



# Methodological Development of a Combined Preparation for Micropaleontological and Sedimentological Studies of Samples From the Proterozoic Record

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The recovery of microfossils from Proterozoic rocks is commonly challenging because of metamorphism. In this study, an application of different methods usually applied on Phanerozoic rocks to test efficiency on recovering microfossil from Proterozoic units is presented. Chemical, physical, and biological factors can influence the recovery of microfossils, thereby becoming a barrier for biostratigraphic and paleoecological studies. Furthermore, low-cost projects with a reduced amount of sample collected, such as drill core sampling, need to optimize the preparation time and sample needed for different analyses. To overcome this challenge, the classical procedure of mineralized microfossil preparation, the palynological technique, and the study of clay mineralogy with the analyses of diagenetic alteration and the search for possible microfossils in thin sections were combined. Three Proterozoic lithostratigraphic units were selected to develop an integrated procedure for preparing samples for micropaleontologic and sedimentologic studies: the Paranoá Group, Mesoproterozoic, and the Bambuí Group, Ediacarian-Cambrian, Brazil, and Nama Group, Ediacaran-Cambrian, Namibia. Recovering individual microfossils from the Paranoá and Bambuí groups has been a challenge for paleontologists. Therefore, most micropaleontological studies have been done as a part of microfacies analyses in thin sections. All sediment fractions were studied in trial for the examination (and picking) of mineralized microfossils, even the finest ones. The microfossil picking was conducted using a stereomicroscope. Three species were recovered following this procedure: *Vetronostocale* aff. *V. amoenum* Schopf and Blacic, 1971, *Myxococoides* sp., and *Melanocyrrillium* sp. Analyses in whole rock samples of residues from water (H<sub>2</sub>O) and hydrogen peroxide 30% (H<sub>2</sub>O<sub>2</sub>) procedures showed similar results when the clay fraction studied was obtained as part of micropaleontological preparation compared with the results from the standard clay mineral preparation method. The clay fraction diffractograms showed that the micropaleontological preparation with H<sub>2</sub>O and H<sub>2</sub>O<sub>2</sub> caused an increase in the intensity of the quartz reflections compared with

untreated samples. Moreover, detailed protocols for organic-walled microfossil preparation and low concentrated acetic and formic acids attacks for mineralized microfossil extraction were presented.

**Keywords:** micropaleontological preparation, sedimentological preparation, proterozoic microfossils, clay minerals, curatorship protocol

## INTRODUCTION

The diversity and preservation of fossil specimens from the Precambrian have been considered rare compared with those recovered from the Phanerozoic (Knoll, 1985; Schopf, 1995). Among other causes, such as taphonomic alterations, which greatly influence the fossil record, the preparation methodology also plays an essential role in recovering. Therefore, this barrier in the study of the Precambrian strata requires methodological considerations because, depending on the method applied, the fossil record may be lost. The present study proposes a protocol to increase microfossil recovery based on a combined methodology focused on micropaleontological and sedimentological analysis integration (Alves, 1987; Campos, 2012; Horne and Siveter, 2016; Leite et al., 2018). Samples from Paranoá and Bambuí groups, Brazil, and Nama Group, Namibia, were analyzed to assess all methods presented in this study.

Because of distinctive micropaleontological recoveries procedures on samples from Phanerozoic to other strata, it is necessary to formalize preparation methodologies for recovery of organic-walled and mineralized microfossils from Precambrian lithostratigraphic units. With mineralized micropaleontological analyses, the residues from the same preparation can be used for clay mineral analyses. This combination accelerates the whole research besides reducing the costs of preparation procedures. The application of this protocol could improve the recovery of microfossils from Precambrian units and, consequently, improve biostratigraphic studies besides combining analyses for micropaleontology and sedimentology for integration and reduction of costs. In the present case, at least three laboratories are working together, Laboratory of Mineralized Microfossils, Laboratory of Organic-walled Microfossils, and Laboratory of X-ray Diffraction, so curatorial procedures must be shared and followed to promote efficiency on data acquisition and analysis integration.

Moreover, it also detailed the curatorship procedures, identification, allocation by collection category, packaging, and housing samples under the policy of the Museum of Geosciences, University of Brasilia. In addition to the management methodology, rules for the transit of samples between laboratories are also described.

## GEOLOGICAL SETTINGS

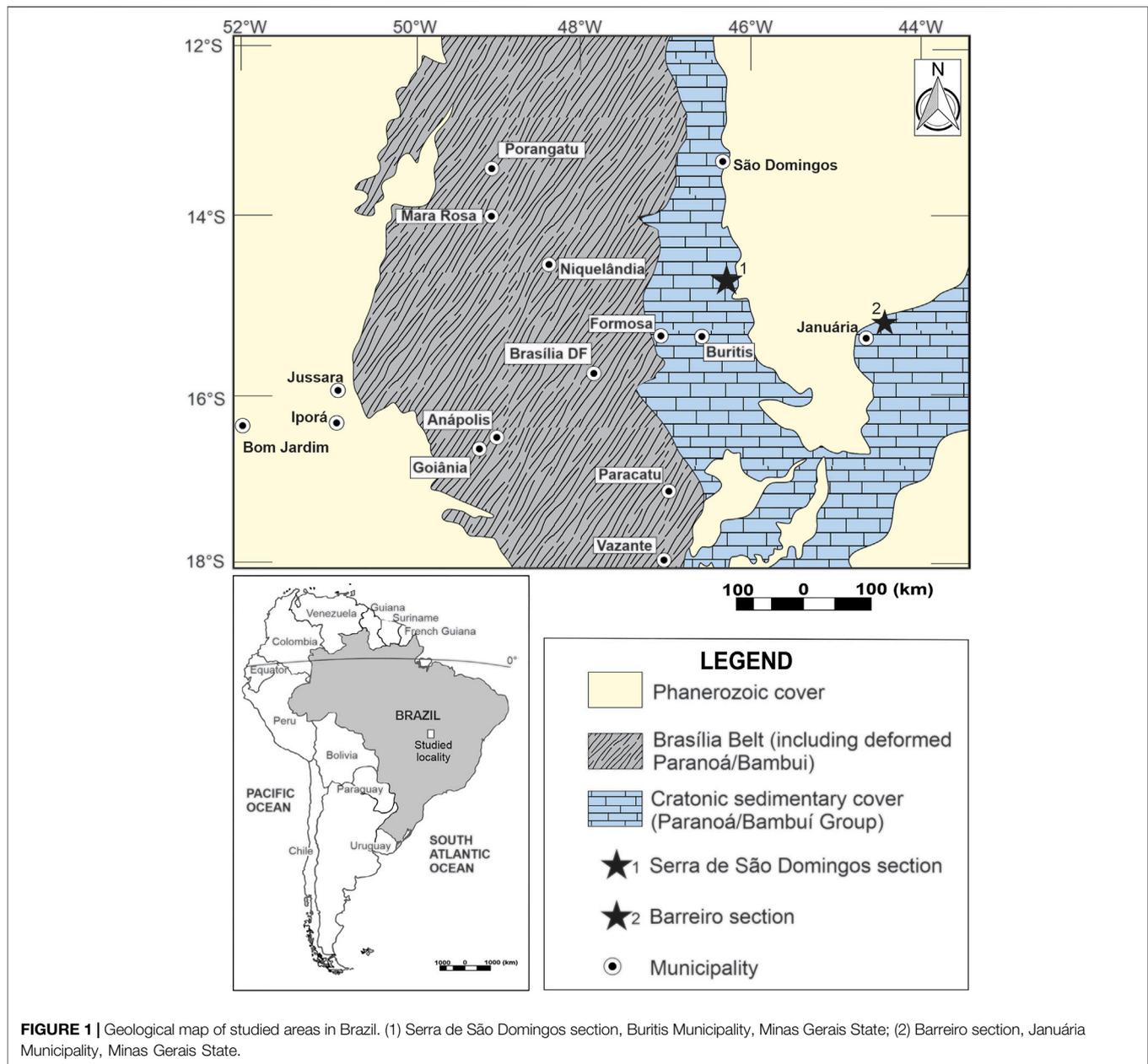
Two localities in Minas Gerais State, Brazil, were studied: the Buritis Municipality, which is part of the Brasília belt within the Tocantins province, and the Januária Municipality, which is located in a nondeformed domain of the São Francisco craton (Figure 1). A thick interval of Meso-Neoproterozoic sedimentary rocks was

deposited along the west portion of the San Francisco craton. These rocks were separated into three stratigraphic units, from bottom to top: Paranoá Group, Jequitai Formation, and Bambuí Group.

The deposition of terrigenous and chemical sedimentary rocks belonging to the Paranoá Group dates from the Mesoproterozoic when the separation of the Rodinia supercontinent generated a passive rift-margin basin, West of the São Francisco craton (Alvarenga et al., 2014). Faria (1995) studied the stratigraphy of the Paranoá Group in the type locality of Alto Paraíso de Goiás and São João D'Aliança municipalities, Goiás State, Brazil; however, the study did not formalize the units according to any stratigraphic code. Thereafter, Campos et al. (2013) formalized 11 stratigraphic units within the Paranoá Group according to the Brazilian Code of Stratigraphic Nomenclature in order to adjust the informal units proposed by Faria (1995). The Paranoá Group consists of, in ascending stratigraphic order, the Ribeirão São Miguel, Córrego Cordovil, Serra da Boa Vista, Serra Almécegas, Serra do Paranã, Ribeirão Piçarrão, Ribeirão do Torto, Serra da Meia Noite, Ribeirão Contagem, Córrego do Sansão, and Córrego do Barreiro formations (Campos et al., 2013) (Figure 3).

After the deposition of the Paranoá Group, because of climatic changes, the Jequitai Formation was deposited under glacial conditions, and their records remain in erosional contact with the Paranoá Group (unconformity) (Uhlein et al., 1995; Caxito et al., 2012). Right above in conformable contact, the carbonated-terrigenous Bambuí Group was deposited in a foreland-type basin generated from the flexure caused by tectonics in the Brasília belt. The Bambuí Group consists of five lithostratigraphic units, from base to top, the Sete Lagoas, Serra de Santa Helena, Lagoa do Jacaré, Serra da Saudade, and Três Marias formations (Dardenne, 1978) (Figure 3). Lately, the Bambuí Group has been attributed to the Ediacaran/Cambrian interval (Pimentel et al., 2011; Warren et al., 2014; Paula-Santos et al., 2015; Moreira et al., 2020; Sanchez et al., 2021).

The Nama Group, Namibia (Figure 2), represents the deposition in a shallow water foreland system; the deposition of the basal portion started around 550 Ma, followed by the deposition of siliciclastic Molasse sediments from the upper portion deposited in 540 Ma (Germs, 1983; Germs and Gresse, 1991). In the central and southern part of Namibia, the Nama Group rests discordantly on the crystalline basement. Its basal portion is represented by a succession of siliciclastic and carbonate rocks with occurrences of skeletal fossils of *Cloudina lucianoi* and other fossils with carbonate skeletons, as well as ichnofossils and palynomorphs in the Kuibis Formation (Germs, 1995; Gaucher et al., 2005). The upper portion of the Nama Group is represented by the Schwarzrand subgroup, which contains the ichnofossil *Phycodes pedum*, *Cloudina*, and palynomorphs (Figure 3) (Germs, 1983; Germs and Gresse, 1991; Gaucher et al., 2005).

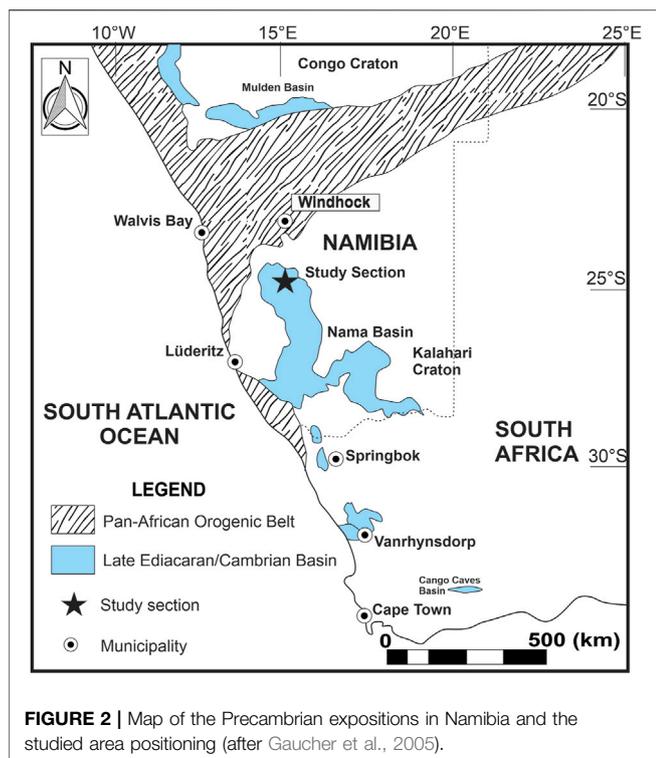


## MATERIALS AND METHODS

The studied material is from three Precambrian units: Paranoá (Mesoproterozoic) and Bambuí (Ediacaran-Cambrian) groups, São Francisco craton, Brazil, and Nama Group (Ediacaran-Cambrian), Namibia. The samples from Brazil were collected in outcrops from Buritis and Januária municipalities, Minas Gerais State (Table 1). Detailed methodological processes for microfossiliferous recovery are discussed in *Preparation Methodologies*.

The same sample was analyzed through different ways to obtain clay minerals information: (1) using the residues from water ( $H_2O$ ) and hydrogen peroxide 30% ( $H_2O_2$ ) micropaleontological preparations; (2) using the standard clay preparation, which initially included material disaggregation with a hammer and

powdering in the Planetary Mill pulverisette by Fritsch for 5 min with 400 revolutions/min. X-ray powder diffraction was carried out on clay fractions. Clay fractions ( $<2\ \mu m$ ) were separated by centrifugation routine at LARIX described by Campos (2012) and modified from Alves (1987). The measurements were undertaken in oriented clay fractions in air-dried conditions. Analyses were performed in a RIGAKU Ultima IV diffractometer equipped with  $CuK\alpha$  radiation, Ni filter, under 35 kV and 15 mA. The samples were scanned at  $5^\circ/min$  velocity, 0.05 stepping ranging from 2 to  $40^\circ 2\theta$  for clay fraction. Mineral phases were identified using Jade XRD 9.0 (Materials Data) with PC-PDF (Powder Diffraction File—PDF for PC—ICDD). Major (M), minor (m), and trace (tr) components were established by comparing the reflection intensities in d: 4.26 Å for quartz, 10 Å for illite, and 7 Å for chlorite.



## CURATORSHIP PROTOCOL

Curatorship procedures must rule the studied material (rocks and fossils samples) management when multiple analysis is performed in different laboratories. This procedure aims to share information about samples, data acquisition, and analysis integration. In this study, the protocol used at the Micropaleontology Laboratory of the University of Brasilia (LabMicro), on curatorship of geological and paleontological samples that become housed at the Museum of Geosciences, was presented. The LabMicro is currently responsible for the Paleontological Collection of the Museum of Geosciences of the University of Brasilia (MGeo), which is subdivided into seven collections: (1) field collection, (2) residual samples, (3) recovered collection, (4) research collection, (5) special collection, (6) didactical collection, and (7) macrofossil collection (Table 2).

Sample curatorship starts during fieldwork. Field sampling is always accompanied by labeling to identify collected samples once they arrive at the laboratory. This is guaranteed by the mandatory completion of an individual sample tag containing data about their recollection site (Figure 4). The field collection comprises materials that have recently arrived at the LabMicro through fieldwork, independent of its immediate use (or not) as research, teaching, and/or training material. If they generate such interest, samples are due to be processed through laboratory work, which will result in both a residual sample and possible recovered fossil assemblage. The residual sample left from preparation is stored in the residual collection in field bags inside storage cabinets, whereas the recovered fossil assemblage is encased in micropaleontological slides to be held

in specific fossil cabinets, consisting of the recovered collection. Research macrofossil and microfossil specimens, used to illustrate taxa in publications such as articles, theses, and reports, are isolated from others either in macrofossil cabinets or microslides that will be deposited at their specific fossil cabinets. In this case, the specimen is relocated into the research collection and recoded with a CP prefix.

Special collection covers fossil material of scientific interest donated or temporarily transferred to the MGeo by partner institutions such as universities, private companies, and other museums. The didactical collection is used in undergraduate and graduate courses given by the Institute of Geosciences, University of Brasilia (IG); it comprises fossil material from other collections at the LabMicro and those collected by professors and students at the IG, as well as third-party direct donations. Finally, the macrofossil collection comprises macrofossil samples that require special conditions for safekeeping because of their size; therefore, they are stored in a cabinet of their own.

Samples arriving at the LabMicro initially get separated into three collections: field, macrofossils, or residual collections (the latter to be prepared for possible microfossil recovery). Once the fossil content is recovered from analyzed samples by picking, it is deposited either on multicelled micropaleontological slides (carbonate/siliceous fossils separated from rock through sieving) or glass microscope slides (organic-walled microfossils concentrated through organic preparation). The possible use of any microfossils on publications requires their relocation into single-celled micropaleontological slides to be stored in the research collection cabinet or the relocation of the entire glass microscope slide (with microfossils of scientific relevance properly marked) into the same space.

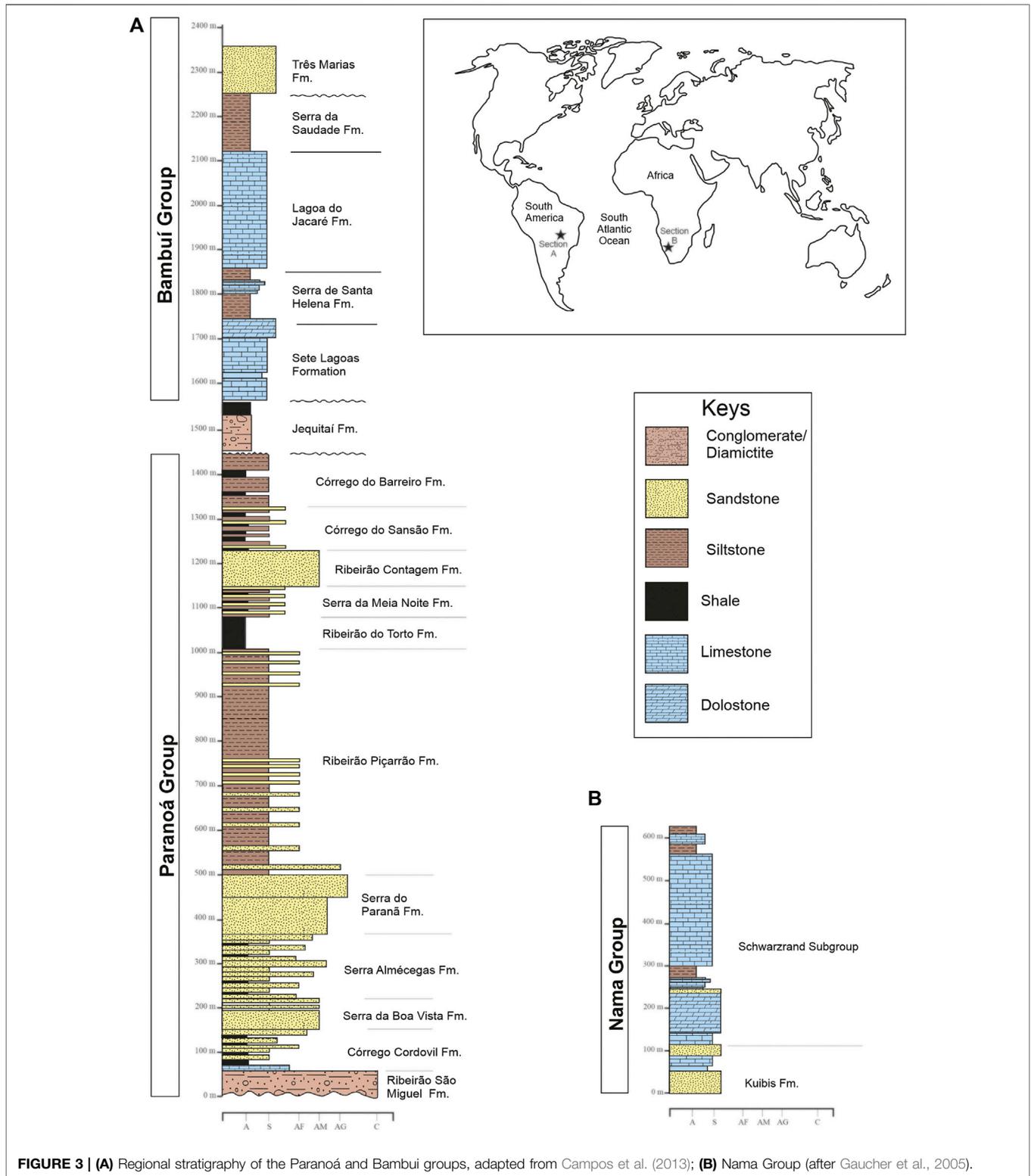
## PREPARATION METHODOLOGIES

Once the initial steps of the curatorial procedure are completed, thin-section slides of the samples are produced for sedimentological/paleontological studies. Subsequently, a mechanical fragmentation of samples can be performed by using several possible methodologies, including soaking them in H<sub>2</sub>O and/or chemical attack with H<sub>2</sub>O<sub>2</sub>, acetic acid 4%–10%, formic acid 10%, hydrochloric acid 36% (HCl), and hydrofluoric acid 40% (HF).

In the present work, both water and oxygen peroxide preparations were performed for mineralized microfossiliferous recovery. After washing both preparations on a sieve set (composed of 630-, 250-, 120-, 80-, and 50- $\mu$ m mesh sieves plus a collecting bucket underneath), each fraction was analyzed on a stereoscopic microscope to pick for mineralized fossil remains.

### Combined Preparation for Mineralized Microfossils and Clay Minerals (H<sub>2</sub>O or H<sub>2</sub>O<sub>2</sub>)

The preparation presented herein aims to recover mineralized fossiliferous remains from disaggregating 30 g of sedimentary



rock samples. This method is commonly used with Quaternary and Cretaceous units (Horne and Siveter, 2016; Leite et al., 2018; Machado et al., 2020). Two distinct sieving procedures were conducted on the same sample for mineralized microfossil

recovery and clay mineral analyses: (1) treatment with water before sieving and (2) attack with hydrogen peroxide before sieving. Both methods aim to disaggregate the rock sample. After sieving both products from the water treatment and

**TABLE 1** | Samples from Ediacaran units analyzed for specific preparations.

Sample	Fossils recovered	Lithotype	Stratigraphic unit	Locality	Applied methods
MP1203	CP965; CP966; CP967; CP968; CP969	Siltstone	Paranoá Group	Serra de São Domingos section, Buritis, Brazil	H <sub>2</sub> O and H <sub>2</sub> O <sub>2</sub> analyses
MP1221	CP970	Siltstone	Bambuú Group, Sete Lagoas Formation	Rio de São Domingos section, Buritis, Brazil	H <sub>2</sub> O, H <sub>2</sub> O <sub>2</sub> , and clay mineral analyses
MP1226	—	Siltstone	Bambuú Group, Serra de Santa Helena Formation	Rio de São Domingos section, Buritis, Brazil	H <sub>2</sub> O, H <sub>2</sub> O <sub>2</sub> , and clay mineral analyses
MP1231	CP971; CP972; CP 973	Siltstone	Bambuú Group, Serra de Santa Helena Formation	Rio de São Domingos section, Buritis, Brazil	H <sub>2</sub> O, H <sub>2</sub> O <sub>2</sub> , and clay mineral analyses
MP2289	CP974	Limestone	Nama Group, Kuibis Subgroup	Namibia	Low concentrated acetic acid
MP2995	CP961	Limestone	Bambuú Group, Sete Lagoas Formation	Barreiro section, Januária, Brazil	H <sub>2</sub> O, H <sub>2</sub> O <sub>2</sub> , HCl, and HF attacks
MP3013	CP914	Limestone	Bambuú Group, Sete Lagoas Formation	Barreiro section, Januária, Brazil	H <sub>2</sub> O, H <sub>2</sub> O <sub>2</sub> , HCl, and HF attacks
MP3034	CP963	Limestone	Bambuú Group, Sete Lagoas Formation	Barreiro section, Januária, Brazil	H <sub>2</sub> O, H <sub>2</sub> O <sub>2</sub> , HCl, and HF attacks
MP3710	CP916	Limestone	Bambuú Group, Sete Lagoas Formation	Barreiro section, Januária, Brazil	H <sub>2</sub> O, H <sub>2</sub> O <sub>2</sub> , HCl, and HF attacks
MP3714	CP917	Limestone	Bambuú Group, Sete Lagoas Formation	Barreiro section, Januária, Brazil	H <sub>2</sub> O, H <sub>2</sub> O <sub>2</sub> , HCl, and HF attacks

**TABLE 2** | Collections into the paleontological collection of the Museum of Geosciences, University of Brasília.

Collection	Code	Material
Field collection	Code gave during fieldwork	Rock sample
Residual collection	MP	Residual rock and organic fractions
Recovered collection	MP (same as the residual collection)	Microfossils recovered but not illustrated in publications
Research collection	CP	Microfossils illustrated in publications
Special collection	Coded according to their previous repository	Microfossils donated and loaned from another institution
Didactical collection	CD	Rock, microfossils, and macrofossils for didactical purposes
Macrofossil collection	MAF	Macrofossils



**SAMPLE FORM - IG/UnB**  
**SAMPLE P \_\_\_\_\_ Am \_\_\_\_\_**  
P = GPS point (arabic); Am = sampling sequence by point

**Special Collection nº UnB-GEO-E \_\_\_\_\_**

City \_\_\_\_\_

**Objective of the fieldwork** \_\_\_\_\_

**Participantes** \_\_\_\_\_

Sampling date \_\_\_\_/\_\_\_\_/\_\_\_\_

Group/Formation/Member/Suite/Complex \_\_\_\_\_

Geographic location \_\_\_\_\_

Lithotype \_\_\_\_\_

Column level \_\_\_\_\_

GPS point \_\_\_\_\_ Zone \_\_\_\_\_ Datum \_\_\_\_\_ Altitude \_\_\_\_\_

Coord UTM \_\_\_\_\_ mL \_\_\_\_\_ mN,

Obs: \_\_\_\_\_ MP: \_\_\_\_\_

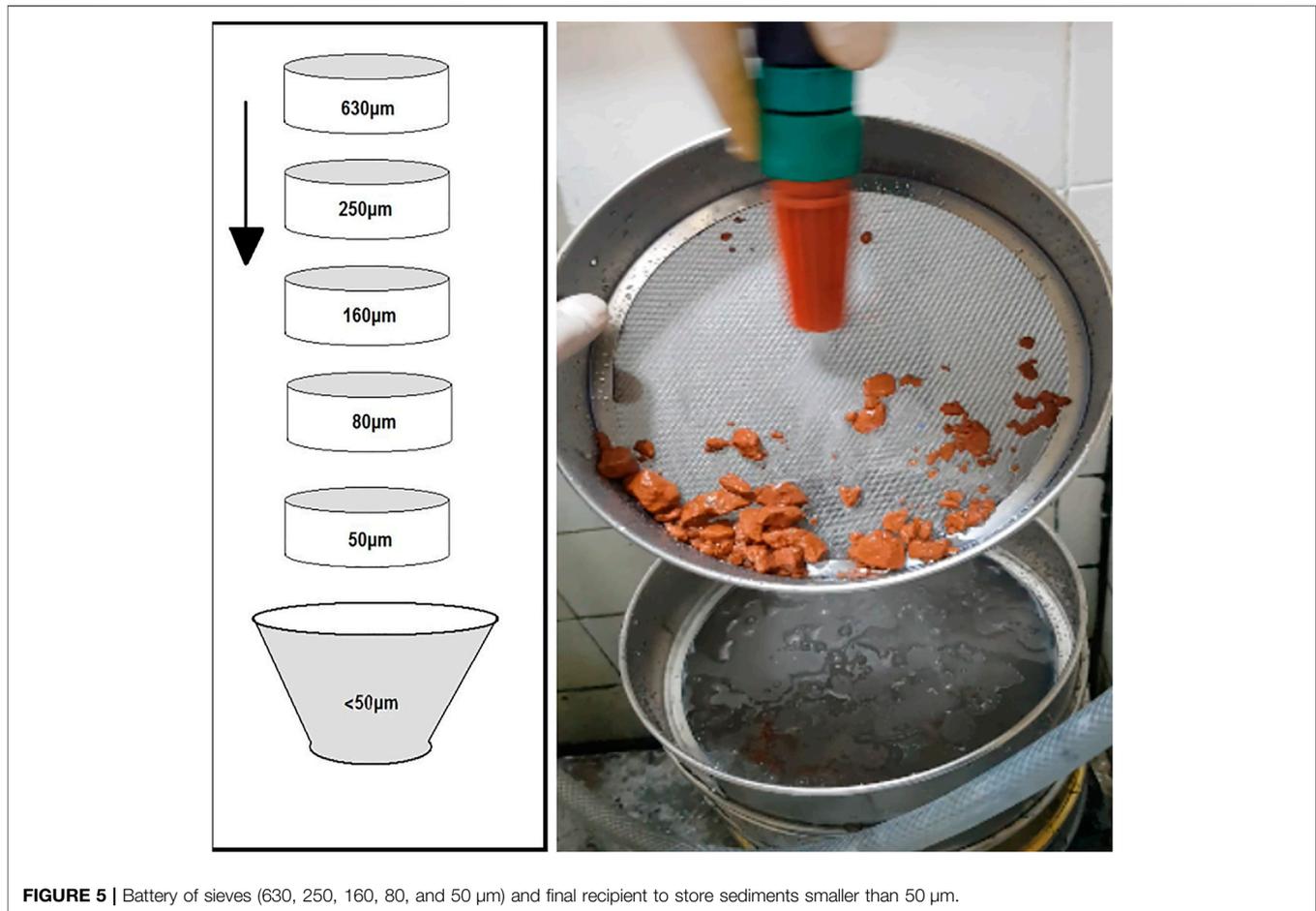
**FIGURE 4** | Sample datasheet used to identify samples during the Laboratory of Micropaleontology fieldwork, University of Brasília, Brazil. The datasheet contains all information needed for further curatorship.

hydrogen peroxide attack, the sedimentary fractions were dried in a laboratory drying oven, and then analyses were performed under a stereoscope microscope to pick microfossils.

After mechanical disaggregation, a single sample followed two preparation routes: (1) left in beaker for 48 h with H<sub>2</sub>O and (2) left in beaker for 48 h with H<sub>2</sub>O<sub>2</sub> 30% (PV). After these procedures, the samples were washed in a battery of sieves (630, 250, 160, 80, and 50 µm) (Figure 5). The fraction smaller than 50 µm were kept in an appropriate container. All fractions were dried in a laboratory drying oven at 60°C and then examined under a stereoscope microscope to pick microfossils. This drying temperature prevents unwanted fragmentation of microfossils. The finest fraction (>50 µm) from both preparations was also analyzed through X-ray diffraction for clay minerals studies.

### Mineralized Microfossiliferous Recovery (Acetic and Formic Acid Attacks)

The traditional study of *Cloudina* species and other tubular carbonate fossils hosted in limestone is performed preferably in two-dimensional (2D) views. This analysis uses polish or thin sections due to the ease of this methodology and quick



**FIGURE 5** | Battery of sieves (630, 250, 160, 80, and 50  $\mu\text{m}$ ) and final recipient to store sediments smaller than 50  $\mu\text{m}$ .

preparation, although studying fossils in 2D views make the 3D morphology reconstruction more complex and less accurate. In some cases, phosphatization processes offer an opportunity to know more about its morphology. The fossil can be easily isolated from the rock matrix by acid attack without destroying the specimen (Hua et al., 2003). In contrast, when the composition of the fossil and that of the matrix are both carbonates, it becomes a challenge to separate the specimen from the rock. This work shows a new methodology of fossil extraction using a low concentrated acetic acid such as vinegar (~4% acetic acid).

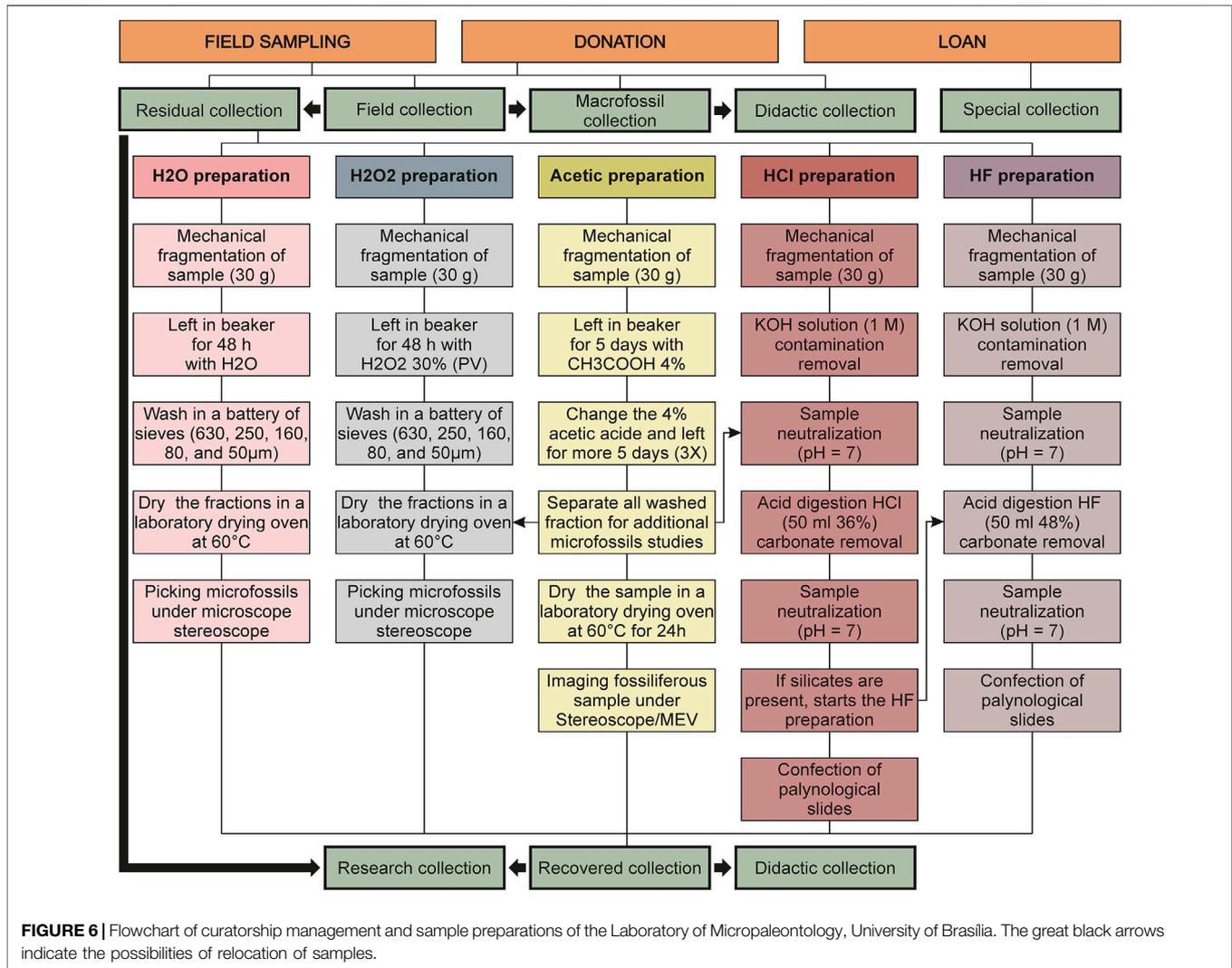
The preparation returned a positive result because acetic acid (4%) attacks the carbonate matrix preferentially, to the detriment of the carapace. Its slightly larger magnesium content is dissolved more slowly than the carbonate matrix. The dissolved fraction of the sample can be separated to analyze the palynological content (Figure 6).

The methodology consists of selecting a fossiliferous sample and introducing acetic acid solution 4% concentration. As the shell composition varies slightly from the matrix's, it allows the acid to act differently, releasing the specimens following the reaction:  $\text{CaCO}_{3(s)} + 2\text{CH}_3\text{COOH}_{(aq)} \rightarrow \text{Ca}(\text{CH}_3\text{-COO})_{2(aq)} + \text{H}_2\text{O}_{(l)} + \text{CO}_{2(g)}$ . A similar, but slower, process occurs in the outcrops of these carbonate fossiliferous rocks, where the

carbonic acid ( $\text{H}_2\text{CO}_3$ ) of the rain erodes the matrix resulting in the eventual exposition of the skeleton. The reaction can be controlled daily by observing the acid's reaction on the fossiliferous sample. The entire preparation cycle takes approximately 15–20 days. The acid must be replaced every 3 days. At the end of the preparation, the sample must be gently and thoroughly washed with running water for approximately 5 min.

After the preparation mentioned previously, the fragments retained in sieves with mesh equal to or greater than 160  $\mu\text{m}$  undergo a new preparation, this time using weak acids, such as acetic acid, to attack limestone, and formic acid, to attack dolomites, both at 10% concentration, with the aim of disaggregating the sample. For each sample to be prepared, it is recommended to use 1 L of 10% diluted acid solution for 200 g of sample. The sample is then placed in a hood, where it will remain until the chemical reaction is complete.

Periodically, after every 24 h of acid attack, it is recommended to change this acidic solution, as it loses its reaction power as the limestone is attacked. The solution that would initially be discarded during the exchange process, as it is a methodological evaluation, is separated for testing in micropaleontology. These tests are carried out with an emphasis on permineralized palynomorphs and for those



microfossils that may be sorted with the aid of a stereoscopic microscope (any particle in suspension).

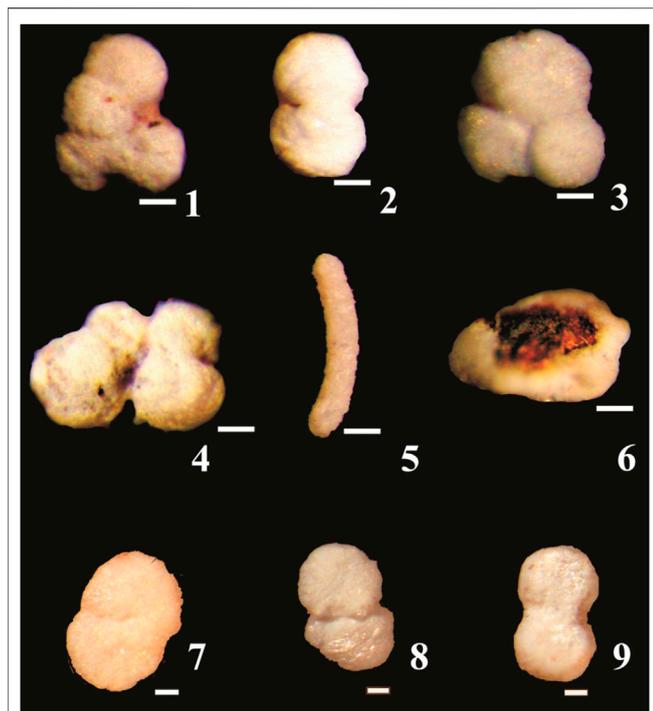
When weak acids are used, the preparation can take up to 2 months to be completed, but instead, the risk of destruction of mineralized microfossils is reduced. After being disaggregated, the material is washed in a battery of sieves. The fraction retained in each sieve is dried in an oven at 60°C and then examined under a stereoscopic microscope.

### Organic-Walled Microfossils Preparation (Hydrochloric and Hydrofluoric Acid Attacks)

Approximately 30 g of sample is used for preparation to recover organic microfossils. Here, the mineral components of the rock are dissolved using two acids: HCl and HF (Figure 6). First, fragmented samples are put in a 400-mL beaker, adding 50 mL of HCl at 36% concentration during 24 h to dissolve carbonates. The next step is to bring the sample solution to a neutral pH value, using distilled water in periodic washings. The neutralization

procedure involves the addition of distilled water to beaker capacity, waiting for the decantation of the organic extract, carefully removing the acid solution; the process is repeated until neutral pH is reached. Then 50 mL of HF at 40% concentration is added to dissolve silicates for 48 h. Then, the washing procedure is repeated. All recovered organic residues are placed in polypropylene tubes and distilled water at pH 7 to further conserve these residues.

After the acid attack process, the final remains are named palynological extract. This material is kept in water solution and, sometimes, when following the classic procedure, needs to be sieved before preparing palynological slides. In synthesis, this traditional procedure uses aleatory organic remains distributed in this solution to prepare palynological slides. Nevertheless, an approach to this classic procedure on picking palynological remains under a stereoscopic microscope is presented. Using a very liner brush (000), it is possible to select specimens to prepare palynological slides with this procedure. There are two ways of making palynological slides: (1) palynological slides created after picking microfossils under a stereoscopic microscope; (2)



**FIGURE 7** | Recovered specimens from the Proterozoic units of the São Francisco craton, São Domingos River section (samples from the Sete Lagoas, MP1221, and Serra de Santa Helena, MP1231, formations) and from São Domingos Hill section (sample from the Paranoá Group, MP1203), Municipality of Buritis, Minas Gerais State, Brazil. **(1–5)** Specimens from the Paranoá Group; **(6)** Specimens from the Sete Lagoas Formation, Bambuí Group; **(7–9)** Specimens from the Serra de Santa Helena, Bambuí Group. **(1–4, 7–9)** *Myxococcoides* sp., respectively, CP965, CP966, CP967, CP968, CP971, CP972, CP 973; **(5)** *Vetronostocale* aff. *V. amoenum* Schopf and Blacic, 1971, CP969; **(6)** *Melanocyrrillium* sp., CP970. Scale bar: 10  $\mu\text{m}$ .

palynological slides without preanalyses under stereoscope microscope, which involves placing a few drops of the recovered organic residue and distilled water on a glass cover. Both types of slides are prepared after putting on a heating plate at 30°C. After the water had evaporated, a few drops of Entellan® resin were added to the coverslip to be completely sealed after contact with the blade. The resin used has the function of drying together with the material mounted on the blade and preventing oxidation of the organic matter and its degradation.

## RESULTS AND REMARKS

The results presented comprehend micropaleontological, mineralized, organic-walled microfossils, and sedimentological data, specifically clay mineral analyses. The oxidation attack was conducted to promote the complete or partial disaggregation of sedimentary rock. When sedimentary rock is composed of organic matrix, the  $\text{H}_2\text{O}_2$  reacts with it and may result in a full or partial disaggregation. In this case, it is possible to recover microfossils in the  $\text{H}_2\text{O}_2$  preparation, even in samples with oxidation considered ineffective (MP1226) and low efficiency (MP1203,

MP1221, and MP1231). After the chemical reaction, the sedimentary material was sieved and by picking finest fractions. The coccoidal structures, as well as tubular and vase-shaped structures, recovered (**Figure 7**) were recognized as fossil content due to their similar morphological and size features assigned to well-described species and genus commonly found in Precambrian units. Besides that, they are very distinct from other grain particles analyzed from the same sample. The species recovered from the Paranoá Group, Mesoproterozoic, from sample MP1203, comprehends *Myxococcoides* sp. (CP965, CP966, CP967, CP968) (**Figures 7.1–4**) and *Vetronostocale* aff. *V. amoenum* Schopf and Blacic, 1971 (CP969) (**Figure 7.5**). One species was recovered in the Sete Lagoas Formation, Bambuí Group, from sample MP1221: *Melanocyrrillium* sp. (CP970) (**Figure 7.6**), and one species was recovered from the Serra de Santa Helena Formation, Bambuí Group, from sample MP1231: *Myxococcoides* sp. (CP971, CP972, CP973) (**Figures 7.7–9**).

The limestone samples of Sete Lagoas Formation, Januária Municipality, did not show a considerable disaggregation effectiveness. The  $\text{H}_2\text{O}_2$  disaggregation method shows more effectiveness on siliciclastic rocks when compared with carbonate rocks. This could be due to the difference in permeability of those two lithotypes. The more permeable the rock, the easier the  $\text{H}_2\text{O}_2$  reacts with the organic matter content. In this context, metamorphism can also affect the  $\text{H}_2\text{O}_2$  disaggregation process as, depending on the metamorphic grade, it could change the rock permeability because of rock compaction.

The finest fraction (>50  $\mu\text{m}$ ) sieved from samples MP1226, MP1221, and MP1231 from three distinct micropaleontological preparations procedures were analyzed: (1) treatment with tap water before sieving, (2) treatment with deionized water before sieving, and (3) hydrogen peroxide attack before sieving. Analyses in whole rock from all three procedures showed similar results when the clay fraction studied was obtained as part of micropaleontological preparation compared with the results from the standard clay mineral preparation method. The total similarities between diffractograms could be verified when both oxidized ( $\text{H}_2\text{O}_2$ ) and nonoxidized (tap water and deionized water) preparations of the same sample (**Table 3**) are compared. The mineral composition of the whole rock sample, determined by X-ray diffraction, shows that all samples have quartz as their major constituent, besides the sample MP1221, which also has calcite as the major component. Illite and albite are minor constituents of all samples.

The standard clay mineral preparation results present changes in reflection intensities compared with the whole rock: the phyllosilicates have higher reflection intensity, which becomes major constituents, whereas the quartz reflection intensity decreases, which becomes a minor constituent. When the standard clay mineral preparation results are analyzed, the clay fraction shows the same composition as the whole rock, but (except for calcite in MP1221) the reflection intensities are opposite to those of the whole rock. Chlorite and illite are major constituents in the clay fraction, whereas quartz and albite are minor constituents (**Figure 8**). The diffractograms of samples MP1226 and MP1221 show a low and ill-defined band at the d~28 position that expands slightly under treatment with

**TABLE 3** | Mineral composition of siltstones in whole rock and clay fraction, indicating the major constituents (M), minor (m), and trace (tr).

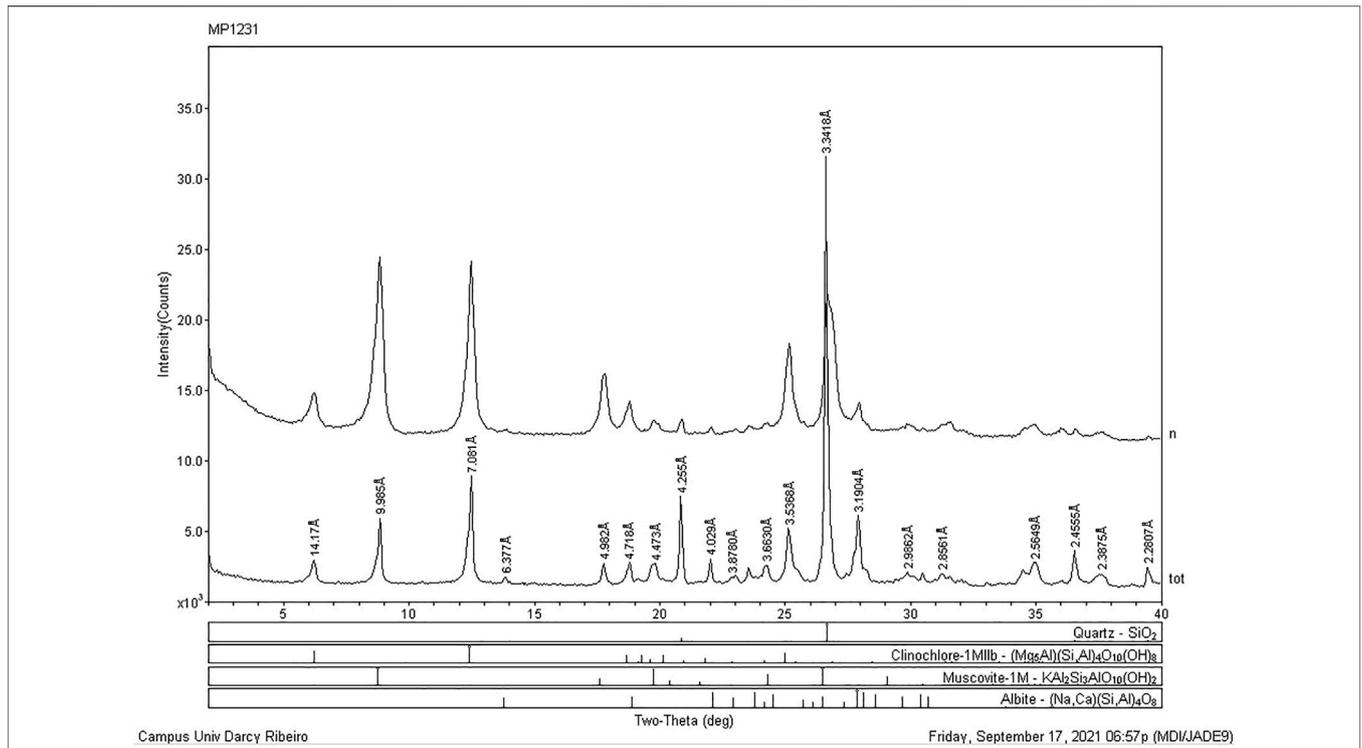
Sample	Preparation	Identified minerals	Whole rock	Clay fraction	
MP1226	Standard clay mineral preparation	Clinochlore (chlorite)	m or tr	M	
		Illite (muscovite)	m or tr	M	
		Quartz	M	M	
		Albite (feldspar)	m	M	
	Micropaleontological residues	H <sub>2</sub> O	Clinochlore (chlorite)	m	m
			Illite (muscovite)	m	m
			Quartz	M	M
			Albite (feldspar)	m	m
		H <sub>2</sub> O <sub>2</sub>	Clinochlore (chlorite)	m	m or tr
			Illite (muscovite)	m	m
			Quartz	M	M
			Albite (feldspar)	m	m
MP 1221	Standard clay mineral preparation	Clinochlore (chlorite)	m	M	
		Illite (muscovite)	m	M	
		Quartz	M	m	
		Albite (feldspar)	M	m	
		Calcite	M	M	
	Micropaleontological residues	H <sub>2</sub> O	Clinochlore (chlorite)	m	M
			Illite (muscovite)	m	M
			Quartz	M	M
			Albite (feldspar)	m	m
			Calcite	M	M
		H <sub>2</sub> O <sub>2</sub>	Clinochlore (chlorite)	m	m or tr
			Illite (muscovite)	m	m
Quartz			M	M	
MP 1231	Standard clay mineral preparation	Clinochlore (chlorite)	M	M	
		Illite (muscovite)	M	M	
		Quartz	M	m or tr	
		Albite (feldspar)	m	m or tr	
	Micropaleontological residues	H <sub>2</sub> O	Clinochlore (chlorite)	M	M
			Illite (muscovite)	m	m
			Quartz	M	M
			Albite (feldspar)	m	M
		H <sub>2</sub> O <sub>2</sub>	Clinochlore (chlorite)	M	M
			Illite (muscovite)	m	m
			Quartz	M	M
			Albite (feldspar)	m	M

glycerol, indicating the presence of interstratified clay mineral, possibly illite/vermiculite. Clay residues obtained from samples treated with H<sub>2</sub>O and H<sub>2</sub>O<sub>2</sub> during the micropaleontological preparation do not maintain this trend. The clay fraction maintains the same intensities as the total sample: quartz remains a major constituent, whereas phyllosilicates are presented as minor or trace constituents (**Figure 9**). This effect can be explained as the effect of disaggregation, dispersion, or release of quartz particles during micropaleontological treatment.

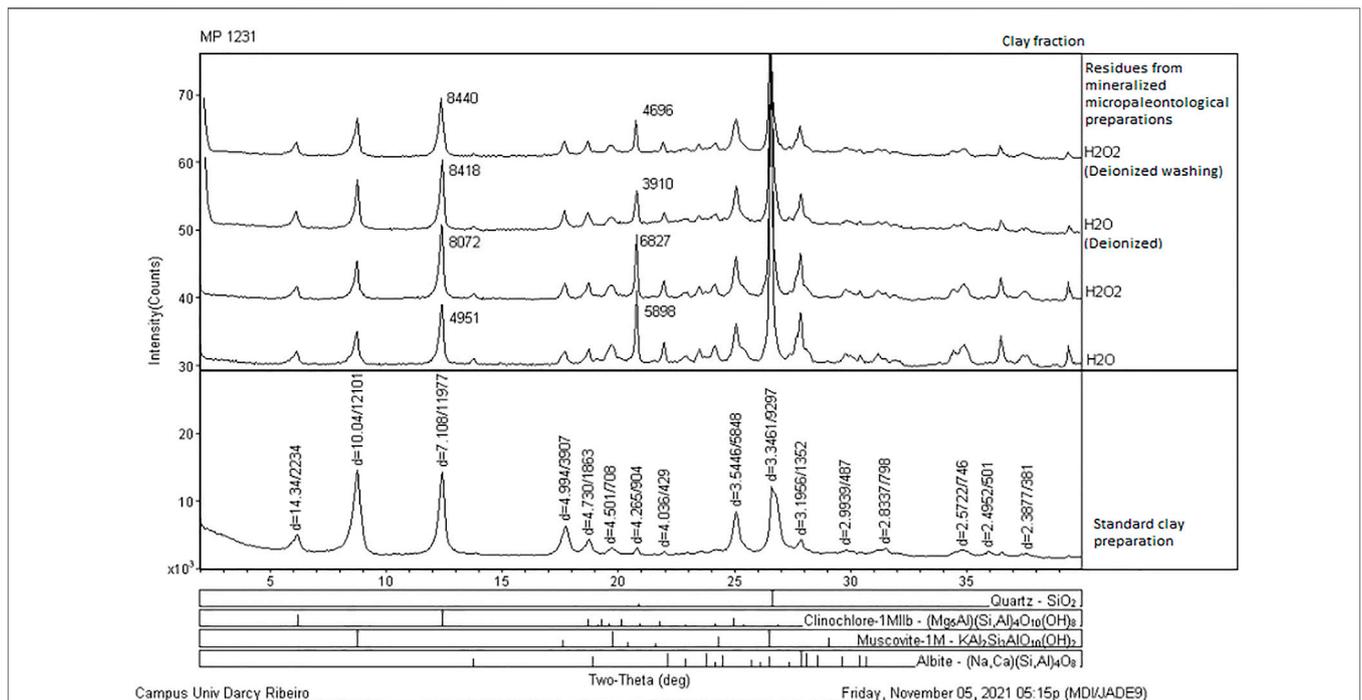
The procedure of analyzing the same sample residue for both micropaleontological and sedimentological approaches as a combined preparation reduces time of maceration and costs, besides being sure that both analyses comprehend the

same depositional interval. This association leads to a more precise paleoenvironmental interpretation. In addition, it can save samples when a small amount is available for multiple analyses.

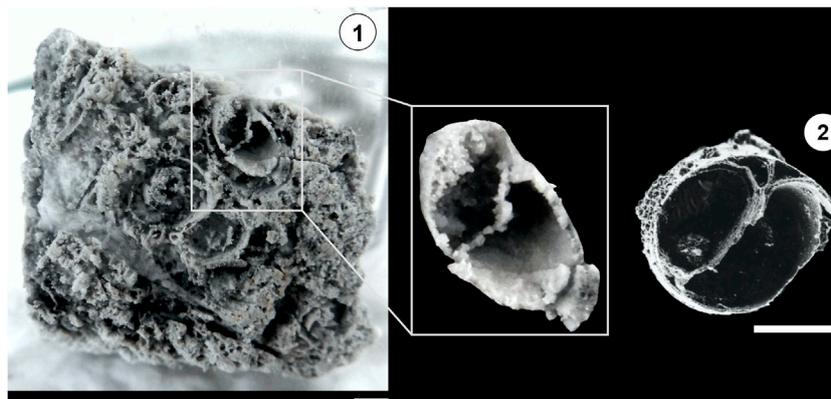
The acetic acid preparation was conducted on a sample from the Nama Group, Namibia; it showed efficient extraction of carbonate *C. lucianoii* skeleton within a carbonate matrix. This extraction technique allowed 3D imaging of the carbonate skeleton (**Figure 10**). This preparation shows similar results compared with phosphatized skeleton preparations from Dengying Formation in China (Hua et al., 2005). Researchers might use this easy, accessible, and environmentally friendly method to conduct 3D studies on carbonate skeleton fossils within limestone rocks. This



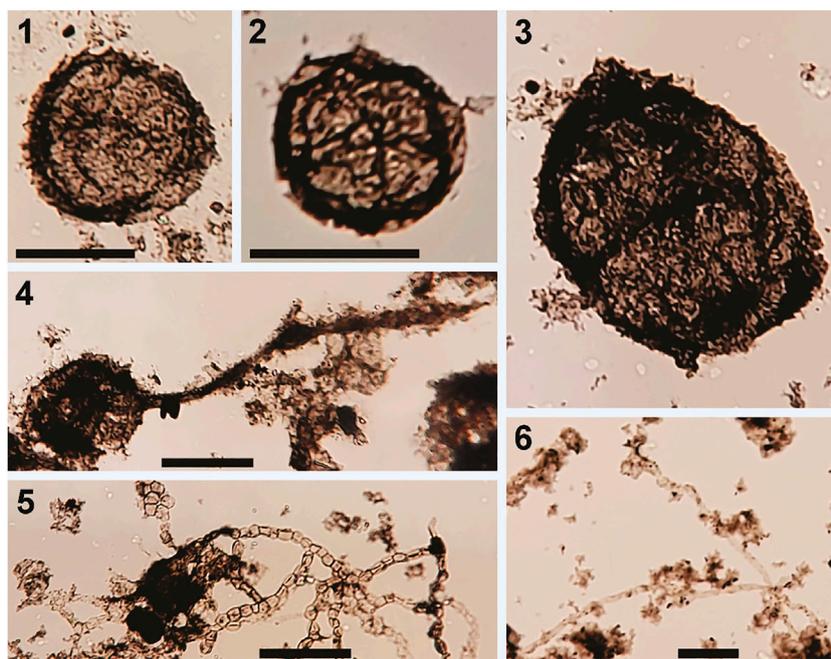
**FIGURE 8** | Diffractograms of sample MP1231 from Bambuí Group, Municipality of Buritis, Minas Gerais State, Brazil. Whole rock (tot) and clay fraction (n); note the variation in reflection intensities.



**FIGURE 9** | Diffractogram of the sample MP1231 from Bambuí Group, Municipality of Buritis, Minas Gerais State, Brazil. Clay fraction (n); note the variation in reflection intensities between the diffractogram of untreated sample (base) and treated samples; note that all treatments have the same effect on the clay fraction.



**FIGURE 10** | Comparison between results of different preparations. **(1)** *Cloudina lucianoii* (Beurlen and Sommer, 1957), CP974. Sample MP2289, carbonate from the type-section of *Cloudina riemkeae* Germs, 1972, Nomtsas Formation, Namibia, (UTM Coord. 33 k 0667883 7358829); **(2)** Phosphatized skeleton of *Cloudina lucianoii* from Dengying Formation, China (Hua et al., 2005). Scale bars: 400 µm.



**FIGURE 11** | Organic-walled microfossils recovered from Sete Lagoas Formation, Bambuí Group, Januária Municipality. **(1–2)** *Leiosphaeridia minutissima* (Naumova, 1949)—CP963; **(3)** *Leiosphaeridia tenuissima* Eisenack, 1958—CP914; **(4)** *Germinosphaera bispinosa* Mikhailova, 1986 - CP917; **(5)** *Ghoshia* sp.—CP916; **(6)** *Siphonophycus robustum* (Schopf, 1968)—CP961. Scale bars: 25 µm.

extraction is possible when the skeleton composition is slightly or entirely different compared with the host carbonate matrix, such as the case of (1) the slightly richer magnesium *C. lucianoii* skeleton from Namibia and (2) the complete different composition of the phosphatized *C. lucianoii* skeleton from Dengying Formation, China (Hua et al., 2005).

The organic-walled microfossil preparation of limestones from Sete Lagoas Formation, Bambuí Group, Januária Municipality, Minas Gerais State, Brazil (Table 1), which

followed the protocol presented in this work, led to the recovery of exquisitely specimens of organic-walled microfossils. The recovered assemblage comprises *Leiosphaeridia minutissima* (Naumova, 1949), CP963 (Figures 11.1, 2) and *Leiosphaeridia tenuissima* Eisenack, 1958, CP914 (Figure 11.3), one acritarch: *Germinosphaera bispinosa* Mikhailova, 1986, CP917 (Figure 11.4), and two cyanobacteria species: *Ghoshia* sp., CP916 (Figure 11.5), and *Siphonophycus robustum* (Schopf, 1968), CP961 (Figure 11.6).

## CONCLUSION

- (1) Efficiency of mineralized microfossiliferous disaggregation using H<sub>2</sub>O<sub>2</sub>: differences in disaggregation efficiency were observed, varying from ineffective (MP1226) to low efficiency (MP1203, MP1221, and MP1231). The lithotype, the amount of organic matter within the matrix, and the metamorphic grade can influence the disaggregation efficacy. The H<sub>2</sub>O<sub>2</sub> disaggregation method shows more effectiveness on siliciclastic rocks when compared with carbonate rocks. This could be due to the difference in permeability of those two lithotypes. The more permeable the rock, the easier the H<sub>2</sub>O<sub>2</sub> reacts with the organic matter content. In this context, metamorphism can also affect the H<sub>2</sub>O<sub>2</sub> disaggregation process as, depending on the metamorphic grade, it could change the rock permeability due to rock compaction.
- (2) Mineralized microfossils recovered using the H<sub>2</sub>O<sub>2</sub> preparation: three permineralized species were recovered: *Vetronostocale* aff *V. amoenum* Schopf and Blacic, 1971 (from Paranoá Group), *Myxococcoides* sp. (from Paranoá Group and Lagoa do Jacaré Formation, Bambuí Group), and *Melanocyrrillium* sp. (from Sete Lagoas Formation, Bambuí Group).
- (3) Organic-walled microfossils recovered from Sete Lagoas Formation, Bambuí Group, using HCl and HF preparation: *L. minutissima* (Naumova, 1949), *L. tenuissima* Eisenack, 1958, *G. bispinosa* Mikhailova, 1986, *Ghoshia* sp., *S. robustum* (Schopf, 1968). The organic residue can integrate organic carbon isotopic studies.
- (4) Mineralized microfossils recovered using acetic acid preparation: *C. lucianoii* (Beurlen and Sommer, 1957).
- (5) Integration of clay mineral and micropaleontology preparations methods: the whole rock diffractograms of siltstones without treatments (standard preparation for clay mineral analyses) or treated in micropaleontological preparation with H<sub>2</sub>O (tap water or deionized water) and H<sub>2</sub>O<sub>2</sub> did not show differences and allow the determination of mineral composition.
- (6) Clay fraction diffractograms of residues from micropaleontological preparation: The clay fraction diffractograms showed that the micropaleontological preparation caused an increase in the intensity of the quartz reflections compared with untreated samples. Samples obtained after micropaleontological treatment may not be suitable for assessing the intensity of diagenesis using the Kübler Index, but they are useful for identifying the mineral assemblage.
- (7) 3D extraction of a skeletal fossil can be possible even when the skeleton is carbonate in a carbonate matrix using weak acetic acid dissolution. This extraction is possible when the skeleton composition is slightly or entirely different from the host carbonate matrix. The organic material released by this preparation can be integrated into palynology studies.

## DATA AVAILABILITY STATEMENT

The original contributions presented in the study are included in the article/Supplementary Material, further inquiries can be directed to the corresponding author.

## AUTHOR CONTRIBUTIONS

MD—Fieldwork, data collection and interpretation, drawing of the figures and writing of the manuscript. RA—Fieldwork, data collection and interpretation, drawing of the figures and writing of the manuscript. DC—Ph.D. supervisor of MD, fieldwork, data interpretation and writing of the manuscript. EG—Data interpretation, drawing of the figures and writing of the manuscript. DW—Fieldwork and review of the paleontological and stratigraphic aspects of the manuscript. CA—Fieldwork and review of the stratigraphic aspects of the manuscript. GG—Fieldwork, data collection. LA—Curator of the Paleontology Collection of the Museum of Geosciences, University of Brasília, review the curatorship protocol and the mineralized microfossiliferous preparation protocol. CV—Data interpretation and review of the manuscript. OJ—Data collection and interpretation of the mineralized microfossils.

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