



Geological Characteristics of the Mesozoic Unconformities in Eastern Heilongjiang, NE China: Implications for the Mesozoic Continental Margin Evolution of Northeast Asia

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Most of the significant petroleum- and coal-bearing sedimentary basins in Northeast Asia originated *via* rifting and thermal subsidence during the Late Jurassic-Early Cretaceous, followed by basin inversion in the Late Cretaceous. However, the tectonic background governing these basin prototype shifts has not been fully explored. The unconformities are excellent archives of plate boundary interactions and geodynamic switches in subduction zones. The Eastern Heilongjiang Province (EHLJ), Northeast China (NE China), comprises a series of Mesozoic-Cenozoic residual basins with well-preserved successions and provides significant insights into the tectonic characteristics and background of Northeast Asia. Mesozoic unconformities and large-scale contractional structures in the basins mark a series of important tectonic transitions in Northeast Asia. Based on the synthesis information of regional Mesozoic unconformities identified in the seismic reflection profiles and field outcrops of EHLJ, the tectonic characteristics and geodynamic background of the Mesozoic continental margin basins in Northeast Asia are analysed. The Middle-Upper Jurassic/basement unconformity (U1) can only be found in some areas of the Sanjiang and Hulin basins. It was a response to the continental collision of Siberia and the northern China–Mongolia tract along the Mongolia–Okhotsk suture during the Jurassic. The Paleo-Pacific Plate rapidly subducted in the NNW direction towards the eastern margin of Eurasia in the early Lower Cretaceous resulting in a mass of strike-slip faults and the widespread absence of deposits (Valanginian) (U2) in the EHLJ. Because of the subduction slab rollback of the Paleo-Pacific Plate during the late Lower Cretaceous, the local asthenospheric material upwelled, and fault and volcanic activities intensified in Northeast Asia. The Lower Cretaceous Dongshan Formation (Fm)/Muleng Fm unconformity (U3-1) reflects a specific scale of bimodal magmatism in the Songliao Basin and the EHLJ. The Pacific Plate subducted in a transformation from NNW to WNW during the early Upper Cretaceous (Cenomanian). The Houshigou Fm (Qixinhe Fm)/ Lower Cretaceous angular unconformity (U3) reflects that on the basins experienced denudation after being extensively uplifted from the subduction events. With the subduction of the Kula Plate, a compression stress field during the later Upper

Cretaceous Period controlled NE China. The basins underwent a widely compressive deformation, accompanied by large-scale thrusts, denudation and deplanation, resulting in Paleogene/Cretaceous unconformity (U4) was formed.

Keywords: unconformities, geodynamic characteristics, Mesozoic, continental margin, Northeast Asian

1 INTRODUCTION

Several significant petroleum- and coal-bearing sedimentary basins exist in Northeast Asian, such as the Songliao, Sanjiang, Hailar, Erlian, and Sanjiang-Middle Amur basins (Cao and Zheng, 2003; Wu et al., 2004; Liu et al., 2006; Jia and Zheng, 2010; Men et al., 2010). Previous studies have proposed that these basins have undergone underwent rifting, thermal subsidence; and then basin inversion since the Mesozoic (Song, 1997; Wang et al., 2016; Guo et al., 2018). It is widely accepted that the two-sided active continental margin tectonics of the Mongolia–Okhotsk Belt, with suturing in the north, and the Pacific Plate subduction in the east during the Mesozoic are caused the series of basin tectonics (Northrup et al., 1995; Maruyama et al., 1997; Song, 1997; Wang et al., 2001; Zhang et al., 2012; Song et al., 2015; Cui et al., 2020).

The NE Asian margin has undergone prolonged Paleo-Pacific subduction and the accretion of island arcs, oceanic plateaus and microcontinents since the early Mesozoic (Meng, 2003; Zhang et al., 2017b). Two decisive factors caused these tectonic movements. One is the closure of the Mongolia–Okhotsk Sea in the north during the Late Jurassic–Early Cretaceous (Halim et al., 1998; Zhang et al., 2001; Ren et al., 2002; Kirillova, 2003; Liu et al., 2004; Metelkin et al., 2007; Yin, 2010; Zhu et al., 2010; Guo et al., 2018); the other is the ancient Pacific Plate subduction beneath the Eurasia Plate in the east during the Mesozoic (Faure et al., 1995; Shao, 1995; Liu et al., 2001; Kirillova, 2003; Li et al., 2003; Huang and Zhao, 2006; Chen et al., 2008; Duan et al., 2009; Cho et al., 2016). Several studies have proposed significant collision between the Okhotomorsk Block and Northeast Asia in the Late Cretaceous (Yang, 2013; Zhang et al., 2017b). This tectonic configuration has created complicated structural patterns within Northeast Asia (Song et al., 1996; Wang, 1997; Cao and Zheng, 2003; Kirillova, 2003; Wen et al., 2008). However, these geological events in Northeast Asia are poorly explained using a successive tectonic and geodynamic model of the eastern Asia continental margin (Yang, 2013; Lee, 2018) because the complex evolution of the NE Asian margin involves deformations of extension, contraction and strike-slip related to active margin tectonics.

The tectonic history, including mantle dynamics, kinematic interactions between converging continental and oceanic plates, and variations in the convergence and slab rollback rates, can be recorded on the surface and at shallow lithospheric depths as tectonic responses (Dilek and Sandvol, 2009; Zhang et al., 2015; Zhang et al., 2017b). As boundaries of stratigraphic units, the unconformities are common phenomena at shallow lithospheric depths (Steno, 1669) and can be divided into angular unconformity, nonconformity and disconformity. Their formation results from sea-level changes and regional tectonic

processes. They are significant in the crust's evolution and deformation mechanisms and provide vital evidence and guidance in reconstructing orogenic phases or dividing geotectonic cycles (Walter, 1952; Heller et al., 1988; He, 2007). Therefore, unconformities are excellent archives of plate boundary interactions and geodynamic switches in subduction zones (Wilson, 1991; Zhang et al., 2015).

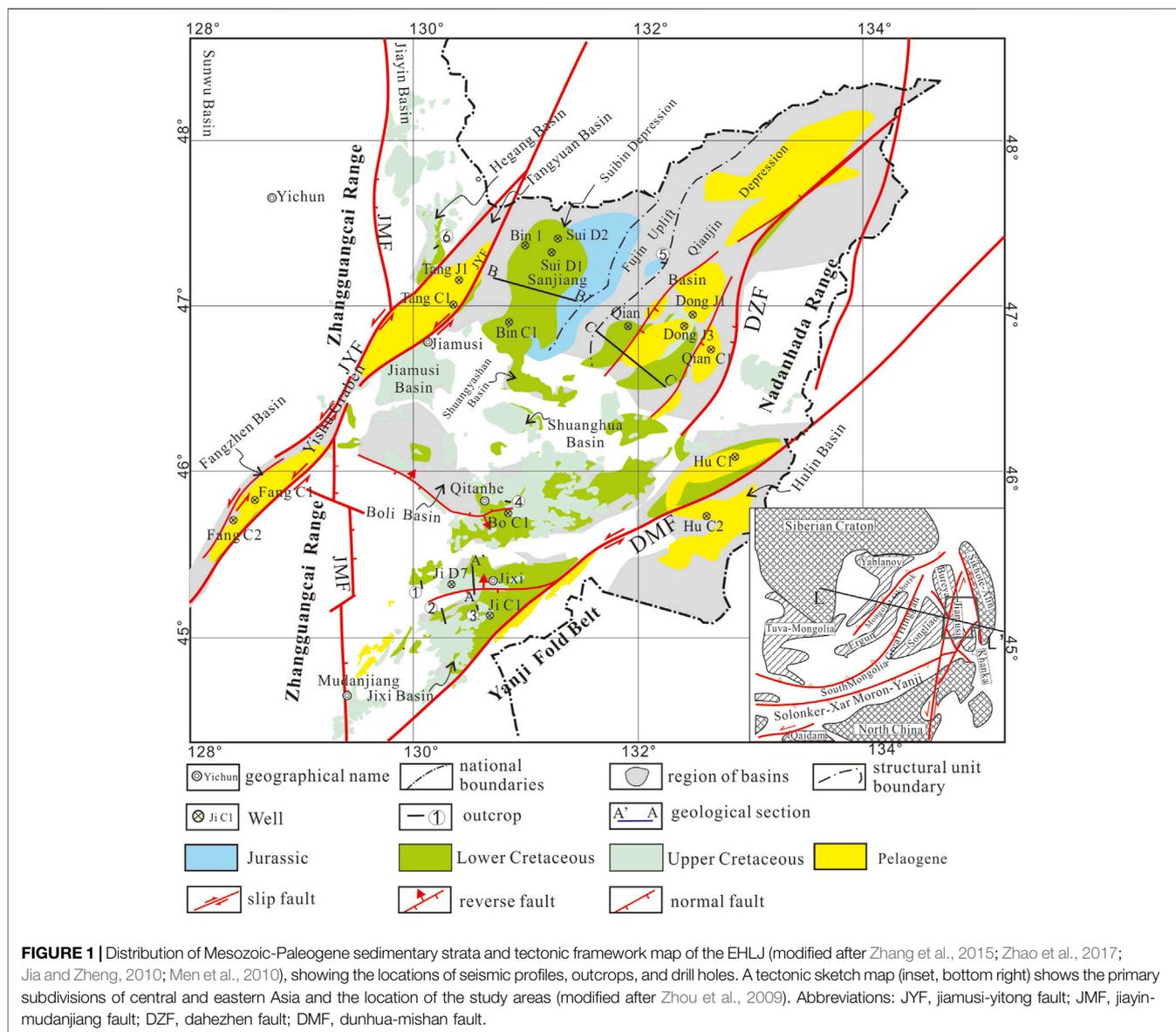
In the Eastern Heilongjiang Province (EHLJ), NE China, a series of Mesozoic–Cenozoic residual basins are distributed in the pre-Jurassic basement (Zhao et al., 2017). With the well-preserved Mesozoic–Cenozoic successions, the basins provide significant insights into Northeast Asia's structure characteristics and tectonic. Mesozoic unconformities and large-scale contractional structures in the basins mark a series of crucial tectonic transitions (Zhang et al., 2012; Zhang et al., 2015; Zhao et al., 2017). This study provides comprehensive synthesis information of regional Mesozoic unconformities from high-resolution seismic profiles and field investigations. Then the tectonic kinematics of Northeast Asia during the Mesozoic are discussed and analysed, and the tectonic characteristics and geodynamic background of the Mesozoic continental margin basin's evolution in Northeast Asia are explored.

2 GEOLOGICAL SETTING

Located in NE China, the EHLJ contains several adjacent Mesozoic–Cenozoic basins, including Sanjiang, Boli, Jixi, Hulin, Hegang, Tangyuan, and Fangzheng basin (Jia and Zheng, 2010; Men et al., 2010). They are separated by the NNE-trending Jiamusi–Yitong Fault (JYF) and the NS-trending Jiayin–Mudanjiang Fault (JMF). West of the EHLJ is the Zhangguangcai Range, and east is the Nadahada Range, between which there is the NS-trending Dahezhen Fault (DZF). Southeast of the EHLJ is the Yanji Fold Belt with the NE-trending Dunhua–Mishan Fault (DMF) between them (**Figure 1**).

The EHLJ's basement comprises three large tectonic units, including the Jiamusi, Khanka, and Bureya blocks (Zhang and Shi, 1993; Zhang and Zhang, 1999; Wang et al., 2009; Meng et al., 2010; Zhang et al., 2010), primarily comprising of the Archean Mashan group and the Late Hercynian to Indo-Chinese epoch Heilongjiang group Complexes and Permian granites (Zhou et al., 2010; Sun et al., 2013). The EHLJ's sedimentary cover the contains marine and terrestrial deposits of the Upper Jurassic–Cenozoic (Kirillova, 2003; Yang et al., 2007; Zhu and He, 2007; Sha and Galina, 2009).

Figure 2 summarises the inter-correlation of stratigraphic units of the EHLJ's sub-basins (Zhao et al., 2017). The



Middle-Upper Jurassic sequence includes terrestrial deposits, including the Suibin (Peide) Formation (Fm) and the Dongrong (Qihulin) Fm. **Figure 2** shows they are found only in the northern Sanjiang basin and the eastern Hulin basin (**Figure 1**). The Lower Cretaceous sequence is widely distributed in EHLJ, including the Didao Fm, the Chengzihe Fm, the Muleng Fm and the Dongshan Fm. However, complete successions are conserved only in the Hulin basin, the Jixi basin, and part of the Sanjiang basin. The Didao, the Chengzihe and the Muleng Fms were deposited in terrestrial environments from the Hauterivian to early Albian (Sha and Galina, 2009). Unconformably overlying the Muleng Fm, the Dongshan Fm is dominated by tuffs and basalts (also Albian). Locally distributed in the Sanjiang, the Hegang, the Jixi, and the Boli basins, the Upper Cretaceous sequence in the EHLJ primarily comprises the Houshigou Fm (also the Qixinghe and Yanwo

Fms). Multi-coloured, sandy conglomerates deposited in a fluvial environment dominate this Cenomanian to Santonian sequence.

3 METHODOLOGY

This study's methods and new integrated data include outcrop surveys, drill core observations, well logging data analysis and 2D/3D seismic profile interpretations.

The study area's stratigraphic information was based on the 1:200000 China Regional Geological Survey Reports. The outcrop observations along the basin margins were from fieldwork. Most structural information about outcrops was collected from investigating exposed sections in several quarries and open-pits that benefit from exploitation depth and variably orientated walls, allowing for 3D analyses.

System	Series	Stage	Age (Ma)	Fangzheng Basin	Tangyuan Basin	Hegang Basin	Sanjiang Basin		Shuangyashan Basin	Boli Basin	Jixi Basin	Hulin Basin	Evolutionary Stage of Basin	
							Suibin Depression	Qianjin Depression						
Cretaceous	Upper	Maastrichtian	72.1	U4 (Inversion)									Local Depression (Compression after Rifting)	
		Campanian								Hailiang ?	Hailiang ?			
		Santonian												
		Coniacian			Houshigou						Houshigou	Houshigou		
		Turonian												
	Lower	Cenomanian	100	U3 (Inversion)									Sag Rifting (Inversion)	
		Albian			Dongshan	Dongshan	Dongshan	Dongshan	Dongshan	Dongshan	Dongshan	Dongshan		
		Aptian			U3-1									
		Barremian			Muleng FM	Muleng FM	Muleng	Muleng	Muleng	Muleng	Muleng	Muleng		Zhushan
		Hauterivian	134			Chengzihe	Chengzihe	Chengzihe	Chengzihe	Chengzihe	Chengzihe	Chengzihe		Shangyunshan
Jurassic	Upper	Valanginian		U2									Sag	
		Berriasian	145											Qihulin
		Tithonian					Dongrong							
Jurassic	Lower	Kimmeridgian											Rifting	
		Oxfordian	164											Peide FM
Basement				U1										
				Mashan and Heilongjiang Complexes and Permian granites										

FIGURE 2 | Stratigraphic classification and comparison of the primary basins in the EHLJ (modified after Sha and Galina, 2009; Zhang et al., 2015; Zhao et al., 2017; He et al., 2009). Chronostratigraphic correlations are from the 2013 Geological Time Scale (Walker et al., 2013). Note: U1–U4 are unconformities; U1 = Jurassic/basement; U2 = Cretaceous/Jurassic; U3 = Upper Cretaceous/Lower Cretaceous; U3-1 = Dongshan Fm/Muleng Fm; U4 = Paleogene/Cretaceous.

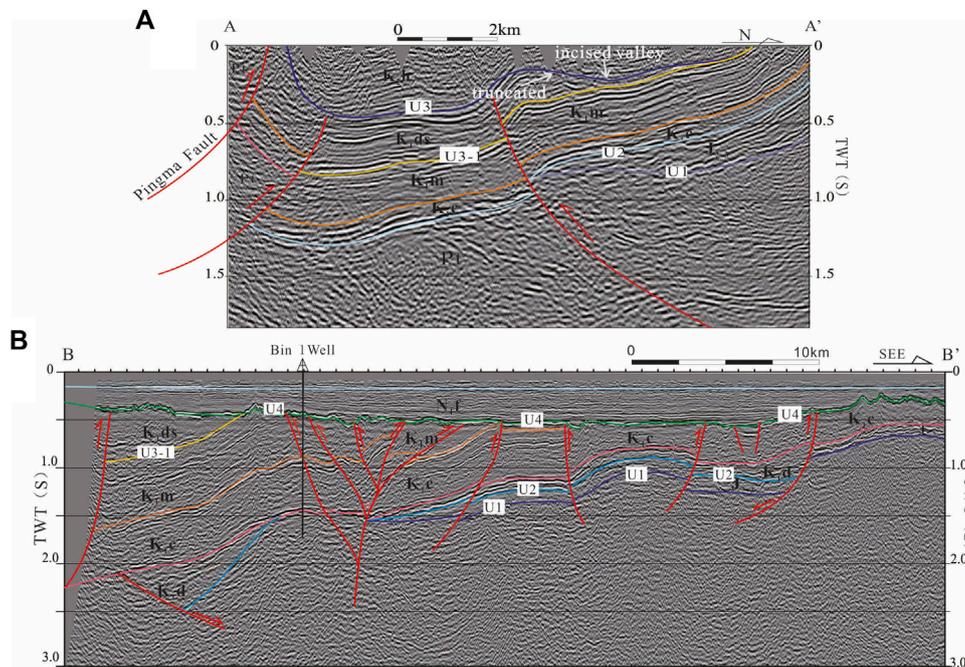


FIGURE 3 | Geological interpretations of seismic reflection profiles A-A' and B-B' (after Zhao et al., 2017) showing five unconformities (i.e., U1, U2, U3, U3-1, and U4). (A) showing four unconformities in Mesozoic. (B) showing five unconformities. Furthermore, these profiles show a transition from fault (half) graben of the K_{1d} period to a depression of the K_{1c} and K_{1m} periods. Symbols and abbreviations: J₂₋₃, Middle-Upper Jurassic; K₁, Lower Cretaceous; K₂, Upper Cretaceous; K_{1d}, Didao Fm; K_{1c}, Chengzihe Fm; K_{1m}, Muleng Fm; K_{2s}, Dongshan Fm; K_{2h}, Houshigou Fm; E_{1-2b}, Wuyun Fm and Xin'Anchun Fm; E_{2d}, Dalianhe Fm; E_{3b}, Baoquanling Fm; N_{1f}, Fujing Fm. See Figure 1 for the location.

Additional information was obtained from numerous road slopes.

The Daqing Oilfield Co. Ltd., a subsidiary of PetroChina provided the seismic reflection data and related drilling data. With a length of over 2,000 km, 3D and 2D seismic reflection data

mostly stretch across the central zone of the basins and were used to analyse their subsurface structures. Furthermore, drilling data and cores were from over 200 oilwells supplied by the Daqing Oilfield, which provided the related geological information for seismic interpretation, and constraining the strata units.

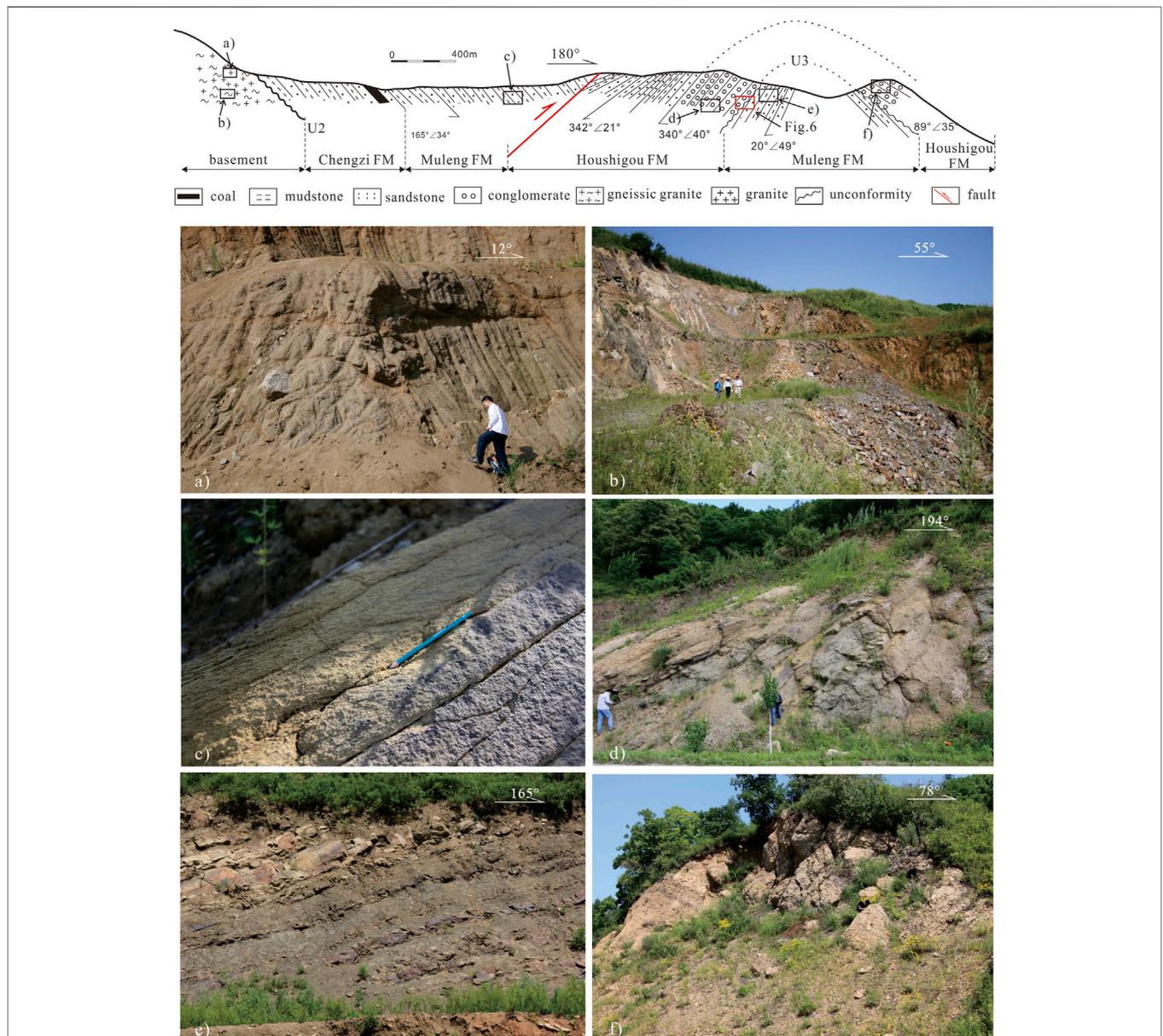


FIGURE 4 | Outcrop on the side of the road from Guangxin Town to Hongxin Town, showing two unconformities between the Lower Cretaceous Chengzihe Fm and the Pre-Mesozoic basement, and between the Houshigou and Muleng Fms. An anticline on the upper outcrop is also shown. **(A)** Proterozoic migmatitic granite, which has been cut by multiple granite dikes. **(B)** Strongly deformed Proterozoic metamorphic rock containing dark grey schist, gneiss, and marble. **(C)** Grey and grey-white coarse sandstone of the Muleng Fm, oblique bedding is developed. **(D)** Interbedded conglomerate and sandstone of the Houshigou Fm. **(E)** The Muleng Fm, the upper part is grey and grey-white coarse sandstone, the lower part is grey-black siltstone and silty mudstone. **(F)** The thick beds of conglomerate of the Houshigou Fm.

4 MESOZOIC UNCONFORMITIES

4.1 Middle-Upper Jurassic/Basement Nonconformity (U1)

The Middle-Upper Jurassic Dongrong and the Suibin Fms in the EHLJ are limitedly distributed in the Sanjiang Basin’s Suibin Depression. However, the Peide and Qihuilin Fms have limited

distribution in the Hulin Basin (Figure 1). Therefore, the U1 nonconformity can only be found in some areas of the Sanjiang and Hulin basins. The main characteristic of the selected seismic profile is that a regime of uniform reflections in the underlying basement changes to chaotic and sub-parallel reflections in the overlying Middle-Upper Jurassic sequence (Figure 3B).

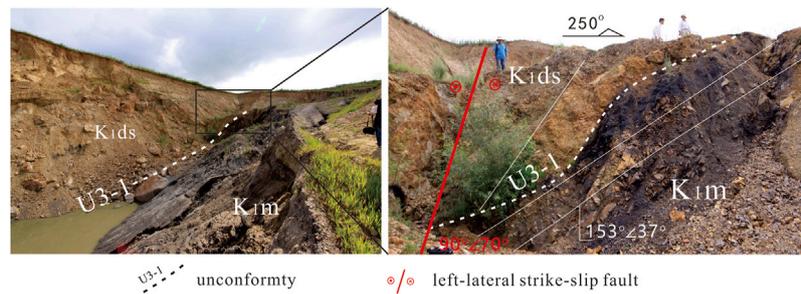


FIGURE 5 | The outcrop in an abandoned mine in Dongxing Town, Hegang City. It shows an unconformity between the Lower Cretaceous Dongshan and Muleng Fms, and a strike-slip fault cutting the strata of an attitude of $90^{\circ}\angle 70^{\circ}$. The right picture showing the Muleng Fm contains thick grey-white sandstone intercalated with a coal seam below the (angular) unconformity, whereas the Dongshan Fm comprises of disorderly conglomerates above the unconformity.

4.2 Lower Cretaceous/Jurassic Angular Unconformity (U2)

The unconformity found at the base of the Lower Cretaceous Didao and Chengzihe Fms is angular to the underlying strata. In the selected seismic section of the Suibin Depression, the Dongrong and the Suibin Fms are chaotic and sub-parallel reflections, whereas the overlying Didao and the Chengzihe Fms are parallel and sub-parallel reflections (**Figure 3B**). Furthermore, the chaotic characteristics in the lower strata turn into parallel and sub-parallel reflections in the Didao and the Chengzihe Fms in the Jixi Basin (**Figure 3A**). Some outcrops show that the Lower Cretaceous Muleng Fm directly covered the Pre-Mesozoic basement (Zhao et al., 2012).

Because Jurassic is only confined to the interior of individual basins and is absent at the basin margins, only the nonconformity between the Chengzihe Fm and the underlying ancient basement can be seen in the field outcrop. A typical outcrop is exposed to the side of the road from Guangyi Town to Hongxin Town (**Figure 4**). Two sets of rocks occur under the surface of the unconformity: one is a Proterozoic migmatitic granite cut by multiple granite dikes; the other is strongly deformed metamorphic rock containing dark grey schist, gneiss, and marble (**Figure 4B**). Overlying the unconformity, the Lower Cretaceous Chengzihe and the Muleng Fms are primarily grey and grey-white coarse sandstone and a coal seam (**Figures 4C,E**). The bedrock's orientation of the Muleng Fm is $165^{\circ}\angle 34^{\circ}$. In the upper outcrop, the Muleng and Houshigou Fms are involved in a large anticline. The Houshigou Fm comprises conglomerate or interbedded conglomerate and sandstone (**Figures 4D,F**) at $340^{\circ}\angle 40^{\circ}$ and $89^{\circ}\angle 35^{\circ}$, respectively. It is speculated that a major thrust fault has developed in the middle of the section.

4.3 Lower Cretaceous Dongshan Fm/Muleng Fm Disconformity or Angular Unconformity (U3-1)

The U3-1 unconformity is on the top of a sequence of the delta plain and lacustrine facies in the Muleng Fm and is overlain by pyroclastic rocks of the Dongshan Fm (Sha and Galina, 2009; Zhang et al., 2016). The 1:2,00,000 China Regional Geological

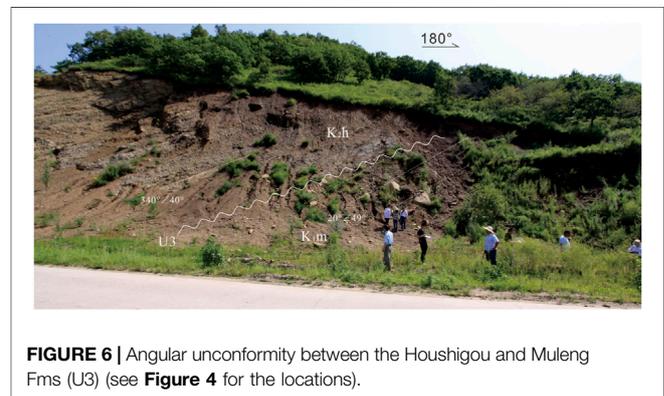


FIGURE 6 | Angular unconformity between the Houshigou and Muleng Fms (U3) (see **Figure 4** for the locations).

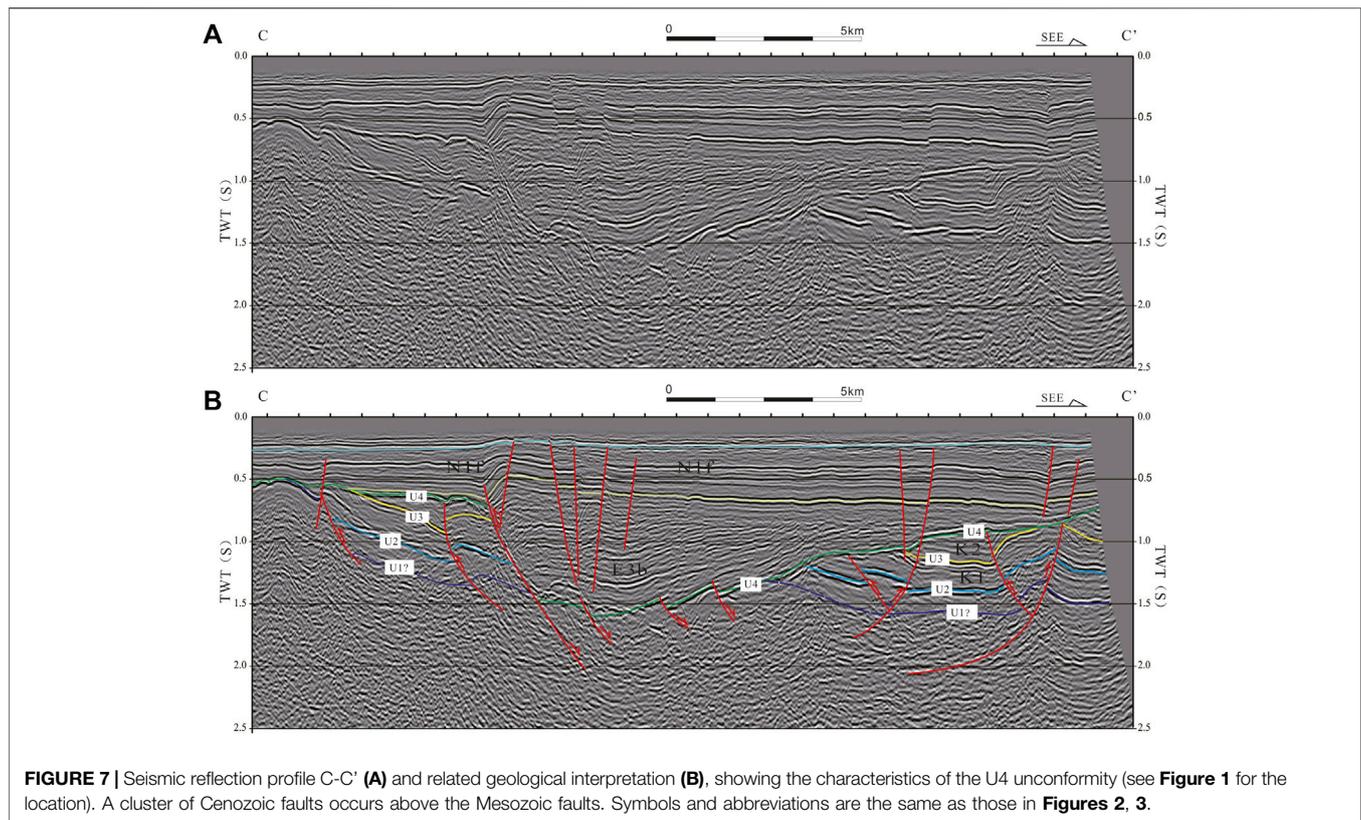
Survey Reports and field outcrops show the conformity, disconformities or local angular unconformities. The seismic reflection profiles in the Jixi and Sanjiang basins show that the lower reflection of the U3-1 has no noticeable erosion, a type of contact with either conformity or disconformity (**Figure 3**).

An angular unconformity occurs the outcrops between the Dongshan and Muleng Fms (**Figure 5**). The underlying Muleng Fm contains thick grey-white sandstone intercalated with a coal seam whose attitude is $153^{\circ}\angle 37^{\circ}$. However, the Dongshan Fm comprises disordered conglomerates, containing mainly of grey-white sandstones and lava (felsite) pebbles and cobbles with round shapes.

4.4 Upper Cretaceous Houshigou Fm (Qixinhe Fm)/Lower Cretaceous Angular Unconformity (U3)

Outcrops with angular unconformity between the Dongshan and the Muleng Fms are widely distributed in the Jixi and the Boli basins. In outcrop 2 (**Figure 4**), the Muleng Fm underlies the unconformity, with an attitude of $340^{\circ}\angle 40^{\circ}$ and the Houshigou Fm overlies the unconformity, with an attitude of $20^{\circ}\angle 45^{\circ}$ (**Figure 6**).

On the selected seismic reflection profile in **Figure 3A**, the strata underlying the unconformity's surface have been denuded, and an incised valley can be observed in some areas.



4.5 Paleogene/Cretaceous Angular Unconformity (U4)

The U4 unconformity is widespread in the EHLJ, noticeable in seismic profiles. For instance, an angular unconformity occurs between the Paleogene (or Neogene) and Cretaceous in Figures 3, 7. Below the unconformity, thrust faults are well developed and the strata exhibit strong denudation and deformation due to the compression associated with the thrusting showing that the EHLJ underwent widespread tectonic activity at the end of the Cretaceous.

5 DISCUSSION

The complex evolution of the NE Asian margin involves deformations of extension, contraction and strike-slip related to active margin tectonics. Some geodynamic models have been previously established (Faure et al., 1995; Northrup et al., 1995; Wang, 1997; Kirillova, 2003; Wen et al., 2008); however there is a lack of sufficient evidence and systematic understanding of the evolution process since the Mesozoic. Based on the five regional Mesozoic unconformities, combined with the geotectonic background (Zhang et al., 2012; Zhang et al., 2015; Zhao et al., 2017), we correlate the identified unconformities with the Mesozoic continental margin in Northeast Asia and establish evolution models (Figures 8, Supplementary Material).

5.1 The Transitional Environment of Convergent Orogeny and Intra-Continental Extension During the Jurassic

The Mongolia-Okhotsk Ocean's northward subduction began at the end of the Lower Jurassic Period and ended during the Middle-Upper Jurassic Period (Zonenshain et al., 1990; Zorin, 1999), and the Ocean began to close (Huang et al., 2021) (Figure 8A). Located at the Southern Ocean's active continental margin (Zorin, 1999), the Ergun-Greater Khingan Range composite plates generated massive N-S thrusts during the subduction, while the Nadanhada Arc Terrane began to accrete onto the East Asian Continent (Whan, 2006; Isozaki et al., 2010; Khanchuk et al., 2016; Yang et al., 2017).

Crustal shortening prevailed in Northeast Asia during the Middle to early Upper Jurassic times in response to the continental collision of Siberia and the northern China-Mongolia tract along the Mongolia-Okhotsk suture (Meng, 2003) (Figures 8A,B). Following the subducting plate's partial melting, magma intrusion and large-scale volcanic activity were generated by lithospheric extension and mantle upwelling due to slab break-off from the Mongolia-Okhotsk subducting plate (Ma et al., 2017; Tang et al., 2018). During this period, a complex Middle and Upper Jurassic passive continental margin rift formed with a near NS direction in Northeast Asia (Figure 8B). Volcano-sedimentary formations of the Suibin and Peide Fms were deposited in the eastern basin, whereas

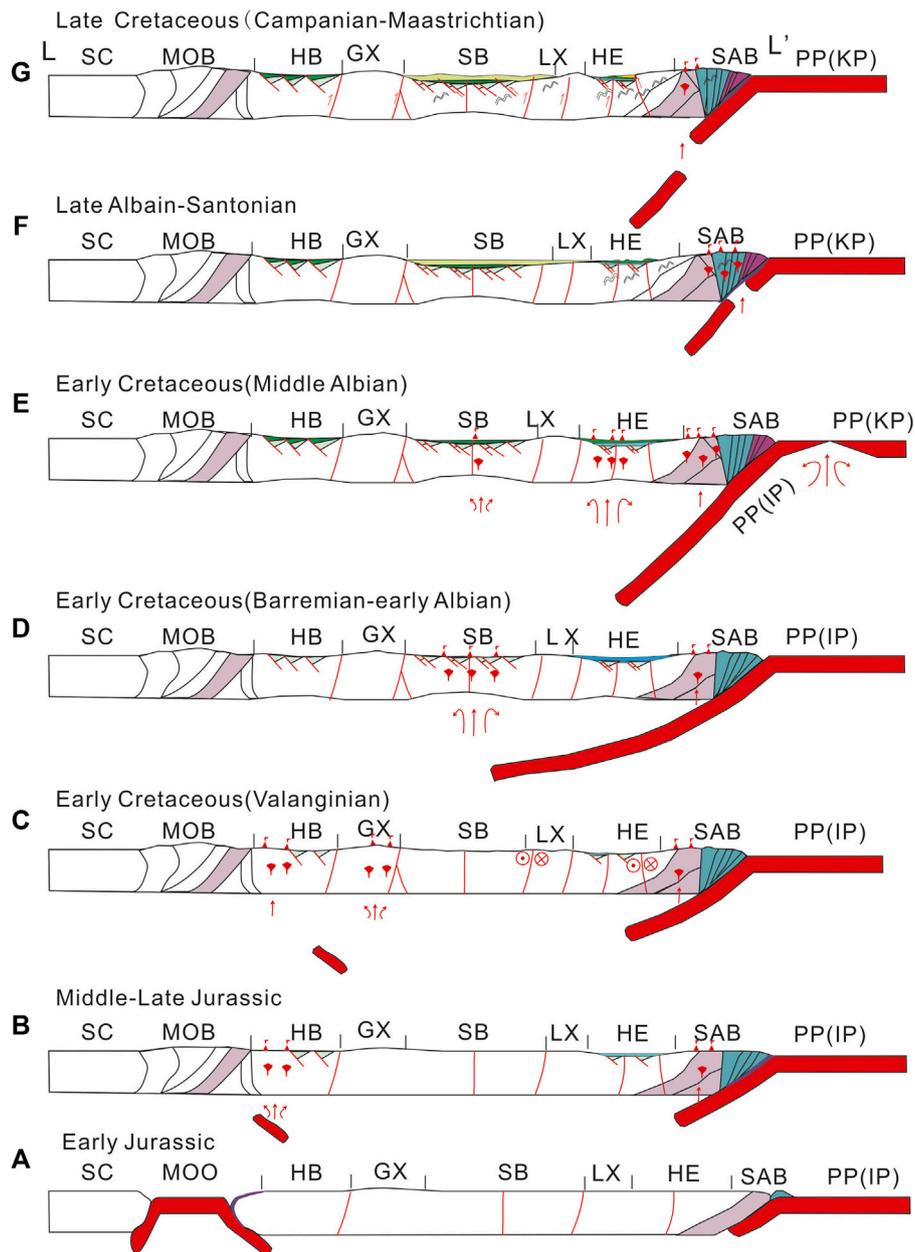


FIGURE 8 | The tectonic evolution of NE China and its geodynamic backgrounds during the Mesozoic era. It shows that two-sided active continental margin tectonics of the Mongolia–Okhotsk Belt suturing in the north, and the Pacific Plate subduction in the east controlled the evolution (Wang et al., 2016). It also shows the changes towards the increasing dominance of the Pacific Plate subduction (Li et al., 2012). Abbreviations: SC, siberian craton; MOO, mongolia–okhotsk ocean; MOB, mongolia–okhotsk orogenic belt; HL, hailer basin; GX, greater khingan range; SB, songliao basin; LX, lesser khingan range; HE, EHLJ; SAB, sikhotealin orogenic belt, PP, pacific plate, also called Izanagi (180–85 Ma) and Kula (85–70 Ma) plates during the Cretaceous (Maruyama et al., 1997); IP, izanagi plate; KP, kula plate.

those of the Tamulagou and Maketouebo Fms were deposited in the western basin (Figure 9).

During the Late Jurassic Period, the eastern basin went into a post-rift depressional stage with continuous subduction of the Pacific Plate (Huang et al., 2021; Li et al., 2021). Consequently, the Sanjiang basin connected with the Amur-Bureya basin and the Dongrong Fm’s littoral facies were deposited during a massive transgression (Jin et al., 2007).

5.2 Adjustments to the Paleo-Pacific Plate Subduction During the Early Lower Cretaceous

In the Lower Cretaceous (Valanginian), the Paleo-Pacific Plate rapidly subducted in an NNW direction towards the eastern margin of Eurasia (Maruyama et al., 1997). During the Asian continental plate convergence, the accreted-plates maintained a

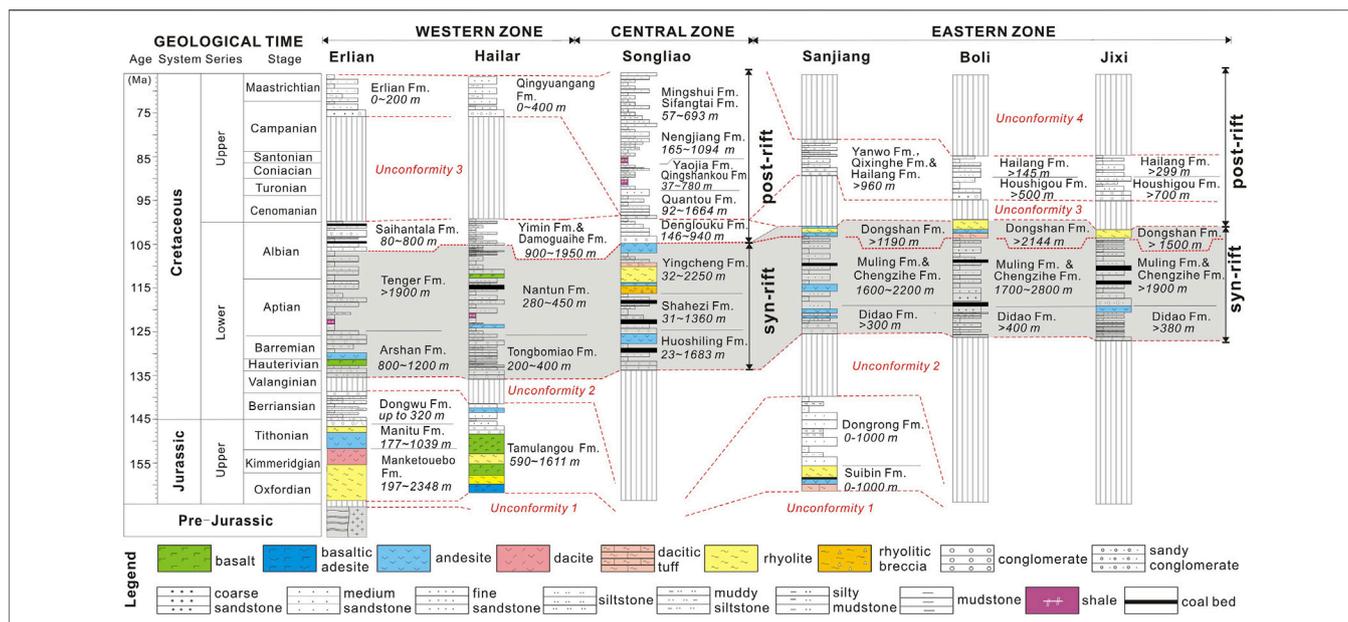


FIGURE 9 | Stratigraphy of seven representative late Mesozoic basins in NE China (Zhang et al., 2017b).

self-rotational motion due to continuous adjustment after accretion (Zhu et al., 2009; Hou et al., 2010; Yang, 2013), whereas there was significant strike-slip deformation between the plates (Halim et al., 1998; Metelkin et al., 2007; Metelkin et al., 2010), which could have been constrained by the Yangtze, North China, Pacific and Siberian Plates. The most significant strike-slips occurred during the Upper Jurassic to Lower Cretaceous, one of which was the Tanlu Fault Zone’s sinistral strike-slip ductile deformation (Uchimura et al., 1996). The strike-slip damaged the integrity of Eurasia’s eastern margin’s lithosphere and thinned of the Lower Cretaceous lithosphere. In Northeast Asia, it generated a wide shear stress field, causing numerous Palaeozoic suture zones to rejuvenate (Li et al., 2004; Zhu et al., 2006; Sun et al., 2008; Guo et al., 2018) (Figure 8C), resulting in a mass of strike-slip faults (Meng, 2003) and the widespread absence of early Lower Cretaceous (Valanginian) deposits (U2).

5.3 The Transition From an Extensional to Compressional Framework During the Middle-Late Lower Cretaceous

5.3.1 The Middle Lower Cretaceous

After the northward and north-westward subduction of the Pacific Plate, Northeast Asia’s the tectonic framework transformed from compressional to extensional during the Lower Cretaceous (Figure 8D). The thick deep crust and lithosphere underwent massive delamination from west to east (Wang et al., 2006; Zhang F. Q. et al., 2010; Zhang J. et al., 2010), resulting in the thinning of the lithosphere and intense

magmatism in Northeast Asia (Kirillova, 2003; Shao et al., 2007; Tang et al., 2018). Numerous igneous rocks and sag basins were formed as a result. During the Lower Cretaceous, intense magma upwelling induced extensional tectonic events, followed by thermal cooling subsidence, causing the Northeast Asia Mesozoic basins to enter the syn-rift stage.

As a significant Cretaceous intra-continental rift basin on the Eurasian continent in the central zone, the Songliao Basin was formed and filled under the control of mantle upwelling and rifting (Li and Liu, 2015; Zhang F. et al., 2010). The basin bounded by normal faults comprised isolated grabens or half-grabens. In ascending order, the syn-rift strata included the Huoshiling (K1h), Shahezi (K1s), and Yingcheng (K1y) Fms (Figure 9) and the sediment thickness reached 4,000 m in the south-eastern uplift zone (Xie et al., 2003).

In the EHLJ, the Jixi, Boli and Sanjiang basins deposited the Didao Fm during the early stage. During the Aptian phase, the basins expanded and underwent subsidence, entering the depression stage. During a widespread transgression, the Mesozoic-Cenozoic residual basins in the EHLJ (i.e., the Jixi, Boli, Jiamusi, Sanjiang, Hegang basins) might have become well connected, regarded as the Cretaceous ‘Grand Sanjiang Unified Lake Basin’ (Sha et al., 2003; Sha, 2007; Wen et al., 2008; Sha and Galina, 2009; Jia and Zheng, 2010; Zhao et al., 2012; Zhao et al., 2017) comprising the coal-bearing strata of the Chengzihe and the Muleng Fms. The Sikhote-Alin Orogenic Belt uplifted, and deep faults in Northeast Asia (such as the Tanlu fault) began a sinistral strike-slip activity. Local volcanic activity occurred at the intersection of the NE and EW orientated faults (Zhao, 2011).

5.3.2 The Late Lower Cretaceous Period

Because of the plate's slab roll-back during the late Lower Cretaceous (after the deposition of the Muleng Fm), local asthenospheric material upwelled, and fault and volcanic activities intensified (Li et al., 2007; He et al., 2009; Zhu Z. P. et al., 2009; Sun et al., 2013; Ling et al., 2017; Tang et al., 2018) (**Figure 8E**). Some scale of bimodal magmatism occurs in the Songliao Basin and the EHLJ, and the corresponding sedimentary strata are the Denglouku and Dongshan Fms. The Dongshan Fm was deposited in the EHLJ basins, containing of intermediate volcanic breccia, lithic tuff, altered basalt, siltstone, fine sandstone, and mudstone (**Figure 9**).

After the late Lower Cretaceous Period (Middle Albian), a massive collisional orogeny occurred in the eastern Eurasian margin (**Figure 8E**). The orogeny is named the Sakawa orogenesis in southwestern Japan (Kobayashi, 1941; Faure, 1985; Okada and Sakai, 2000) and the Minchen movement in south-eastern China (Gu, 2005). Based on research data, Vaughan (1995) highlighted that the collisional orogeny on the Pacific Rim occurred at 110–100 Ma, and generated wide thrust nappes and nonconformities. Sato et al. (2002) argued that the geodynamic and magmatic processes in Eurasia's east margin in this stage were controlled by the Izanagi and Pacific Plates' transition in configuration and style of convergence styles.

5.4 Compression System During the Upper Cretaceous Period

During the early Upper Cretaceous (Cenomanian), the Kula Plate began impinging on Eurasia and subducting after the Izanagi Plate entirely subducted under the Eurasia Plate, whereas the Pacific Plate subducted in a transformation from NNW to WNW (Maruyama et al., 1997; Northrup et al., 1995; Zhang et al., 2017a). These geodynamic movements shifted the Eurasian Plate's eastern margin into an Andean-type continental margin (Sato et al., 2002). Based on thermochronological data in the Songliao Basin, Song et al. (2015) and Zhao et al. (2013) proposed that the Pacific-induced deformation at 87–89 Ma encompassed the entire eastern Asia from the subduction boundary to the hinterland. During these episodes, the large deep thrust faults (Zhang and Dong, 2008; Sun et al., 2010) developed under the large-scale orogeny domination (Han et al., 2008; Shi et al., 2008). In the EHLJ, after being extensively uplifted, some basins, such as the Jixi and the Boli Basins, became separated and experienced erosion. However, the Songliao Basin entered a post-rift stage (**Figure 9**).

The Upper Cretaceous stratigraphy in some Northeast Asia basins reveals internal angular unconformities and folds and thrust/reverse faults (Cho et al., 2016), indicating that the basins underwent compressional deformation episodes (**Figure 8F**). Two unconformities were displayed in the EHLJ, including those of the early Cenomanian–Turonian and the early Campanian–Maastrichtian consistent with the event of two discrete regional contractions in the Upper Cretaceous (Zhang et al., 2017b). In the Songliao Basin, two significant episodes of drainage pattern switches were found in the Upper Cretaceous

strata, of which the thickness is up to 5 km, although no unconformities occurred in such a complete deposition (Wang et al., 2016; Zhang et al., 2017b).

With the subduction of the Kula Plate, a compression stress field entirely controlled NE China during the late Upper Cretaceous (**Figure 8G**). The basins underwent widely compressive deformation, accompanied by large-scale thrust, denudation and deplanation (Zhao et al., 2017). The intensive structural contraction lasted until the Paleogene, and divided the Cretaceous basin into individual ones, forming the present basin framework (Zhang et al., 2017b).

6 CONCLUSION

The data obtained seismic reflection profiles and field outcrops have revealed several regional Mesozoic unconformities in the EHLJ, NE China. The unconformities indicate that the two-sided active continental margin tectonics of the Mongolia–Okhotsk suture in the north during the Upper Jurassic and the Pacific Plate subduction in the east during the Cretaceous controlled the Mesozoic tectonic evolution of Northeast Asia.

- (1) The U1 reflects the transitional environment from a convergent orogeny to an intra-continental extension in Northeast Asia during the Jurassic in response to the continental collision of Siberia and the northern China–Mongolia tract along the Mongolia–Okhotsk suture. The Lower Cretaceous/Jurassic angular unconformity (U2) results from numerous Palaeozoic suture zones and the activity of strike-slip faults at the eastern margin and in the neighbouring area of Eurasia, and indicating that the Paleo-Pacific Plate rapidly subducted in the NNW direction towards the Eurasia's eastern margin.
- (2) During the late Lower Cretaceous Period, subduction slab rollback of the Paleo-Pacific Plate resulted in the upwelling of local asthenospheric material, intensifying the fault and volcanic activity. The Lower Cretaceous Dongshan Fm/Muleng Fm unconformity (U3-1) reflects a scale of bimodal magmatism in the Songliao Basin and the EHLJ. During the early Upper Cretaceous (Cenomanian), the Kula Plate began impinging on Eurasia and subducting after the Izanagi Plate entirely subducted under the Eurasia Plate, whereas the Pacific Plate subducted in a transformation from NNW to WNW. Some the basins in the EHLJ, such as the Jixi and the Boli Basins, became separated at that epoch and experienced denudation after being extensively uplifted (U3). With the subduction of the Kula Plate, compression stress field controlled NE China during the later Upper Cretaceous. The EHLJ basins underwent widely compressive deformation, accompanied by large-scale thrusts, denudation and deplanation, and consequently Paleogene/Cretaceous unconformity (U4) was formed.

DATA AVAILABILITY STATEMENT

The original contributions presented in the study are included in the article/**Supplementary Material**, further inquiries can be directed to the corresponding author.

AUTHOR CONTRIBUTIONS

XZ contributed to the conception of the study and field works; CL and YJ performed the interpretation of seismic; XZ, HZ, and FW wrote the manuscript; PC helped perform the analysis with constructive discussions.

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SUPPLEMENTARY MATERIAL

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