been collected in the past years to improve the geoelectric field model needed for space weather ground impact studies (Hübert et al., 2024c). The modelling approach follows the methodology described in Kelbert et al. (2017); Campanyà et al. (2019); Malone-Leigh et al. (2023); Hübert et al. (2024a).



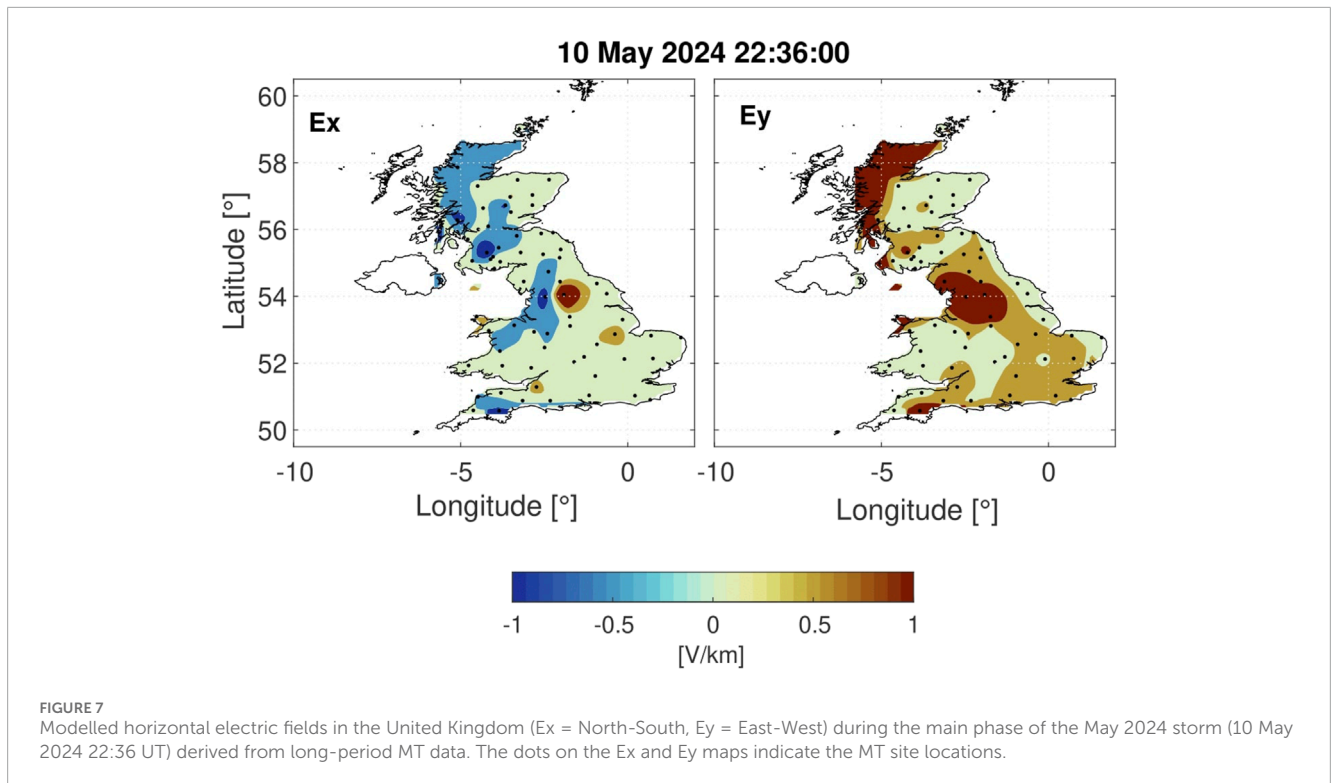
Using the MT data at 70 sites across the country and spherical elementary current system analysis (Amm, 1997; McLay and Beggan, 2010) to interpolate the magnetic field variations every minute based on the United Kingdom geomagnetic observatories and variometer sites, the geoelectric field variations during the ‘Gannon’ storm were modelled. Figure 7 shows the interpolated geoelectric field in N-S (E_x) and E-W (E_y) direction based on the convolution of MT transfer functions with the interpolated geomagnetic field at the peak of the storm 10 May 2024 22:36 UT, when the total geoelectric field estimated across Britain reached its maximum value. The highest amplitudes for the geoelectric field (> 1 V/km) are estimated for the northwest of Scotland and northern England with peak values reaching 3.7 V/km in the Peak District, East of Manchester.

4 Geomagnetically induced currents (GIC)

Rapid variations in the magnetic field, particularly during severe space weather events, induce electric fields in the conductive subsurface which generate quasi-steady direct currents known as geomagnetically induced currents (GICs) (Thomson et al., 2005). The induced currents can flow through low-resistance grounded infrastructure such as the high-voltage power transmission network (Pirjola, 1985), gas pipelines (Pulkkinen et al., 2001) and railways (Patterson et al., 2024), causing damage via overheating and even

harmonic disturbances (Rodger et al., 2020; Clilverd et al., 2020). GICs in the United Kingdom high-voltage power grid are computed based on a representation of the network topology provided by the Electricity Ten Year Statement (ETYS) 2022. The network model consists of over 430 substations with approximately 1,300 connections for the 400 and 275 kV electrical networks across the whole of the United Kingdom and the 132 kV network in Scotland, and 2044 transformer windings (Kelly et al., 2017). Using the modelled electric field values, described in the previous section, the size of GIC flowing through each substation in that power network can then be estimated during the geomagnetic storm. The assumptions of the model and validation are discussed in Hübert et al. (2024b) who validated the grid model using the differential magnetometer method (DMM) and demonstrated that knowledge of the geoelectric field is the major limiting factor at the present time. The geoelectric field used to compute GIC is not uniform in this simulation, as we use the MT transfer functions from 70 sites in Britain to build up a measurement-based map of the geoelectric field (Figure 7).

The estimated maximum absolute GIC values at each substation ($|GIC|$) during the ‘Gannon’ storm are shown in the upper panel of Figure 8. The largest $|GIC|$ tend to occur at the end of long lines and corner nodes, which are typically along the coast in Britain. The southwest and east-central coast of England, as well as northwest coast of Wales experience particularly large GICs. Note that most substations have multiple transformers with a common earthing mat, in which case the GIC flow is usually, but not always, divided



relatively equally between them. The |GIC| modelled at six United Kingdom substations reached over 60 A on the evening on the 10th May, during the main phase of the event, with largest value of 63 A at a substation near Norwich (East coast of England). This particular substation has five transformers and the largest transformer current modelled at this site is around 15 A. The Stalybridge substation (northern England) had the largest estimated transformer |GIC| of 41 A, on the high-voltage winding of a 400–275 kV transformer. The uncertainties on the modelled GIC values are strongly dependent on the geoelectric field model. From previous work using DMM measurements, we suggest the uncertainty is on the order of 10%–20% (Hübner et al., 2024b).

GIC measurements are only available from a few substations in Scotland, but the validity of the GIC model was previously provided by the comparison with line GIC measurements (Thomson et al., 2005; Hübner et al., 2020). There are, at present, no GIC measurements in England and Wales. This inhibits our capability to fully validate the modelled results during the May 2024 storm. The lack of GIC measurements at substations is a limitation and will be addressed in future years; this aspect of the modelling should improve when more data are available to compare with (e.g., Mac Manus et al., 2022).

5 Auroral observations

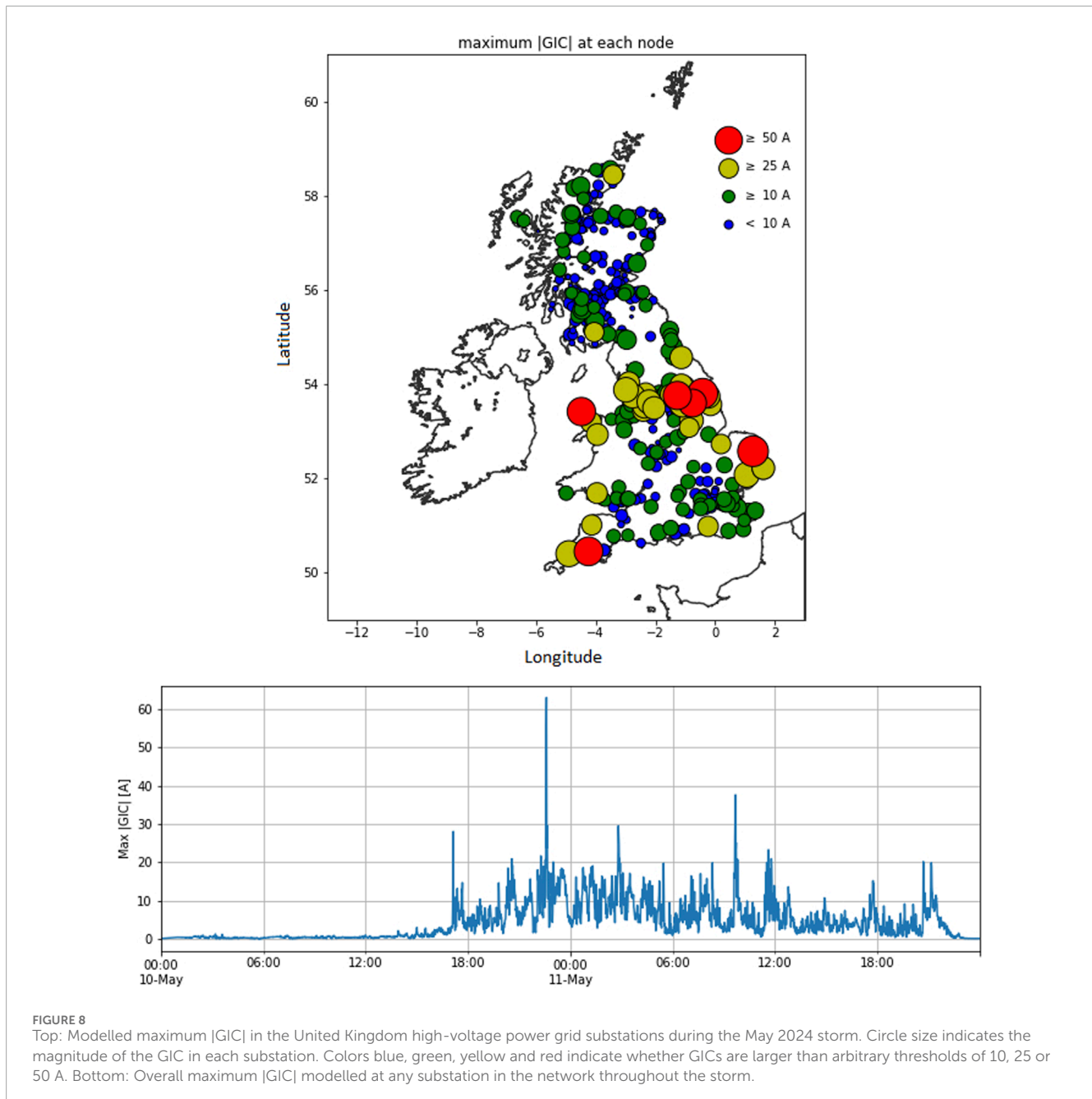
During periods of enhanced solar activity, electrically charged particles within the solar wind collide with the magnetosphere and travel along the magnetic field lines down to the magnetic poles, where they interact with atoms in the Earth's atmosphere. The collisions and interactions between solar particles and atmospheric

molecules result in excitation (heating) which causes them to glow with different colours depending on the molecular species. This phenomenon is commonly known as the northern lights or aurora borealis (when observed in the northern hemisphere). Although auroral displays are usually concentrated around the polar regions, they can sometimes be seen at lower latitudes when geomagnetic activity is significantly elevated. The best chances for aurora observations are during severe geomagnetic storms, such as the 'Gannon' 2024 event (Hayakawa et al., 2025). Thanks to the various aurora alert services available around the world, including the BGS aurora alert predictions for the United Kingdom issued on the afternoon of the 9th May², millions of people had the opportunity to observe and photograph the auroral displays during the May storm. The Skywarden catalogue³ contains over 180 observations recorded across a wide range of latitudes, between 30° and 60°, in both hemispheres. Thousands of people in the United Kingdom stayed up late at night on the 10th May taking photos and videos of the northern lights all across the country. Figure 9A shows a few examples of aurora sightings in Scotland provided by BGS staff in Edinburgh (latitude: 56°N).

The System for Capture of Asteroid and Meteorite Paths (SCAMP) United Kingdom, part of the FRIPON network (Colas et al., 2020) consisting of 17 cameras located across England and Wales between 54.7°N and 50.8°N, also managed to record the aurora expansion. Though the cameras do not usually capture auroral light, as the exposure settings are optimized for bright

² https://geomag.bgs.ac.uk/data_service/space_weather/alerts/alert_2024-05-09.html

³ <https://taivaanvahti.fi/>



meteorite trails, they became visible after image processing to enhance the gain and remove noise. As the cameras take only intensity images there is no color determination. In Figure 9B, images from 10 of the SCAMP cameras are arranged in rows of descending geographical latitudinal order from Alston (53.8°N) at the top to Cowes (50.8°N) at the bottom. The images show a 180° fish-eye lens view of the visible night sky every 10 min. The peak of the aurora is clearly visible around 22:40 UT on the 10th May when all of the cameras capture auroral arcs and more complex patterns. The peak coincides with a large deviation of over 500 nT in the horizontal magnetic components at Eskdalemuir (upper panel) and Hartland (lower panel). The auroral oval expands rapidly over 30 min before retreating to higher latitudes post-midnight.

6 Extreme storms comparison

Analysis of previous severe space weather events can help understand how different features of a geomagnetic storm (such as its solar origin, timing of commencement, main phase duration) can result in varied level of impact on ground-based infrastructure. Lessons learned from past events help further develop efficient forecasting and nowcasting systems and improve current mitigation strategies employed by industry. In this section we present a comparison, locally in the United Kingdom, between the May 2024 storm and other extensively studied events: September 2017 (Dimmock et al., 2019; Piersanti et al., 2019; Alfonsi et al., 2021), October 2003 (Pulkkinen et al., 2005; Thomson et al., 2005; Xue et al.,

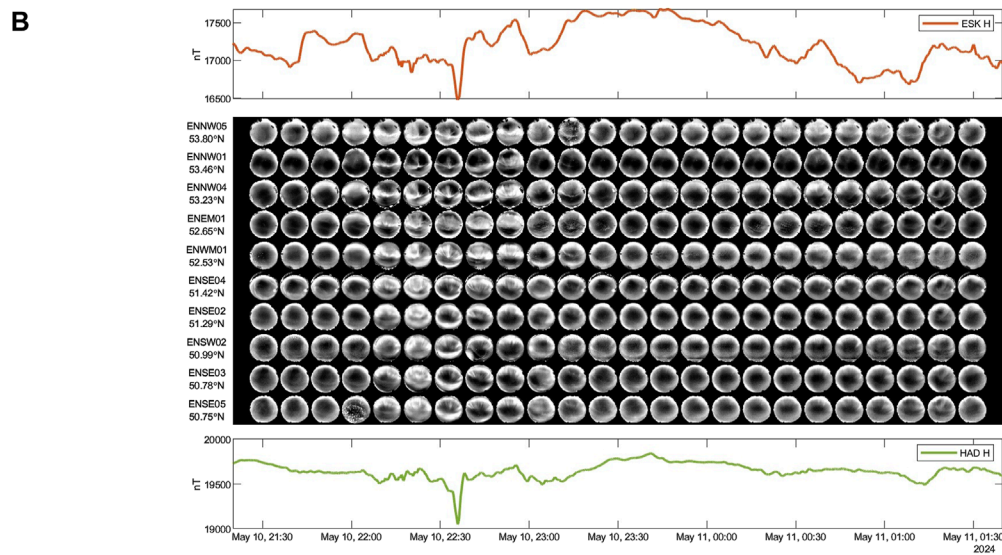
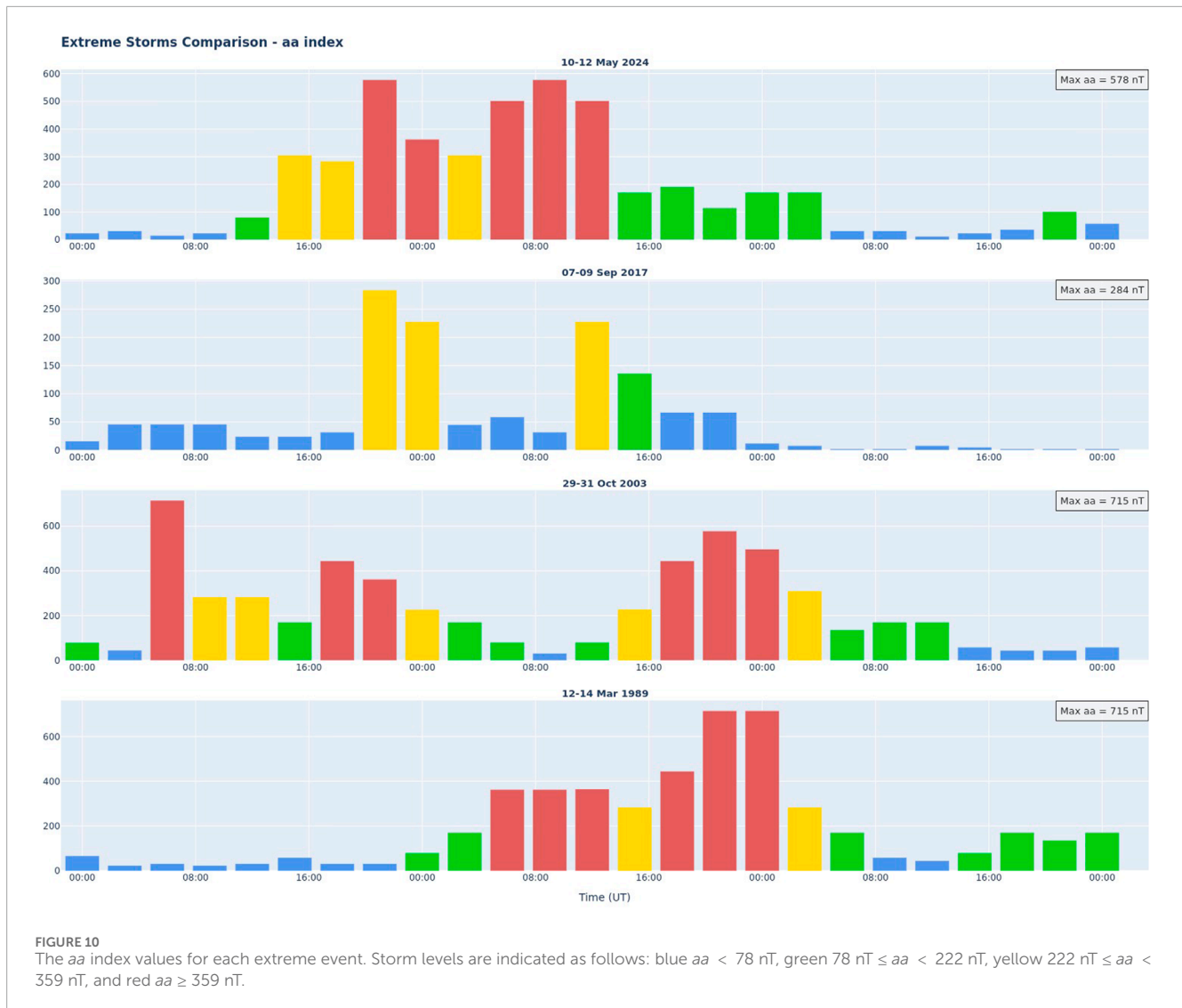


FIGURE 9
(A) Photos of the northern lights on 10–11 May 2024 provided by BGS colleagues. Credit: Natalia Gomez-Perez (top left), Victoria Thompson (top right, bottom left), Chris Turbitt (bottom right). **(B)** SCAMP-FRIPON 10-min images from meteorite cameras captured at Alston (top row) in order of descending latitude to Cowes (bottom row) from 21:30 to 01:30 UT on the 10–11 May 2024 showing the rapid expansion of the auroral oval. The time series show the Eskdalemuir (upper) and Hartland (lower) geomagnetic observations of the horizontal component with 1-s cadence.



2023), March 1989 (Allen et al., 1989; Shirochkov et al., 2015; Boteler, 2019; Love et al., 2022), as well as the Carrington storm of September 1859 (Tsurutani et al., 2003; Giegengack, 2015; Hayakawa et al., 2023; Beggan et al., 2024; Thomas et al., 2024). We also present the overall storm ranking on a global scale in terms of the aa^* index, dating back to 1868.

6.1 Geomagnetic aa and aa^* index

The K index is a measure of disturbance in the horizontal magnetic field component over a 3-h interval, ranging from 0 to 9, with 4 or less being quiet and 5 or more indicating a storm (Bartels et al., 1939). The 3-hourly aa index is derived from the K indices from two approximately antipodal observatories and has units of nT (Mayaud, 1972; Mayaud, 1980). Over the years, to ensure data quality and continuity, the observatories used in the index derivation, were replaced with new observatories. These often operated concurrently to minimize any non-homogeneity of the time series (Clilverd et al., 2002). At present, the observatories

are Hartland in the United Kingdom, operated by the BGS, and Canberra in Australia, operated by Geoscience Australia. The index extends further back in time (to 1868) than any other planetary index time series. The aa indices are estimated in real-time by BGS with final values available weekly following the publication of the Canberra K indices. Definitive aa are published by the International Service for Geomagnetic Indices (ISGI)⁴, a service of the International Association of Geomagnetism and Aeronomy (IAGA). The aa indices for each geomagnetic storm are presented in Figure 10. Note that values for 2024 are still provisional at the time of writing, though will not change significantly. The aa index usually follows the same pattern as the K index, however it shows more variation throughout each storm since it is not restricted to a specific maximum value. Therefore, different peak aa indices can be observed depending on the level of geomagnetic disturbance of each event, which arguably serves as a better indicator of space weather severity. However, unlike ap , the aa index is limited as a true

⁴ <https://isgi.unistra.fr/>

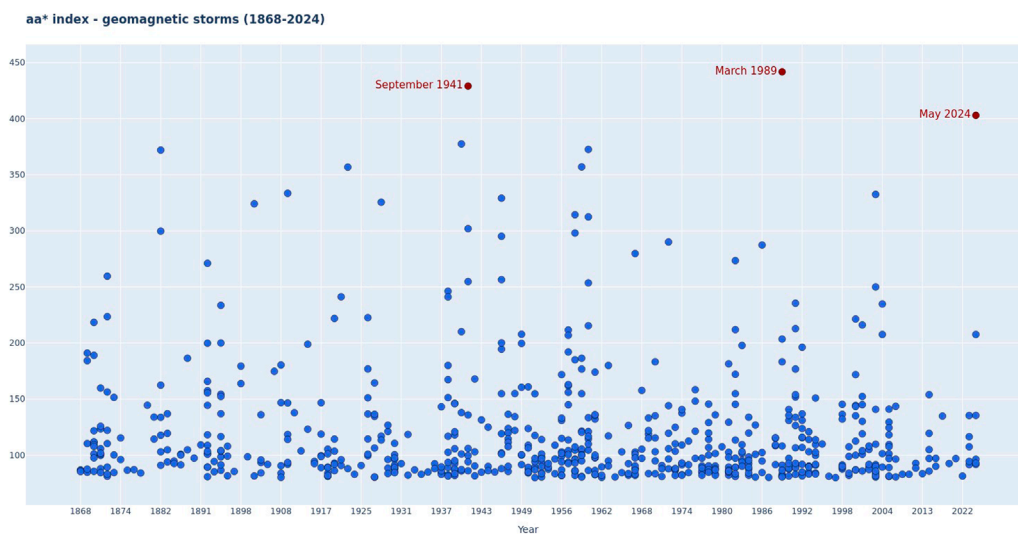


FIGURE 11
The aa^* index values showing extreme geomagnetic storms ($aa^* \geq 80$ nT) from 1868 to 2024. Top three events are highlighted and labeled in red.

planetary activity measure, as it is based on only two positions on the Earth's surface. For statistical studies and to estimate global levels, this can be overcome by computing eight-point running means to generate 24-h average values of aa , denoted aa^* , while retaining the same 3-h resolution (Clilverd et al., 2002).

In the afternoon (all times are provided in UT) of the 10 May 2024 the aa index gradually increased to storm levels (> 222 nT), reaching the maximum value of 578 nT in the late evening. The geomagnetic disturbance continued throughout the morning and early afternoon of the 11th May, before declining to moderately elevated levels in the late afternoon. On the morning of the 12th May, geomagnetic activity fully returned to quiet conditions. In comparison, despite being one of the more extreme events of solar cycle 24, the September 2017 storm only reached maximum of $aa = 284$ nT. The Halloween 2003 storm is one of the most extensively studied space weather events of the 21st century. The SSC arrived around 06:00 UT on the 29th October and caused a rapid increase in the aa index, reaching maximum value of 715 nT soon after the shock. Another gradual enhancement in geomagnetic activity followed in the afternoon of the 30th October, with $aa > 400$ nT, before slowly returning to quiet conditions in the morning of the 31st. The geomagnetic storm of March 1989 remains one of the most significant space weather events in modern era, strongly affecting the northern regions of North America and Europe. Geomagnetic disturbances caused damage to the transformers in the Hydro-Québec power system, leaving millions of people without power for several hours (Bolduc, 2002). Unlike the October 2003, this storm started with steady increase in the aa index from midnight 13 March 1989, reaching high values for over 24 h, with maximum $aa = 715$ nT during late evening.

The aa^* geomagnetic index is computed back to 1868 and used to identify the most extreme events. Figure 11 shows the index values for any events that exceeded 80 nT throughout the entire aa^* time series. Three most extreme storms are highlighted in red, all of which reached over 400 nT. In comparison with other historical events, the

'Gannon' 2024 storm ranks third (403 nT) in terms of the aa^* index, after March 1989 (442 nT) and September 1941 (429 nT). Note that the Carrington 1859 storm cannot be included as it predates the start of the aa and aa^* time series.

6.2 Horizontal magnetic field variation

The horizontal magnetic field intensity variation (H) and more specifically its rate of change minute-to-minute (ΔH) can be used as a proxy for GIC magnitude, since large geoelectric fields are associated with high ΔH values (Fiori et al., 2014). The H (nT) and ΔH (nT/min) are defined by Equations 1 and 2:

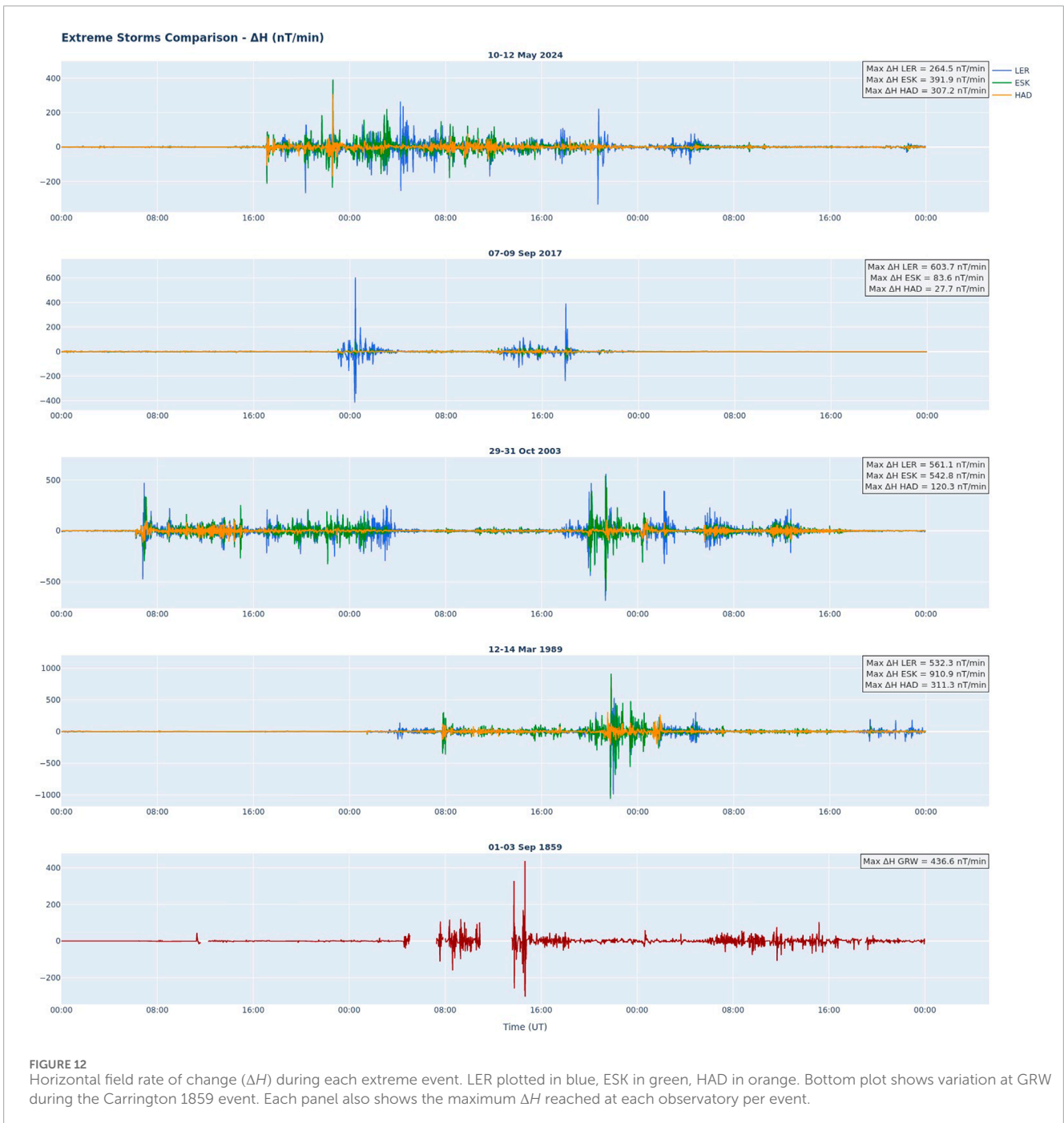
$$H = \sqrt{B_x^2 + B_y^2} \quad (1)$$

$$\Delta H = \frac{H(t + dt) - H(t)}{dt} \quad (2)$$

where B_x and B_y correspond to the northward and eastward surface magnetic field, respectively and dt is equal to 1 min which corresponds to the magnetic field data cadence. For the direct magnetic field comparison, digitized H measurements from Greenwich (GRW) observatory in London during the Carrington September 1859 storm (Beggan et al., 2024) are also included in the analysis.

Figure 12 shows the calculated rate of change of the external H component for each event from all three United Kingdom observatories, as well as GRW during the Carrington storm (bottom plot). Each panel also contains information about the maximum values of ΔH reached at United Kingdom observatories during respective storm. Similar to the index analysis in previous subsection, the external field comparison indicates the differences in duration and severity of each event, but it also illustrates their latitudinal extent.

The geomagnetic storms of May 2024 and March 1989 are similar in terms of duration and morphology, both lasting more than



30 h with continuous progression and largest disturbances observed during the main phase. The September 2017 and Halloween 2003 events include two distinct enhancements in geomagnetic activity, several hours apart, with strong variations seen in the early stages of each storm. Higher latitudes tend to be more susceptible to space weather effects due to the expansion of the auroral oval and substorm occurrence (Skone and De Jong, 2000; Pirjola, 2005). The analysis of the geographical latitudinal differences (for United Kingdom observatories, around 10°) captures the spread of the auroral oval, indicating the intensity of the solar wind-magnetosphere interactions. The maximum ΔH values show that during May 2024

and March 1989, ESK observatory (55.3°N) measured the highest variations. Peak ΔH for ESK during March 1989 reached over 900 nT/min, in contrast to 532 nT/min at LER (60°N). For the other two storms, we see the largest variations observed at LER, decreasing with latitude. The September 2017 storm was particularly intense only at the highest latitude, where ΔH reached over 600 nT/min.

There remain large uncertainties regarding the fidelity and quality of GRW data during the Carrington event due to the unknown response of the historic instrumentation compared to modern sensors. There are also long gaps throughout the main phase of the storm where the field intensity was off the scale of the chart.

The horizontal field at GRW shows a sharp response to the solar flare around 11:15 UT on the 1st September, followed by a clear SSC in the morning of the 2nd September. It is important to note that it is not possible to determine the true maximum ΔH because of the data gaps when (we assume) the variations were greatest. Therefore the assessment of the storm severity should be treated as a minimum estimate. Stewart (1861) suggests a rate of change at over 700 nT/min from an off-page visual observations at the time, which is more than twice that of the March 1989 peak values in HAD, similar latitude to GRW.

7 Discussion and summary

Geomagnetic field measurements from the three United Kingdom geomagnetic observatories and three variometers sites indicate a clear SSC at 17:08 UT on 10 May 2024. Magnetic field values show a strong, long-lasting response to the main phase of the ‘Gannon’ storm before gradual return to quiet activity levels around 12:00 UT on the 12th May, over 30 h after the shock arrival. The geoelectric field data from two available monitors also follows a similar pattern, reaching maximum values in the early stages of the storm progression on the 10th May, at 22:36 UT. Modelled electric fields across the United Kingdom, based on the MT measurements and transfer functions, also show the strongest response at that time, with the most extreme values estimated in the northwest of Scotland and northern England, and higher variations observed in the East-West component.

Based on the real-time magnetic field measurements and the modelled geoelectric fields, GICs flowing in the United Kingdom high-voltage power network during the ‘Gannon’ event were estimated. Peak GIC values were observed in long transmission lines and corner nodes, usually situated at coast lines. Several substations in southwest and east-central England, as well as northern Wales showed GICs exceeding 60 A, with largest value of 63 A at a substation near Norwich (East coast of England). Since the leading edge of the auroral oval moved quickly to lower latitudes, GICs in the Scottish part of the network (above 55°N) did not show any extreme enhancements during the storm.

The comparison, locally in the United Kingdom, with previous extreme events, namely, September 2017, October 2003, March 1989 and September 1859 storms, shows how diverse a storm progression can be, and how different aspects of storm morphology can result in varied impacts. Globally, the ‘Gannon’ event ranks third in severity since 1868, based solely on the aa^* index, and was more in line with a 1-in-30 years storm based on rate of change of the magnetic field. Analysis of other indices suggests that its magnitude puts it closer to a 1-in-12.5 years event, and a 1-in-41 years event based on duration (Elvidge and Themens, 2025). However, the magnetic field rate of change was not sufficiently large or rapid to cause any reported issues or damage to technical infrastructure in the United Kingdom.

Data availability statement

The raw data supporting the conclusions of this article will be made available by the authors, without undue reservation.

Author contributions

EL: Data curation, Investigation, Visualization, Writing–original draft, Writing–review and editing. CB: Data curation, Investigation, Visualization, Writing–original draft, Writing–review and editing. GR: Data curation, Methodology, Visualization, Writing–review and editing. SR: Data curation, Investigation, Visualization, Writing–review and editing. VT: Data curation, Visualization, Writing–review and editing. EC: Data curation, Investigation, Methodology, Writing–review and editing. LO: Data curation, Visualization, Writing–review and editing. JH: Data curation, Investigation, Methodology, Software, Visualization, Writing–original draft, Writing–review and editing. AS: Visualization, Writing–review and editing, Data curation.

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Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

The handling editor Walach declared a past co-authorship with the author LO.

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