



Analysis of Increasing Flash Flood Frequency in the Densely Urbanized Coastline of the Campi Flegrei Volcanic Area, Italy

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Esposito G, Matano F and Scepi G (2018) Analysis of Increasing Flash Flood Frequency in the Densely Urbanized Coastline of the Campi Flegrei Volcanic Area, Italy. Front. Earth Sci. 6:63. doi: 10.3389/feart.2018.00063 Climate change is affecting the frequency and severity of extreme meteorological and geo-hydrological events hitting the coastal zone of the Campi Flegrei volcanic district (southern Italy), which is prone to a wide spectrum of natural hazards, including volcanism, earthquakes, ground deformation, flash floods, landslides, and coastal erosion. This study documents the trend of flash floods affecting the town of Pozzuoli, located along the coastline of the Campi Flegrei volcanic area, during the 1970-2014 time period. An archive research together with the collection of data published on news websites and social media allowed understanding of the triggering and evolution mechanisms of flash floods in the area, as well as the most recurrent damages. Rainfall data collected by the rain gauge located within the Pozzuoli watershed were also analyzed. Results of this study show an increased frequency of flash flood events occurred in the study area. The variation in flash flood frequency is likely not related to urbanization changes, as no increase of the urban area occurred after the year 2000. The observed increase of flash flood events in recent years (2000-2014) can be reasonably ascribed to variations in the rainfall regime. Our research provides a contribution for the understanding of flash flood processes in the study area as well as relevant information for hazard and risk assessment.

Keywords: climate change, flash flood, extreme precipitations, Pozzuoli, Italy

INTRODUCTION

Changes in extreme weather and climate events, including an increase in warm temperature extremes and an increase in the number of heavy precipitation events, have been observed since about 1950 in many regions worldwide (IPCC, 2014). Specifically, changes in precipitation regimes are modifying the magnitude, frequency, and timing of floods at both regional and local scales. Blöschl et al. (2017), for example, evaluated a shift in the timing of river floods in Europe, during the past 50 years, as well as Diakakis (2014) observed in recent decades an increase in flood frequency associated with increases in extreme daily precipitation around Athens (Greece).

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Extreme floods and flash floods are the most frequent and among the costliest and deadliest natural disasters worldwide (Gaume et al., 2016). In the Mediterranean area, coastal sectors of the central-western countries (e.g., France, Italy, Spain) result the most exposed to high impact and high magnitude events because of climatic and topographic conditions (Llasat et al., 2010; Tarolli et al., 2012; Gaume et al., 2016). Here, floods result frequently in fatalities and significant economic losses due to the high concentration of urban settlements and population.

The Italian coasts were recurrently affected by flash floods in the last years, especially during the late summer-autumn period. The latest event occurred on September 10, 2017 and hit the urban area of Livorno, a port city located along the Tuscan coastline. Such event was triggered by a quasi-stationary mesoscale convective system (MCS) characterized by a V-shaped plume, which released about 230 mm of rain in 3h (Regione Toscana, 2017). During this flood, eight people died. Previously, other deadly events affected urban areas widespread along the Italian coasts, such as the major flash floods occurred in 2000 in Calabria (13 deaths—Federico et al., 2003; Montani et al., 2003), 2009 in Siciliy (38 deaths—Aronica et al., 2012), 2011 in Tuscany and Liguria (19 deaths—Cevasco et al., 2012; Silvestro et al., 2012; Faccini et al., 2015). These and other events, also affecting inland areas of Italy, may be correlated with a general tendency toward a decrease in the number of rainy days, and an increase in the frequency of high-intensity precipitation events. According to Brunetti et al. (2004), these trends were observed since the end of the nineteenth century and resulted amplified in the last decades.

Ducrocq et al. (2014) mostly relate the occurrence of heavy precipitation events in the Mediterranean region to the MCSs development. Most of these systems occur in the autumn season (September to November), when the Mediterranean Sea is still relatively warm with high evaporation rates (Mariotti et al., 2002). The warm sea surface temperatures increase air-sea surface heat fluxes, which in turn moisten and destabilize the marine atmospheric boundary layer (Ducrocq et al., 2014). The warm and moist low-level air flows that impact with reliefs surrounding the Mediterranean Sea are then uplifted, initiating a cycle of condensation and convection processes that can be renewed at the same location for hours, as long as the flux of low-level moisture remains intact (Ducrocq et al., 2014). In addition, flash floods can be also triggered by short-lived convective cells with a limited areal extent (typically <100 km²), or by occasional heavy rainfall developing within large-scale perturbations lasting several days (Gaume et al., 2016).

Location and magnitude of flash floods are influenced by rainfall properties (i.e., spatial extent, duration, intensity), as well as by topography, land use, soil types, and the initial soil moisture content of the affected watersheds (Hapuarachchi et al., 2011). In many regions, natural watersheds underwent massive urbanization associated to strong modification of local drainage patterns. The permeable soils were replaced with impermeable surfaces such as roads, roofs, parking lots, and sidewalks, reducing infiltration of water into the ground and accelerating the runoff rates (Konrad, 2003). These modifications increased both likelihood of flash floods and risk conditions. In spite of this, people living in flood or landslide prone areas are expected to increase in the coming years because of population growth and urban development (United Nations, 2011). On the other side, the globally averaged surface temperature is also projected to rise over the twenty-first century under all assessed emission scenarios (IPCC, 2014). For these reasons, it is very likely that the climate change-related effects, such as extreme floods, flash floods, and rapid-moving landslides (e.g., soil slips, debris flows, rock falls) will occur more often, potentially leading to loss of human life and economic damages.

In this work, we have analyzed the historical frequency of flash flood events affecting the urban area of Pozzuoli, a densely populated town located near Naples in southern Italy (**Figure 1**). In a previous paper (Esposito et al., 2015), we described the flash flood event that hit the Pozzuoli town on November 6, 2011, by analyzing meteorological conditions, rainfall properties, ground effects and results obtained by simulating the flow propagation. The new research presented in this paper is aimed at verifying the eventual increase of flash flood events in the 1970–2014 time interval and related causes, as well as at characterizing the flash flood-triggering rainfall properties. The work wants also give a valuable contribution to the more wide research topic concerning the identification of an increasing trend in the frequency of extreme floods in the central Mediterranean region under the current climate change scenario.

THE POZZUOLI STUDY AREA

Pozzuoli is the most important town of the Campi Flegrei volcanic area, located in the western sector of the wide urban area of Naples (Italy). Very important roman monuments such as amphitheater Flavius and Temple of Serapis are present in the historical town center of Pozzuoli. More than 81,000 people inhabit this territory, with a population density of about 1,800 inhabitants per square kilometer. The town center is located along the coastline, and spreads out over the surface of four volcanic edifices formed during the past explosive activity of the still active Campi Flegrei caldera (e.g., Di Vito et al., 1999): Astroni, Cigliano, Gauro, and Solfatara (Figure 1). These cones extend from the sea level to an altitude of 325 m a.s.l. Since historical times, slope instability processes including shallow landslides and rock falls involved the volcanic slopes, as well as sediment-laden flows (hyperconcentrated flows and streamflows - Pierson and Costa, 1987), affected the drainage network near the coastline. Nevertheless, only a few studies related to geo-hydrological events in the whole Campi Flegrei area have been made in relatively recent years (e.g., Beneduce et al., 1988; Calcaterra et al., 2003; Di Martire et al., 2012), given that the scientific community dedicated more interest toward volcanic and seismic processes.

The flash flood hazard in the urban area of Pozzuoli have been analyzed in a previous work (Esposito et al., 2015). Authors evaluated in 360 ha the extent of the partially-urbanized watershed involved by flash floods, and in 18.6° its average slope angle, reaching values of about 70° within the drainage channels located in its central zone. The geological framework is characterized by welded pyroclastic rocks like tuff and



FIGURE 1 | Aerial view and location of the Pozzuoli town within the Campania Region, southern Italy. The most important infrastructures and position of the rain gauge are also indicated. The contour interval is 25 m.

ignimbrite, and locally by lava domes/flows. Such lithologies are often mantled by incoherent primary volcaniclastic deposits constituted by ashes, scoria, and pumices, and by reworked pyroclastics and colluvial soils, together with alluvial deposits derived from the water-driven surface erosion of the slopes. The climatic setting is of Mediterranean type, with hot, dry summers and moderately cool rainy winters. Mean annual temperatures are in the order of 16°C and the mean annual rainfall is of ca. 875 mm. The most intense rainy period is between October and December, when heavy rainfall occur (**Figure 2**).



The rain regime in this part of Italy is characterized by typical rainfall patterns, among frontal rainfalls, hurricane-like rainfalls or isolated convective storms (De Luca et al., 2010). The seasonal effects of rainfall on shallow deposits and unsaturated pyroclastic soils have been studied in Cascini et al. (2014). A typical trend of soil suction is characterized by very high values (>30 kPa) from June to August due to the intense sun radiation, high values from September to October (20-30 kPa), intermediate value (10-20 kPa) in November and December, and the lowest values of soil suction (<10 kPa) from January to May (Cascini et al., 2014). No direct measurements of soil moisture were available for the pyroclastic soils outcropping in the study area. Some relative soil moisture data for a 5 cm layer in pyroclastic soils, measured at Bagnoli and Melizzano sites (see BAG and MEL in the inset of Figure 1), show high values during winter to spring rainy periods and very low values during summer dry periods (Brocca et al., 2011; Dorigo et al., 2011). Relationships between soil suction data and soil volumetric water content have been studied in several sites characterized by unsaturated pyroclastic soils in Campania region, such as the slopes near Cervinara (Greco et al., 2013), Nocera Inferiore (De Chiara et al., 2016), and Sarno (De Vita et al., 2013; Napolitano et al., 2016) (inset in Figure 1), showing opposite trends. Generally, such soils can be affected by erosion phenomena at the end of the dry season (August to October), during intense rainstorms, with a consequent triggering of flash flood events dominated by hyperconcentrated flows (Cascini et al., 2014; Vennari et al., 2016).

MATERIALS AND METHODS

In order to characterize the flash flood-triggering rainfalls and precipitation regime, in this study we have analyzed the rainfall data collected by a rain gauge located within the Pozzuoli watershed (**Figure 1**). Given its location, this was able to capture representative flash flood-triggering rainfall data reported in the Results section. Since 1970 and up to 1999, this rain

gauge was included in the national network of the Italian Hydrographical Service, regularly reporting only daily rainfall values. In the year 2000, the rain gauge was included in the real-time hydrological monitoring network of the Campania Regional Agency of Civil Protection, and was set for collecting rain data every 10 min. For this reason, the rainfall dataset used in this study consists of two subsets. The first subset includes daily rainfalls related to the 1970-1999 time interval, that were extracted from Hydrological Annals-Part I, available on the Institute for environmental Protection and Research (ISPRA) website (http://www.acq.isprambiente.it/annalipdf/). The second dataset includes rainfall data collected in the 2000-2014 time interval with a temporal resolution of 10 min, which were supplied by the Campania Regional Agency of Civil Protection (http://centrofunzionale.regione.campania.it/). Analysis of the frequency distribution of daily rainfall values with respect to different thresholds allowed characterizing the rainfall regime of the study area, as well as addressing identification of the occurred flash flood events.

A first list of flash flood events hitting the Pozzuoli urban center, obtained from previous scientific studies and inventories, such as AVI (Aree Vulnerate in Italia) (Guzzetti et al., 1994) and SICI (Sistema Informativo Catastrofi Idrogeologiche) databases (http://sici.irpi.cnr.it/gndci.htm), as well as from local news websites, was presented in the previous work of Esposito et al. (2015). In order to obtain an updated and complete list of flash flood events occurred in the 1970–2014 time interval, it was performed an archive research at the National Library of Naples. Here, information related to flash flood events (rainfall qualitative data, description of damages, photos, etc.) were collected from the most relevant local newspaper (namely "Il Mattino"), edited in Naples since 1892.

In order to address the archive research by avoiding time wasting efforts, we selected an arbitrary daily rainfall threshold of 40 mm that allowed us to identify rainy days with high probability of flash flood occurrence. This kind of threshold corresponded to the lowest value of daily rainfall measured in a day for a flash

flood event occurred in 2009, as reported in Esposito et al. (2015). Our assumption was hence that all events of the first subset should be occurred in days characterized by cumulated rainfall amounts equal or higher than 40 mm. At the same time, we verified if such amount was reached or exceeded at the turn of two consecutive days, also to include potential rainstorms occurred between midnights. In this way, we selected a series of dates that were investigated through the historical archive research. With the aim of identifying all the occurred flood events, we addressed the archive research by visually checking also the newspaper editions related to a brief period (2-3 days) after the day characterized by high probability of flash flood occurrence. After this, in order to avoid data gaps and verify effectiveness of the adopted threshold, we expanded the archive research by checking days characterized by cumulated rainfall amounts higher than 30 mm. An accuracy analysis aimed at validating the used threshold and at verifying relationships between a series of other thresholds and occurrence of flash flood events was then performed.

Supplementary data for the more recent events were collected from web sources, including news websites and social networks. It is worth noting that, in order to avoid an information bias toward the last decade when reports from web sources were widespread, identification of the flash flood events was primarily obtained by means of the archive research, and supplementary materials were only used for better characterizing the magnitude of processes, affected zones, and consequent damages from a qualitative point of view. After the identification of flash flood events, the effective daily rainfall and, when available, other properties, such as maximum 1-h rainfall and 10-min peak storm intensity, were collected. Both flash flood and rainfall data were finally elaborated and represented by means of specific plots, aimed at highlighting related behaviors. A comparison performed in GIS environment between the 1:10,000 topographic map of the study watershed, dated 1965 and available on the Città Metropolitana di Napoli website (http://sit.cittametropolitana.na.it/), and orthophotos related to the years 2000 and 2012, available on the National Geospatial archive of the Italian Ministry for the Environment (www.pcn. minambiente.it/mattm/), allowed quantifying the urban growth occurred through the considered time spans.

RESULTS

Daily Rainfall Regime

The rainfall dataset analyzed in this study is related to the period between January 1970 and December 2014, and includes 13,078 daily rainfall measurements. The first subset (1970–1999, 30 years) encompasses 7,891 days of measurements, while the second subset (2000–2014, 15 years) encompasses 5,187 days. Data related to some years (i.e., years 1972–1974, 1991–1996, 2000) were not included in the analysis because daily rainfall measurements were discontinuous.

The cumulated annual rain time series shows a clear variability around an average value of 875 mm/year (**Figure 3**), and stronger oscillations can be observed during the 1970–1990 period.

The daily rain time series (**Figure 4**) does not show a particular statistical trend, but it is characterized by both seasonal and monthly fluctuations, as well as by several daily peaks also higher than 100 mm. Specifically, it can be noted that daily peaks are more frequent but less severe during the 1970–1999 time interval, and are less frequent but more severe during the 2000–2014 (**Figure 4**).

Analysis of the frequency distribution of daily rainfall values is represented with respect to classes defined according to different rain thresholds (**Table 1**), by using a bi-logarithmic diagram (**Figure 5**). The two analyzed datasets do not show important





Class	Daily rain threshold (mm)	Cumulate number (1970–2014)	Cumulated frequency (1970–2014)	Relative frequency (1970–2014)	Cumulate number (1970–1999)	Cumulated frequency (1970–1999)	Cumulate number (2000–2014)	Cumulated frequency (2000–2014)
<i>P</i> > 0	0.001	13078	100.00	71.13	7891	100.00	5187	100.00
P > 0.1	0.2	3775	28.87	6.68	2153	27.28	1622	31.27
P > 1	1	2902	22.19	13.95	1687	21.38	1215	23.42
<i>P</i> > 10	10	1077	8.24	4.95	631	8.00	446	8.60
P > 20	20	430	3.29	1.86	260	3.29	170	3.28
P > 30	30	187	1.43	0.83	112	1.42	75	1.45
P > 40	40	78	0.60	0.23	51	0.65	26	0.50
P > 50	50	48	0.37	0.18	34	0.43	13	0.25
P > 60	60	25	0.19	0.07	18	0.23	7	0.13
P > 70	70	16	0.12	0.05	11	0.14	5	0.10
P > 80	80	9	0.07	0.04	6	0.08	3	0.06
P > 100	100	4	0.03	0.03	1	0.01	3	0.06

differences in frequency distribution. Taking as reference daily thresholds of 40 and 75 mm, the frequency of days with relatively higher cumulated amounts is similar for the two studied periods: 51 days with rain >40 mm (0.65% of total daily measurements) and 9 days with rain >75 mm (0.11%) during the 1970–1999, 26 days with rain >40 mm (0.50%), and 5 days with rain >75 mm (0.10%) during the 2000–2014 (**Table 1**).

However, by considering the full series, a clear change in the threshold/frequency relationship occurs around thresholds between 10 and 20 mm, as represented by the two different power-laws in **Figure 5**. At the high frequencies (f = 10-100), corresponding to daily rainfalls lower than 10 mm, the curve is very steep indicating a very rapid decrease in frequency for an increase of daily rain. At the low to very low frequencies (f = 0.01-10), corresponding to daily rainfalls higher than 20 mm, the curve flattens and shows a slight linear increasing behavior since the 30 mm threshold.

Monthly distribution of the relatively most rainy days (rain >40 mm/day) occurred within the entire 1970–2014 time period shows a strong seasonality (**Figure 6**), concentrating mostly in the autumn season. However, some differences can be observed between the two analyzed periods, such as a strong peak recorded in October only during the 1970–1999 period.

Inventory of Flash Flood Events

The most dramatic flood event of the last century occurred on October 25, 1918 (AVI database; Guzzetti et al., 1994), causing the collapse of a building and 13 fatalities in the urban center of Pozzuoli. For this event, no rain data is available and the archive research gave us only a qualitative description of the event: "an extreme rainstorm with torrential rain and strong winds caused the collapse of a two-floor building located in San Leonardo street, in the town center, at 5 a.m." (source: "Il Mattino," issue dated 26–27 October 1918). Beside this





main event, and probably other minor events occurred in the first half of the last century, the results of the archive research revealed that in the 1970–2014 time interval nine flash flood events hit the Pozzuoli town (**Table 2**). Among these, four events occurred in the 1970–1999 (i.e., 30 years) and five events in the 2000–2014 (i.e., 15 years) time spans. Analysis of the record showed that events mostly occurred in the late summer-autumn period (78%), whereas in the spring (11%), and summer seasons (11%) they were less frequent. No event has been recorded in the winter season. This seasonal distribution is coherent with the average monthly distribution of the most severe rainfalls occurred in the 1970–2014 time period (**Figure 6**).

Analysis of information related to the most recent events (e.g., 2011 event; Esposito et al., 2015) supplied a clear explanation of the flash flood dynamic in the study area. Specifically, enhanced runoff processes initiate during heavy rainstorms along the volcanic hillslopes, eroding pyroclastic soils and supplying sediment-laden flows downstream. In the upper part of the watershed, these flows pour down along hollows and natural channels (**Figure 7**), where soil slips can also affect hilly slopes and channel sides. Downstream, the sediment-laden flows

TABLE 2 | Flash flood events occurred in the Pozzuoli area in the 1970–2014 time interval and related daily rainfalls.

Flash flood event	Daily rainfall (mm)	% of annual precipitation	2 days rainfal (mm)	
15/10/1971	84.2	9	88.8	
26/08/1975	76	9	101.6	
18/11/1985	81	8	148.2	
10/10/1987	115.6	15	122	
15/09/2001	161.6	20	163.2	
04/03/2005	128	11	133	
14/09/2009	40	4	74.8	
30/07/2010	76.4	7	76.4	
06/11/2011	116.6	14	117.2	

The "2-days rainfall" column reports the overall rainfall amounts measured in both the flash flood-related day and in the previous day.

have been reported along the road network connecting the neighborhoods located up in the watershed with the littoral zone (**Figure 7**). In many cases, flows were able to drag motorcycles and cars together with people. The documented rapid flows induced heavy damages to the railway crossing the urban center and to the buildings realized close to the coastline, or in proximity of the flooded roads (Esposito et al., 2015). Deposits of mud reported in the littoral zone reached up to one meter in thickness, whereas other amounts of sediments flowed into the sea. No fatalities have been recorded during the recognized events.

The accuracy analysis revealed that the adopted threshold of 40 mm was suitable to identify flash flood events occurred in the study area by means of the performed archive analysis. This emerged after testing further thresholds by means of several accuracy statistics commonly used for assessing models of weather forecasting, landslide susceptibility and other (Frattini et al., 2010), such as the False Positive rate (FP), True Positive rate (TP), True skill statistic score (Peirce, 1884; Hanssen and Kuipers, 1965), Efficiency (Finley, 1884), and the Odd ratio skill score (Yule, 1900). Such statistics allowed assessing the model performance by comparing observed data (occurrence or absence of a flash flood response after a daily rain exceeding a given threshold) with the model (i.e., selected thresholds) by means of a contingency table (Table 3). Results of the accuracy analysis are shown in Table 4, with reference to thresholds ranging from 30 to 80 mm.

Generally, FP and TP give an insufficient performance statistics, because they ignore false negatives and false positives, respectively (Frattini et al., 2010). In this study, however, the 30 and 40 mm thresholds gave a full sensitivity. The same result was given by the odd ratio skill score that conveniently ranges between -1 and 1 and takes prior probabilities into account (Frattini et al., 2010), resulting in better scores for rare events like the analyzed flash floods. The True skill statistic score, that uses all elements of contingency table and does not depend on event frequency, showed better results for the 40 mm threshold (**Table 4**). The Efficiency measured the percentage of events that were correctly classified by the used model (threshold), but all

the scores resulted very high because were heavily influenced by the most common class (true negative). It was therefore not evenhanded.

Properties of the Flash Flood-Triggering Rainfalls

Properties of rainfalls related to the four flash floods occurred in the first analyzed time interval are reported in **Table 5** (gray rows). Daily amounts range from a minimum of 76 mm recorded on August 26, 1975 to a maximum of 115.6 mm recorded on October 10, 1987. The 1-h rainfall amount is available for the 1971 and 1987 events only, resulting in 30 and 40 mm respectively. The 10-min peak storm intensity was recorded for the 1971 event only and resulted in 60 mm h⁻¹. These four flash floods were recognized within a selection of 51 days characterized by daily rainfall amounts equal or >40 mm (**Figure 8**), 47 of which did not correspond therefore to any event on checked newspapers.

Properties of rainfalls related to the five flash floods occurred in the second analyzed time interval (subset 2000-2014) are also reported in Table 5 (orange rows). Daily amounts range from a minimum of 40 mm recorded on September 14, 2009 to a maximum of 161.6 mm recorded on September 15, 2001. The 1-h amounts range from a minimum of 34.4 mm recorded on September 14, 2009 to a maximum of 80 mm recorded on September 15, 2001. The 10-min peak storm intensities range from a minimum of 54 mm h^{-1} recorded on March 4, 2005 to a maximum of 140.4 mm h^{-1} recorded on September 15, 2001. These five flash floods were recognized within a selection of 26 days characterized by daily rainfall amounts equal or >40 mm (Figure 9), 21 of which did not correspond therefore to any event. With reference to these 26 days, a torrential character of precipitations is generally recorded, as highlighted in Figure 9, where daily rains, 1-h maximum amounts and peak intensities recorded in 10 min are compared. These observations were obtained thanks to the availability of rain gauge measurements characterized by a high temporal resolution (i.e., 10 min), since the year 2000. For this latter time interval, we have observed that, in the study area, the flash flood triggering rainfalls consist of two types. The type 1 includes abundant rainfall persisting on the same area for several hours, with variable intensities that can reach peaks in 10 min up to 100 mm h^{-1} ; the type 2 includes short-duration, high-intensity rainstorms able to release significant amounts of rain in <1 h. Examples of both types are shown in Figure 10.

A general analysis of rainfall data related to flash floods (**Table 5**) revealed that the 1-h rainfall amount and the 10-min intensity exceeded the 30 mm and 54 mm/h, respectively, as well as daily rainfall usually showed values higher than 76 mm, except for the 2009 event (40 mm).

Land Use Changes

Processes described in the previous section were favored by a severe modification of the natural drainage pattern together with an increased urbanization of the watershed that occurred through the second half of the XX century.





Drainage channels, in fact, were deeply modified and many of them were transformed in roads or were buried by means of culverts, compromising their capability of drain flood discharges. The hydrographic network is partly natural in the hilly sectors of the watershed, on the upstream side of the highway, and became completely anthropized downstream, where roads (**Figures 1**, 7, and **11**) have replaced natural channels. For this reason, in all the recent inventoried events, we found that the channelized flows leaved their natural path to overwhelm the urban center with high energy downstream.

TABLE 3 | Contingency table used for flash flood model evaluation (modified after Frattini et al., 2010).

Observed	Predicted					
	class 0 () daily rain under threshold	class 1 (+) daily rain exceeding threshold				
class 0 (–)absence of FF	() true negative (tn)/absence of FF in a day with daily rain under threshold	(+ -) false positive (fp)/absence of FF in a day with daily rain exceeding threshold				
class 1 (+)occurrence of FF	$(-\ +)$ false negative (fn)/occurrence of FF in a day with daily rain under threshold	$(+\ +)$ true positive (tn)/occurrence of FF in a day with daily rain exceeding threshold				

TABLE 4 | Classification of rainy days with different thresholds combined with flash flood events and results of related accuracy statistics (the best scores are in bold letters).

Rain threshold (mm)	True positive	False positive	False negative	True negative	False positive rate (1-specificity)	True positive rate (sensisivity)	True skill statistics (Pierce's skill score)	Efficiency	Odds ratio skill score
	tp	fp	fn	tn	FP = fp/(fp + tn)	TP = tp/(tp + fn)	TP – FP	(tp + tn)/T	(tp*tn) – (fp*fn)/(tp*tn) + (fp*fn)
30	9	174	0	12,895	0.01331	1	0.98669	0.96241	1
40	9	73	0	12,996	0.00559	1	0.99441	0.96994	1
50	8	42	1	13,027	0.00321	0.88889	0.88568	0.97218	0.99919
60	8	18	1	13,051	0.00138	0.88889	0.88751	0.97397	0.99966
70	8	8	1	13,061	0.00061	0.88889	0.88828	0.97717	0.99985
80	6	3	3	13,066	0.00023	0.66667	0.66644	0.97494	0.99977

 TABLE 5 | Rainfall properties related to the flash flood events occurred in the

 Pozzuoli area.

Flash flood event	Daily rainfall (mm)	1-h rainfall (mm)	10-min peak storm intensity (mm h ⁻¹)
15/10/1971	84.2	30	60
26/08/1975	76	-	-
18/11/1985	81	-	-
10/10/1987	115.6	40	-
15/09/2001	161.6	80	140.4
04/03/2005	128	37	54
14/09/2009	40	34.4	57.6
30/07/2010	76.4	40.6	126
06/11/2011	116.6	44.2	67.2

Gray rows refer to the 1970–1999 subset, orange rows to the 2000–2014 subset.

According to Pesaresi and Marta (2014), the main part of an urban sprawl occurred over the last decades of the XX century (1970–2000). Later, a strong decrease in the growth of artificial surfaces took place. The quantitative GIS analysis performed in this study revealed that the urbanized surface of the Pozzuoli watershed (area 3,615,454 m²) raised of 31.2% in the 1965–2000 time interval, and of only 0.7% in the 2000–2012 time span (**Table 6**). **Figure 11** shows the areas involved by such urban developments.

DISCUSSION AND CONCLUSIONS

In the Pozzuoli area, the oldest and most severe (13 fatalities) flash flood event occurred in 1918 (for location see Figure 11), but no rain data is available for that time. After this, 9 events have been recognized between 1970 and 2014; all of them caused several damages to roads, buildings and the railway in the town center with a comparable level of severity. Data presented in this research highlight as the frequency of flash flood events varied in the Pozzuoli area throughout the 1970-2014 time interval. The most relevant finding is that the frequency increased in the 2000-2014 time span (5 events in 15 years; 0.33 ev./yr) respect to the 1970-1999 time span, when 4 events occurred in 30 years (0.13 ev./yr). Cumulated frequency of days with rainfall amounts equal or higher than 40 mm (threshold selected for the archive research) shows a constant trend through the analyzed time interval (1970-2014), without significant variations (Figure 12). Conversely, cumulated frequency of flash flood events highlights an increasing trend since the year 2000, while the urbanization of the area increased only before the year 2000 (Figure 12).

The lower occurrence of flash floods in the first period cannot be related to a lack of historical data. In fact, information related to the occurrence of all the inventoried flash floods were reported by the same authoritative source for both the analyzed time periods. Such newspaper had always been careful to report any local incident that had an impact on society, such as fatalities and injuries, road diversions, rail delays, flooding





and other damages typical of flash floods, as testified by the example of the oldest event noticed to (i.e., the 1918 one). However, the greater amount of data supplied by web sources for events occurred in the last decade, resulted very useful for better understanding the triggering and evolution mechanisms of flows, the affected areas and types of damages. Specifically, as reported by Esposito et al. (2015) for the 2011 event, video footages captured with smartphones by people living in different areas of the Pozzuoli watershed gave us relevant information about flow types, their paths, and the variable depths reached within the urban center. Many studies emphasized the valuable use of web sources, especially social media for flood mapping (Brouwer et al., 2017 and references therein), streamflow estimation (Restrepo-Estrada et al., 2018), and damage assessment (e.g., Diakakis et al., 2016). According to Pennington et al. (2015), the use of social

media enables to acquire detailed information about the impacts, remediation and triggers at the time of the event, many of which could be otherwise unknown.

The strategy of determining a daily rainfall threshold for addressing the archive research and extending inventory of flash floods occurred in the study area demonstrated to be very effective in terms of time optimization and accuracy of results. The adopted threshold of 40 mm resulted in fact very accurate from a statistical point of view, as reported in **Table 4**. Moreover, taking into account overall results of the accuracy analysis, as well as properties of rainfalls triggering the identified flash floods (**Table 5**), higher daily thresholds of 60 or 70 mm would have led to less valuable results. In referring to this consideration, however, it is worth underlining that application of such method requires a preliminary set of events, which is essential to take



adequate decisions about the daily rainfall threshold that should be used. The only analysis of daily rainfall data series is not wholly sufficient.

In order to address causes of the increased flash flood trend, it was assessed the urbanization progress occurred in the studied watershed throughout the 1965-2012 time interval. The analysis of topographic maps and orthophotos of the Pozzuoli area documented the high level of urbanization, representing a complex and critical factor for risk management in the area. Our GIS analysis allowed assessing that urbanized surface of the Pozzuoli watershed raised from 24 to 55% in the 1965-2000 time interval. A strong decrease in the growth of artificial surfaces took place instead during the 2000-2012 time span, when the urbanization has increased only for 0.7% of the watershed area (Figure 11). In the light of this, the frequency of flash floods and the rate of urbanization followed opposite behaviors, since 2000 (Figure 12), indicating that the temporal distribution of events is ostensibly controlled by different factors. The initial soil moisture content may also have a marginal role, given that in this period is relatively low in the whole region and that the flood-triggering rainfalls are often characterized by high rain rates exceeding the soil's infiltration capacity. In the volcanic and peri-volcanic areas of the Campania Region, moreover, this behavior is enhanced within hillslopes that during the summer season are hit by wildfires generating soil water repellency conditions (e.g., Esposito et al., 2017). In small and highly urbanized watersheds like that of Pozzuoli, the hydrological response to high-intensity rainstorms is very rapid and almost all the rain supplies overland flows characterized by discharges that are often not able to be drained by artificial channels and culverts. Unfortunately, the studied watershed is not gauged with instruments for flow discharge measurements, so that we have not accurate data about this. Analysis of the collected multimedia data also highlighted that flow discharges were increased by a relevant concentration of fine sediments, mostly derived from surface erosion of volcanic soils and pyroclastic units outcropping in the hilly sector of the study area.

Taking into account the findings described previously, the observed increase in the flash flood trend can be reasonably ascribed to some variations in the rainfall regime.

The frequency of days with most abundant cumulated rainfall (i.e., \geq 40 mm) is very similar for the two analyzed periods, and a generally constant trend can be observed since 1970 (**Figure 12**). The factor that has probably played the major control on the increase of flash flood events in the study area has been a change in the rainfall intensity. However, intensity data are fully available only for the 2000–2014 dataset, and are therefore not sufficient to definitively confirm this hypothesis. As consequence, the use of future 10-min registrations would help to define appropriate rain intensity trends.

In the Eastern Mediterranean, a significant rise in the flood frequency, partially related to the increased rainfall intensity (Philandras et al., 2010), has been documented by Diakakis (2014) in the metropolitan area of Athens (Greece). According to this Author, such increment mostly occurred in recent decades (i.e., from 1970 to 2010), with a concentration of events in the autumn season, as in the current case history. In Athens too, heaviest damages occurred in specific locations mostly near the city's natural drainage network, parts of which have been



converted to culverts or streets and were derelict for decades (Diakakis et al., 2016).

Even though it was not possible to calculate significant intensity trends in this study, a torrential character of subhourly precipitations recorded since the year 2000 (e.g., **Figures 9, 10**) is evident. The 1-h amounts range from 30 to 80 mm, while the 10-min peak storm intensities range from 54 to 140.4 mm h^{-1} . According to Tranfaglia et al. (2016), for example, the meteorological event that hit the cities of Naples and Pozzuoli on September 15, 2001 was undoubtedly one

of the most intense events ever recorded on such territory. This event, as well as those triggering the other flash floods, were rather restricted on the most urbanized areas of the

TABLE 6 | Quantitative urbanization changes in the Pozzuoli watershed during the 1965–2012 time span.

year	Urbanized area (m ²)	Urbanized area (% watershed)	Increase (m ²)	Increase (%)	
1965	878,506	24.3	()		
2000	2,008,033	55.5	+1,129,527	+31.2	
2012	2,032,568	56.2	+24,535	+0.7	

Campi Flegrei coastline. This circumstance may be explained by considering these areas as Urban Heat Islands (UHI) that release high amount of suspended particles together with sensible heat in the atmosphere, contributing to increase its mechanical turbulence and to create favorable conditions for severe convective rainstorms (Chandler, 1965). At this effect should be associated the contribution of the rugged topography of the volcanic hillslopes that, also in other parts of the Campi Flegrei (e.g., Santo et al., 2012), influenced both the development of convective cells and triggering of flash floods. According to Marchi et al. (2010), orography plays an important role in regulating atmospheric inflow to the storm and in controlling storm motion and evolution, as well as topographic relief is



necessary for promoting flow concentration along drainage ways, which result in high unit discharges and relevant geomorphic effects.

The increased frequency of flash floods here described seems to do not follow the trends recognized at larger scales in the Mediterranean region, probably due to the limitations of the available datasets and some complex overlapping signals (e.g., decadal and inter-decadal variability). In fact, Madsen et al. (2014) indicate that there is some evidence of a general increase in observed extreme precipitation, whereas there are no clear indications of significant increasing trends at regional or national levels of observed extreme streamflow in Europe. Gaume et al. (2016) also detected no significant trends in the frequency and magnitude of extreme floods for the Mediterranean area. According to Llasat et al. (2016), the lack of a common or significant trend is a general result found in the majority of works dealing with flood time series analysis, as pointed out by the IPCC, both in their reports on Extremes (IPCC, 2012) and in the Fifth Report (IPCC, 2014).

In spite of this, some increments were clearly identified at local scales. For example, other than findings of Diakakis (2014) discussed previously, Llasat et al. (2013) calculated positive trends for extraordinary floods (0.13 floods yr⁻¹) occurred between 1981 and 2010 in north-east Spain, south-east France and southwest Italy. Petrow and Merz (2009) detected increasing trends in flood frequency in some parts of Germany, and correlated them with temporal changes in atmospheric circulation patterns. Wilhelm et al. (2012) identified flash flood events that affected an alpine watershed in the NW French Alps, during the 1740-2007 time span, and discovered that most of the extreme flood events took place since the beginning of the twenty-first century, with the strongest occurring in 2005. These authors also related the occurrence of high magnitude torrential flood events in high-altitude watersheds to the climate warming. An increment of flash floods in the 1986-2015 time span was also found by Vennari et al. (2016) for other parts of the Campania region.

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Taking into account this framework, it seems that local conditions in terms of atmospheric circulation, topography and vulnerability are prevailing on factors acting at larger scales. This requires more efforts to the scientific community engaged in the evaluation of climate change-related effects, also to predict future social consequences.

In summary, results presented in this contribution highlight as urbanization of the Pozzuoli watershed has become relatively stable since the year 2000, whereas frequency of flash floods has increased. This supports the hypothesis of local changes in the extreme precipitations. Information provided by archive chronicles and web sources demonstrate as the current urban setting plays a fundamental role in making the Pozzuoli area prone to future flash flood events that, probably, will be increasingly favored by the occurrence of more extreme rainstorms, as already forecasted for rainfall-induced landslides in other parts of southern Italy (Gariano et al., 2017).

AUTHOR CONTRIBUTIONS

GE performed the archive research. GS and FM performed the statistical analysis; GE, GS and FM processed rainfall and flash flood data; GE and FM wrote the paper.

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