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Spatio-temporal evolution of pulsating aurora observed using a ground-based imagers

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We investigated the pulsating aurora observed on 7 January 2014, by a narrow field-of-view (FOV) high-time resolution ground-based white-light imager and all-sky low-time-resolution imager operated at Poker Flat, AK (geographic: 65.1°N, 147.4°W). The pulsating aurora showed very notable characteristics, such as frequency drift in their pulsation with time and drifting of the entire pulsating auroral structure in space. We find that (i) the entire pulsating auroral patch was observed to drift northward at a velocity of approximately 76 m/s, which aligns closely with the local convection velocities obtained from Super Dual Auroral Radar Network (SuperDARN) data, consistent with the idea that the patch motion is primarily due to E×B convection. (ii) The duration of persistence for each pulsation in the pulsating aurora is found to be ~1 s (iii) The auroral pulsation frequency abruptly increases from ~0.0625 Hz to ~0.5 Hz, closely aligning with the broadening of the frequency band observed at the Dawson (DAWS) ground magnetometer location. Wavelet analysis of DAWS magnetic field data, recorded at similar magnetic local time (MLT) and L-values, reveals the presence of Pc-1 geomagnetic pulsation (Pc-1). This connection suggests that the drift in auroral pulsation frequency may be driven by the evolution of Pc-1 waves, which are influenced by changes in the local plasma environment. The broadening of the frequency band may indicate dynamic variations in magnetospheric ion composition or plasma density. This interplay underscores the role of Pc-1 waves and magnetospheric dynamics in determining the auroral pulsation characteristics.

KEYWORDS

pulsating aurora, ground imagers, pulsation drift, inner magnetosphere, Pc-1 waves

1 Introduction

Pulsating auroras (PAs) are unique optical phenomena that are characterized by a quasiperiodic rise and fall in auroral intensities over a period of 2-20 s with a typical scale size of ~10–200 km (Royrvik and Davis, 1977; Yamamoto, 1988). These are commonly seen in polar regions during the post-midnight hours, close to the equatorward boundary of the auroral oval. Their occurrence is affected by geomagnetic activity, which can expand their local time range, occasionally extending to the dayside during periods of intense

geomagnetic disturbances (Royrvik and Davis, 1977; Oguti et al., 1981; Jones et al., 2011). PAs typically occur as a series of pulsations lasting approximately 30 s to a few minutes. Each pulsating auroral structure occurs independently in terms of location, timing, lifespan, velocity, and phase, without synchronization or connection to a nearby pulsating auroral structure (Scourfield et al., 1972). PAs occur most frequently during the recovery phases of the substorm and in the morning sector. PAs are a type of diffuse aurora that is caused by precipitation of energetic particles (Smith et al., 1980; McEwen et al., 1981) from the plasma sheet without additional acceleration in the low-altitude magnetosphere (Davidson, 1986a; Davidson, 1986b; Huang et al., 1990). It is considered to be quasiperiodic because, within a single pulsation train, the time intervals between consecutive maxima are not uniform, often exhibiting significant variability from one pulsation to another. Based on rocket and low-altitude spacecraft measurements made simultaneously with auroral imaging, such auroral pulsations are known to be caused by quasi-periodic precipitation of electrons with energies of tens of keV to ~ 100s of keV into the upper atmosphere (Johnstone, 1978; Sandahl et al., 1980; Samara et al., 2010; Miyoshi et al., 2010). Miyoshi et al. (2015) proposed a model to explain that the PAs could also be the optical manifestation of subrelativistic electron precipitation into the middle atmosphere. The light intensity of pulsations is often described by rapid rise and decay times relative to the pulse duration, creating the appearance of the aurora switching on and off. In general, there is notable variability in the shape of pulsation trains, with on-time and off-time durations differing between pulses, with larger variations typically observed in off-time (Davidson and Chiu, 1991).

Several mechanisms have been proposed to explain the precipitation of energetic electrons associated with PA. Modulation of the rate of pitch angle scattering due to wave-particle interactions near the equator causes the pulsations (Nishimura et al., 2010; Nishimura et al., 2011; Coroniti and Kennel, 1970). So, the rate of change of pitch angle scattering could directly affect the pulsation frequency. Considering that the characteristic energy of the PA electrons is of the order of 100s of keV, their cyclotron resonance with whistler mode chorus waves, lower band chorus (LBC) waves especially, is a plausible candidate to cause such high energy electron precipitations (e.g., Thorne et al. (2010), Nishimura et al. (2010), Nishiyama et al. (2011) showed that auroral intensity from all-sky imager (ASI) data closely correlates with chorus wave amplitude measurements from THEMIS spacecraft, confirming that 100 eV to 10 keV electrons resonate with electron cyclotron harmonic and whistler mode waves. In addition to chorus waves, electromagnetic ion cyclotron (EMIC) waves are plasma waves that can scatter energetic protons and electrons into Earth's atmosphere (e.g., Cornwall, 1965; Summers et al., 1998; Halford et al., 2016; Yahnin et al., 2021). They are typically generated near the magnetic equator in the inner magnetosphere by anisotropic ions, primarily at lower L-shells, and are bounded by ion gyrofrequencies into three primary bands: hydrogen, helium, and oxygen (e.g., Blum et al., 2012; Gary et al., 1995). These waves are typically spatially confined in radial extent (e.g., L-shell) in the magnetosphere but can extend in magnetic local time (MLT) (e.g., Mann et al. (2014), Blum et al. (2017). While H^+ -band and He^+ -band EMIC waves are more frequently observed, O⁺-band EMIC waves are not rare, particularly during geomagnetically active periods (Saikin et al., 2015; Usanova et al., 2016; Usanova et al., 2018). These waves have been reported in the outer plasmasphere at L = 2-5 from Van Allen Probes Electric and Magnetic Field Instrument Suite and Integrated Science and Electric Fields and Waves data (Yu et al., 2015).

Pickett et al. (2010) observed EMIC rising tone emissions in association with Pc-1 waves. The characteristic of Pc-1 waves, observed by the Cluster spacecraft on 30 March 2002, are consistent with their identification as EMIC triggered chorus emissions. These Pc-1-triggered chorus waves could potentially cause electron scattering into the loss cone, leading to the generation of PAs. Ultra-low frequency (ULF) waves, with periods ranging from tens of seconds to a few minutes, can influence the generation of PA by modulating whistler-mode chorus waves. Li et al. (2011) demonstrated that these waves affect chorus wave growth by altering electron density or the anisotropy of the resonant electrons. This periodic modulation aligns with the pulsation period of typical PA. Li et al. (2011) highlighted that density variations, whether induced by ULF modulations or other mechanisms, play a critical role in driving PA.

The dynamics of PAs often exhibit a clear post-midnight drift pattern, moving eastward at speeds ranging from a few hundred meters per second to a few kilometers per second. This drift is dawnward after midnight and shifts to duskward behavior before midnight (Nakamura and Oguti, 1987). Importantly, the speed of patch drift has been reported to align closely with the $E \times B$ drift speed (Nakamura and Oguti, 1987), indicating that the movement of these auroras is influenced by the distribution of cold plasma in the magnetosphere more than the eastward magnetic drift of energetic electrons. The eastward drift observed before midnight suggests that the characteristics of pulsating auroral structures are shaped by magnetospheric conditions, emphasizing the role of cold plasma dynamics in auroral behavior (Yang et al., 2015). In pulsating and propagating auroras, precipitation is likely driven by pitch angle scattering, with poleward propagation reflecting the outward movement of strong scattering regions in the loss cone distribution. At the ionospheric level, this propagation occurs at 10-30 km/s, corresponding to hundreds to thousands of km/s in the distant magnetosphere (Oguti and Watanabe, 1976).

Estimating the size of pulsating auroral patches provides insight into the regions of chorus intensification near the magnetic equator. Previous research has suggested that the transverse scale of individual chorus elements varies significantly, with estimates of approximately 100 km at L ~ 4.4 (Santolík et al., 2004) and about 2,800–3,200 km at L ~ 11 (Agapitov et al., 2010). However, *in situ* observations from limited satellite data may not accurately capture the full spatial extent of these chorus enhancements. The modulation region's association with the magnetic equator has been established through time-of-flight analyses and magnetically conjugate PA observations, suggesting that lower-band chorus waves are crucial for the precipitation of electrons responsible for PAs (Nishimura et al., 2011).

Previous studies, such as Samara et al. (2017), have provided valuable insights into the PA observed on 7 January 2014, highlighting the presence of high-frequency peaks in the PA and linking these features to electron dynamics, including electrons that bounce back and forth between the two hemispheres. Khazanov et al. (2017), Khazanov et al. (2021b) further suggested that the auroral emissions are not solely driven by magnetospheric

processes, indicating a more complex interaction. These studies have contributed to the understanding of PAs, yet the factors influencing the spatial and temporal drift of their pulsation frequencies remain poorly understood for this event. Our study builds upon these previous works by providing more detailed spatial and temporal characterization of the pulsating aurora through highresolution ground-based imaging data from Poker Flat, AK. We incorporate additional *in situ* observations from the ground-based magnetometer data from Dawson, which share similar magnetic local time (MLT) and L-values, to offer new insights into the role of Pc-1 waves, in modulating the pulsation frequency. Section 2 presents the event and datasets, while Section 3 provides details on the observations and analysis. Sections 4 and Section 5 discuss the results, and the conclusions are summarized in the final section.

2 Data and methodology

The PA event on 7 January 2014, was analyzed using the Multispectral Observatory Of Sensitive EMCCD (MOOSE) imagers, equipped with Andor Ixon DU-888 EMCCD (Electron Multiplying Charge Coupled Device) cameras. Two white-light imagers with varying fields-of-view (FOV) were deployed at Poker Flat, AK (geographic: 65.1°N, 147.4°W; geomagnetic: 65.7° N, 96.6°W). The ASI operated at 3.3 frames per second (512×512 resolution, ~300 ms exposure) with 2×2 pixel binning, while the narrow FOV imager, with a 4° FOV, captured at 56 frames per second $(128 \times 128 \text{ resolution}, 16 \text{ ms exposure})$ using 1×1 pixel binning. This configuration was selected to balance temporal resolution at large scales (all-sky) with high temporal and spatial resolution at small scales (narrow FOV). The narrow FOV imager, pointed at the magnetic zenith near the center of the all-sky images, was used without a filter to maximize sensitivity to prompt emissions. Previous work by Samara et al. (2012) indicated that a BG3 filter, which blocks emission at 557.7 nm and 630.0 nm, does not significantly improve the visibility of rapid pulsations. The system provided angular resolutions of 0.054° per pixel (~50 m per pixel) for the narrow FOV imager and 400-600 m per pixel for the ASI near the zenith, assuming auroral emissions occurred at an altitude of 100 km. More details on MOOSE EMCCD can be found in Michell et al. (2014), Michell and Samara (2015).

The Canadian Array for Realtime InvestigationS of Magnetic Activity (CARISMA) is a network of ground-based magnetometers deployed across North America Mann et al. (2008). The array includes induction coil magnetometers that monitors 3-D vector magnetic fields and fluctuations at the Earth's surface, which are sensitive to EMIC wave activity. These magnetometers have a resolution of <0.2 pT/Hz^{1/2} at 1 Hz and cover *L*-shells of *L* = 3.6-6, spanning roughly 4 h of MLT. The magnetometers provide data at 20 samples per second. The closest CARISMA magnetometer station to the ground-based auroral imager is near Dawson city (DAWS). The DAWS data follows the geographic coordinate system, and the H-component corresponds to the horizontal geomagnetic direction.

The Super Dual Auroral Radar Network (SuperDARN) is utilized to measure ionospheric plasma convection by detecting Doppler-shifted radio signals reflected from field-aligned irregularities in the F-region ionosphere (Greenwald et al., 1995; Chisham et al., 2007). In this study, we use data from the Prince George (PGR) radar, located at (53.9812°N, 122.5920°W), near the ground-based imagers in British Columbia, Canada to analyze the plasma flow in the region near the ground-based auroral imagers. The radar operates by transmitting high-frequency radio waves and measuring the Doppler shift of returned signals to determine line-of-sight plasma velocities. For this analysis, we focus on beams 1, 2, 12, 14, and 28 of the PGR radar, as it aligns with the observed auroral structures and provides a clear view of the plasma convection dynamics. The beams having a significant number of data points during this interval have been selected. The use of a limited number of beam ensures consistency in tracking velocity variations along a fixed direction, minimizing uncertainties associated with multi-beam averaging (Ruohoniemi and Baker, 1998).

3 Observation and analysis

The solar wind and geomagnetic conditions during 1,600–1,800 UT on 7 January 2014 are shown in Figure 1. The panels represent 1-min resolution data from the OMNI database (King and Papitashvili, 2005) for (a) solar wind velocity (Vsw), (b) the z-component of the interplanetary magnetic field (IMF Bz), and (c) solar wind dynamic pressure (Psw). The geomagnetic activity indices shown in (d) AL-index and (e) AU-index are obtained from Davis and Sugiura (1966). A weak substorm with AL~-500 nT was in action with a weak southward IMF ~-4 nT. The yellow shaded region highlights the time period from 16:26 to 16:28 UT on 7 January 2014, during which a strong PA was observed by ground-based auroral imagers at Poker Flat, AK.

Figure 2 presents the auroral observations that combine the white-light images from both (a) all-sky and (b) narrow FOV cameras, captured simultaneously during the pulsation-on phase of the study. The keogram cut location is marked as a colored vertical line in both images. Panel (c-f) represents the snapshots of the PAs at different times, taken from the narrow FOV imager. This study investigates the movement of auroral structures to capture their spatial evolution. At the same time, it analyzes rapid temporal variations using a narrow FOV with high frame rates.

Figure 3 provides a multiscale view of a 3-min keogram of the PA observed at Poker Flat, AK, on 7 January 2014. Panel (a) displays the PA captured with a 0.33 s exposure using the all-sky white-light imager, showing multiple patch-like features near the zenith in the north-south keogram. Panel (b) presents a zoomed section between the two black dashed horizontal lines in panel (a), focusing on the region between 109 and 129 km. The black dashed horizontal lines in panel (c). Panel (c) depicts a high-time-resolution keogram from the narrow FOV white light imager at the same location, revealing quasi-periodic oscillations beginning at 16:26:20 UT.

We observe three key dynamic features that evolve in both space and time. The first and most prominent feature is the northward movement of the pulsating auroral patch. The second is the duration of persistence of each individual pulse within the PA, and the third is the variation in the PA frequency over time. The spatial location of the maximum intensity and a slight time delay in pulsating aurora peaks are observed in Figures 3b,c. The observed differences between Figures 3b,c can be attributed



and **(e)** AU-index for the period of between 1,600–1,800 UT on 7 January 2014. The shaded yellow region indicates the duration of the PA observed in the ground-based auroral imager at Poker Flat, AK.

to several factors related to instrumental characteristics, viewing geometry, and optical distortions. The temporal resolution disparity between the two imagers plays a key role, as the 180° FOV imager operates at 3.3 Hz, while the narrow 4° FOV imager runs at 56 Hz. The higher frame rate of the narrow FOV imager allows it to capture rapid variations in auroral intensity with greater precision, whereas the lower frame rate of the wide FOV imager may introduce a small lag in detecting peak intensities, leading to the observed ~1 s delay in pulsating aurora variations. Additionally, the spatial resolution and FOV differences contribute to discrepancies in intensity localization. The wide FOV imager, covering a much larger area, has coarser spatial resolution and integrates emissions from a broader region, potentially averaging out finer details that are better resolved by the narrow FOV imager. The exposure time difference also plays a role, with the wide FOV imager having a longer exposure per frame compared to the high-frame-rate narrow FOV imager, which could smooth out intensity variations over time. Moreover, optical distortions in the narrow FOV imager due to lens curvature may introduce small spatial shifts in the recorded intensity patterns, affecting the precise location of peak auroral emissions.

To calculate the drift velocity of the PA, we employ an algorithm that systematically tracks pulsation-on patches in the imager. First, we identify these patches by detecting pulsations with intensities exceeding 700 counts, marked by blue dots in Figure 4 (top). Next, we determine the maximum intensity for each pulsation at each time instant, represented by magenta dots. To refine the analysis, we extract the maximum values within 0.1-s intervals and identify the corresponding pixel locations at peak intensity for each time step. These values, shown as red data points in Figure 4 (bottom), are then converted into distance measurements. A regression line (blue) is fitted to these data points, and its slope provides a drift velocity of 76.41 m/s, indicating northward motion of the pulsating patch. Overall, the PA patch predominantly drifts northeast, with a weaker eastward component. This eastward drift is further evident in Figures 2c–f, where a gradual shift in the patch's position over time can be observed.

Figure 5 (left) provides a detailed visualization of the observational setup, highlighting key regions and instruments involved in the study. The map projection displays the FOV of the PGR radar in red, overlaid with the FOV of the MOOSE ASIs at PFRR in blue, allowing for a comparison of the areas covered by both instruments. This layout is crucial for understanding the spatial relationship between the two FOVs and how they complement each other. The map also incorporates the MLT, which is essential for contextualizing the timing of observations in relation to the Earth's magnetic field. Additionally, the figure marks the location of the DAWS (green), a vital ground-based station, providing further context for the observations. Together, these elements offer a comprehensive understanding of the observational context and the geographic positioning of key instruments in the study. Figure 5 (right) represents the velocity versus time plot recorded by beam 1, 2, 12, 14 and 28 of the Prince George (PGR) radar between 16:27-16:28 UT. The beams having a significant number of data points have been selected. Negative values of the velocity indicate westward plasma flow, while positive values represent eastward plasma flow. During the observed period, the velocity measurements show a strong eastward flow aligning well with the movement of the PAs shown in Figure 3c. The confidence levels of these measurements are determined by the signal-to-noise ratio (SNR) and the fitting accuracy of the radar echoes, both of which fall within the standard acceptable range for SuperDARN measurements. This agreement reinforces the reliability of the velocity measurements and supports the interpretation of the ionospheric plasma motion during the observed interval.

Figure 6 (top) displays the keogram intercept taken ~3.25 km over a 2-min interval from Figure 3c. The green line represents the intensity counts, while the red curve illustrates the smoothed version of this data, which was obtained by applying a low-pass filter with the cut-off frequency of 2 Hz and removing the background counts. Blue circles mark the identified peaks within the data, and magenta circles indicate the Full Width at Half Maximum (FWHM) of these peaks. The FWHM for each pulsation-on interval helps us to investigate the magnetospheric counterpart of the observed PAs. The time duration obtained from the FWHM gives the duration of persistence (τ) of on-time for each pulsation of aurora. Figure 6 presents the histogram for the duration of persistence of each pulsation observed during 16:26:00 UT-16:28:00 UT. The histogram clearly shows that the duration of pulsation on-times varies from pulse to pulse, with the majority having a τ value near 1 s. A couple of events with $\tau > 2$ s come from two or more pulsation peaks close to each other. Figure 6 (top panel) presents the FWHM for each pulsation, with the τ > 5 pulsation resulting from overlapping peak intensities observed between 16:27:28 and 16:27:32.3 UT.



(a) Images of the PA event that captured at 16:26:21 UT on 7 January 2014. All sky image with 180° FOV, and (b) 4° FOV is shown. The colored vertical line in both images represents the location of the keogram in the north-south direction. (c–f) represents the motion of the PAs with time from the narrow FOV imagers.

To analyze the temporal evolution in the pulsation frequency, we present Figure 7a1 which is same as the green curve in Figure 6 (top). To investigate the temporal evolution of pulsations we have performed wavelet transform of the data by following the methodology of Torrence and Compo (1998). Figure 7a2 shows the continuous wavelet transform of the auroral intensity data after applying a Butterworth band-pass filter of 0.1–8 Hz. During 16:26:20 UT to 16:26:50 UT, the PA spans a broad frequency range, approximately from 0.0625 Hz to 0.5 Hz. However, during 16:26:50 UT-16:27:30 UT the frequency of the pulsation abruptly narrows around 0.5 Hz, with no low-frequency (<0.25 Hz) components present.

Figure 7b1 displays the D-component of the ground magnetic field data from the DAWS station, while Figure 7b2 presents the corresponding wavelet after applying a band-pass filter ranging from 0.1 to 0.9 Hz. During 16:26:50–16:27:00 UT interval, a significant wave power enhancement is observed with a broad band of frequncies within the 0.125–0.5 Hz range. In addition to this, we also observe strong Pc-1 waves with a dominant frequency of 4 Hz between 16:27:15 and 16:27:30 UT, which does not appear to have any significant impact on the observed PA. This alignment suggests a potential connection between the Pc-1 waves and the observed features in the auroral intensity curve.

4 Discussion

Our observations of the pulsating aurora on 7 January 2014 provide valuable insights into the spatio-temporal evolution of the pulsating aurora and their characteristics. The frequency drift and spatial drift observed in the pulsating aurora are suggestive of the magnetospheric dynamics such as wave-particle interactions. The combination of ASI and narrow FOV imagers having high time and space resolution helps distinguish unique features of auroral forms, identifying their locations and motion patterns, and correlating them with the magnetospheric processes. For example, the PAs observed on 7 January 2014 is spatially localized until 16:27:00 UT as shown in Figures 3b,c. After this, a rapid movement of the pulsating structure appears in the northeast direction through both ASI and narrow FOV imagers Figures 2c-f. The PA events examined in this study occur during the recovery phase of the substorm, aligning with the findings of Partamies et al. (2019), which reported that a majority (64%) of events with decreasing patch sizes also take place during this phase. The spatial drift motion of the PAs, with speeds of thousands of m/s, cannot be explained by obliquely propagating chorus waves, which typically cause the 3±1 Hz modulation due to their perpendicular speed exceeding the Alfven speed. Instead, slow-mode Alfven waves could be a potential source of modulation



responsible for generating these rapid motions Fukuda et al. (2016), Chaston et al. (2002). Semeter et al. (2008) captured the horizontal distributions and temporal variations of discrete auroras, particularly in the context of Alfvén wave propagation. Their study identified that the variations in auroral intensity and structures are linked to Alfvén wave dynamics and quasi-electrostatic parallel potential drops.

We estimate that the FWHM for most pulsations is close to 1 s. The duration of the pulsation observed in this study closely aligns with previous findings on the precipitation time of electrons with 5 keV-500 keV at L = 6 in the dipole magnetic field model used by Saito et al. (2012). They found that the precipitated electrons resonate with whistler waves with $\omega = 0.4^* | \Omega_{e,eq} |$, where ω is the wave angular frequency, and $\Omega_{e,eq}$ is the electron cyclotron frequency at the equator where whistler chorus elements are launched in the model. Nishimura et al. (2010), Nishimura et al. (2011) discovered a one-to-one correspondence between in situ chorus intensity observed near the equator and pulsating auroral luminosity at the ionospheric footprint. Simultaneous groundbased observations of VLF waves and optical images also shows correspondence between them (Tsuruda et al., 1981; Hansen and Scourfield, 1990; Tagirov et al., 1998; Ozaki et al., 2012). Humberset et al. (2016) found that the on-time of pulsations did not show a strong correlation with the maximum intensity of the patches, suggesting that the energy deposition in a single pulsation fluctuates and is not constant throughout the lifetime of the patch. They proposed that the highest-energy electrons reach the ionosphere first, creating an energy dispersion pattern that governs the temporal dynamics of PAs. Scourfield et al. (1983) demonstrated that pulsating auroral forms drift with the same $E \times B$ velocity as the background cold plasma, as confirmed by observations from the Scandinavian Twin Auroral Radar Experiment (STARE), which monitors electron flow and ionospheric plasma drifts.

During 16:26:50–16:27:00 UT, the PAs exhibited a distinct frequency drift from 0.0625 Hz to 0.5 Hz (Figure 7a2). This frequency increase coincided with an intensification and broadening of wave power in the D-component of the magnetic field, which falls within the Pc-1 range. Since EMIC waves are a key source of Pc-1 waves and are typically generated near the magnetic equator (Sakaguchi et al., 2013; Pickett et al., 2010; Nomura et al., 2016), understanding their role in modulating PA characteristics is important. Unfortunately, the absence of conjugate satellite observations for this event limits our ability to determine the exact location and mechanism of wave-particle interactions and their propagation in the magnetosphere. Investigating these processes remains an important subject for future studies.

Although the white-light auroral imagers used in this study do not distinguish between proton and electron precipitation, the observed Pc-1 activity suggests a possible connection between the pulsating aurora and proton precipitation (Sakaguchi et al., 2008; Nomura et al., 2012; Yahnin et al., 2007; Yahnin et al., 2016). Since PAs exhibit quasi-periodic intensity variations on timescales of seconds (Yamamoto, 1988; Nishiyama et al., 2014), their modulation could be driven by (i) chorus wave intensity (Trakhtengerts, 1999; Li et al., 2011), (ii) electron cyclotron harmonic waves (Meredith et al., 2009; Liang et al., 2010), or (iii) EMIC waves (Yahnin et al., 2007; Sakaguchi et al., 2008; Nomura et al., 2012). Our findings align with Nomura et al. (2016), who observed proton auroral pulsations using an all-sky imager and a ground magnetometer in Athabasca, Canada, showing a oneto-one correlation with Pc-1 rising tones. However, due to limited



observational detail, clear rising tone signatures are not evident in our event, making the association between the waves and PAs less conclusive. These rising tones intermittently scatter magnetospheric protons in the equatorial region, supporting the hypothesis that Pc-1 waves modulate auroral pulsation frequencies.

Another possible explanation involves Pc-1 wave interactions with energetic protons, which may trigger chorus emissions (Pickett et al., 2010), similar to the non-linear VLF chorus generation mechanism proposed by Omura et al. (2008), Omura et al. (2009). EMIC-triggered emissions can result from a non-linear absolute instability of L-mode EMIC waves interacting with energetic protons (Omura et al., 2010). The subsequent modulation of chorus waves could enhance electron scattering, producing high-frequency PAs. Figure 8 provides a schematic representation of the observed Pc-1 waves and their conjugate pulsating aurora signatures on the ground. Previous studies have demonstrated that wave modes, such as Pc-1 waves, can modulate the precipitation of energetic electrons and influence auroral dynamics. Specifically, Pc-1 waves have been shown to trigger chorus waves that could scatter electrons into the loss cone, leading to the generation of PAs (Pickett et al., 2010; Nishimura et al., 2010). Our findings of the modulation of the pulsating aurora's frequency drift align with the hypothesis that interactions between Pc-1 waves and chorus waves drive electron precipitation (Thorne et al., 2010; Li et al., 2011). This connection reinforces the notion that the modulation of chorus waves by Pc-1 waves could drive the observed drift in pulsation frequency, offering a deeper understanding of the mechanisms governing highly structured auroral precipitation.

Samara et al. (2017) demonstrated that the peaks in PAs are caused by electrons that bounce back and forth between the two hemispheres. Although our and Samara et al. (2017) studies analyze the same event and utilize similar imaging data, our focus differs from that of Samara et al. (2017) in exploring the role of structured primary precipitation in shaping these features. Khazanov et al. (2017), Khazanov et al. (2021b) suggested that the aurora is not only driven by pure magnetospheric processes. The magnetosphere-Ionosphere-Atmosphere (MIA) coupling of the precipitated electrons and their interplay between the northern and southern hemispheres is also an additional mechanism that proves



FIGURE 5

left: A comprehensive view of the observational setup showing the map projection of the PGR radar FOV (red), overlaid with the FOV of the MOOSE ASIs at PFRR (blue). The map provides a clear context for the observed region, showing the magnetic local time (MLT) alongside the location of DAWS (green). *right*: Velocity plots for diffreent beam numbers of the SuperDARN PGR, Canada HF radar for the time during which the PAs were observed to be drifting.



FIGURE 6

Top: A green line plot shows the keogram intercept taken at 3.25 km over a 2-min interval. The red curve represents the smoothed green line plot after applying a low-pass filter and removing the background counts. Blue circles indicate the identified peaks, while magenta circles represent their FWHM. *bottom*: Histogram of the pulsation-on time for the PAs that occurred between 16:26:00 UT-16:28:00 UT.



FIGURE 7

(a1) Auroral intensity versus time, and (a2) continuous wavelet transform of the north-south cut of the keogram centered at zenith from the narrow FOV MOOSE imager. (b1) the magnitude of the D-component of the magnetic field, and (b2) its wavelet from the DAWS station on the ground.



to be an important contributor in the formation of different kinds of aurora (Khazanov et al., 2020; Khazanov et al., 2021a). While this mechanism explains the presence of high-frequency faint peaks, our study investigates whether the observed low-frequency main peaks in PAs are driven by temporally structured wave-particle interactions.

5 Conclusion

This study has presented high-speed, narrow FOV imaging observations of auroras in the zenith associated with a weak substorm. The observations were made with a narrow FOV camera incorporating an electron-multiplying CCD (EMCCD) detector and a prompt emission filter. The high resolution of the instrument has allowed for a quantitative analysis of spatial and temporal phase coherence in the elemental pulsating auroral patch. The strength of EMCCD lies in its ability to resolve signals that are both faint and highly short-lived. To complement the bigger picture of the observed auroral emission, we also employed data analysis from the ASI at the same location.

We observed that the auroral frequency suddenly increased from approximately 0.0625 Hz–0.5 Hz, whereas the pulsating auroral patch itself spatially drifted in the northern direction at a velocity of 76.41 m/s. The persistence of each pulsation was calculated to be around 1 s, which is consistent with electron precipitation driven by chorus waves.

Wavelet analysis of the ground magnetometer data from DAWS station suggests that the observed waves in the Pc-1 band. This alignment suggests that EMIC waves in equatorial region could be influencing the auroral modulation. Thus, the results indicate a potential interaction between PAs and Pc-1 waves.

In conclusion, the observed PAs's frequency drift and spatial movement, along with the wavelet analysis of magnetometer data, suggest that the Pc-1 frequency band, play a significant role in modulating the auroral emissions. This interaction likely contributes to the observed pulsation frequency drift, providing important insights into the dynamic coupling between magnetospheric waves and ionospheric auroral processes.

Data availability statement

The raw data supporting the conclusions of this article will be made available by the authors, without undue reservation.

Author contributions

MP: Writing – original draft, Writing – review and editing. AB: Writing – review and editing, Writing – original draft. MS: Funding acquisition, Writing – review and editing. RM: Data curation, Writing – review and editing. EM: Writing – review and editing. S-BK: Writing – review and editing. LB: Writing – review and editing.

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Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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