

# Tropical Cyclones Intensity Prediction in the Western North Pacific Using Gradient Boosted Regression Tree Model

#### Gangya Zhu<sup>1</sup>, Qinglan Li<sup>1</sup>\*, Wei Zhao<sup>1</sup>, Xinyan Lv<sup>2</sup>, Chuanhai Qian<sup>2</sup> and Qifeng Qian<sup>2</sup>

<sup>1</sup>Shenzhen Institute of Advanced Technology, Chinese Academy of Sciences, Shenzhen, China, <sup>2</sup>National Meteorological Center, Beijing, China

As an artificial intelligence method, machine learning (ML) has been widely used in prediction models of high-dimensional datasets. This study proposes an ML method, the Gradient Boosted Regression Tree (GBRT), to predict the intensity changes of tropical cyclones (TCs) in the Western North Pacific at 12-, 24-, 36-, 48-, 60-, and 72-h (hr) forecasting lead time and the model is optimized by the Bayesian Optimization algorithm. The model predictands are the TCs intensity changes at different forecasting lead times, obtained from the best track data of the Shanghai Typhoon Institute (STI) and the Joint Typhoon Warning Center (JTWC) from 2000 to 2019. The model predictors are the synoptic variables, climatological and persistent variables derived from the reanalysis data obtained from the National Centers for Environmental Prediction (NCEP), and the sea surface temperature (SST) data obtained from the National Oceanic and Atmospheric Administration (NOAA). The results show that the GBRT model can capture the TCs intensity changes well for the succeeding 12-h, 24-h, 36-h, and 72-h. Compared with the traditional multiple linear regression (MLR) model, the GBRT model has better performance in predicting TCs intensity changes. Compared with the MLR model,  $R^2$  of the GBRT model for TCs intensity forecast increases by an average of 8.47% and 4.45% for STI data and JTWC data. MAE (RMSE) drops by 26.24% (25.14%) and 10.51% (4.68%) for the two datasets, respectively. The potential future intensity change (POT), the intensity changes during the previous 12 h (Dvmax), Initial storm maximum wind speed (Vmax), SST, and the Sea-Land ratio are the most significant predictors for the GBRT model in predicting TCs intensity change over the Western North Pacific.

Keywords: tropical cyclones, intensity change, prediction, Western North Pacific, gradient boosted regression tree, Bayesian optimization

# **1 INTRODUCTION**

Tropical cyclones (TCs) usually bring substantial economic losses and casualties through strong winds, heavy rainfall, and storm surge during and after TCs landfall, especially in the coastal area which is vulnerable to TCs (Zhang et al., 2009). Therefore, an accurate and timely forecast of TCs track and intensity is crucial for local disaster mitigation. During the past century, scientists and meteorological administrations tried to improve their ability in TCs observation technology, forecasting techniques, and understanding TCs intensification

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> \***Correspondence:** Qinglan Li ql.li@siat.ac.cn

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TABLE 1 | The 26 potential predictors available for the GBRT model.

Predictor	Description				
Lat	Initial storm latitude				
Lon	Initial storm longitude				
Vmax	Initial storm maximum wind speed				
Dvmax	Intensity change during the previous 12 h				
JDATE	Absolute value of Julian day				
SHRS	Area-average (200–800 km) wind shear at 850–500 hpa				
SHRD	Area-average (200–800 km) wind shear at 850–200 hpa				
LSHRS	SHRS times the sine of the latitude				
LSHRD	SHRD times the sine of the latitude				
USHRS	Area-average (200–800 km) zonal wind shear at 850–500 hpa				
USHRD	Area-average (200–800 km) zonal wind shear at 850–200 hpa				
MPI_Vmax	MPI <sup>a</sup> times the initial intensity				
POT	Maximum possible intensity - initial intensity				
RHHI	Area-average (200–800 km) relative humidity at 500–300 hpa				
RHLO	Area-average (200–800 km) relative humidity at 850–700 hpa				
SPD	Storm translational speed				
T200	Area-average (200–800 km) temperature at 200 hpa				
D200	Area-average (200–800 km) divergence at 200 hpa				
WVF500	Area-average (200–800 km) water vapor flux at 500 hpa				
WVF850	Area-average (200–800 km) water vapor flux at 850 hpa				
Z500	Area-average (200–800 km) vorticity at 500 hpa				
Z850	Area-average (200–800 km) vorticity at 850 hpa				
U200	Area-average (200–800 km) zonal wind at 200 hpa				
V200	Area-average (200–800 km) meridional wind at 200 hpa				
SST	Area-average (0-800 km) sea surface temperature				
SL ratio	Ratio of sea area with a certain radius				

<sup>a</sup>MPI, refers to the maximum potential intensity, following the methodology employed by DeMaria and Kaplan (1994b) and Knaff et al. (2005).

mechanism (Emanuel, 2018). In the past 30 years, with the application of meteorological satellites and the popularity of ensemble forecasts for TCs track, the track prediction has been greatly improved (Zhang and Krishnamurti, 1997; Fraedrich et al., 2003; Langmack et al., 2012; Jun et al., 2017). However, the intensity forecast is still poor compared with the track forecast and is a huge challenge around the world (DeMaria et al., 2014).

TABLE 2 | Sample numbers for different TCs' lead time prediction models.

Lead time	Source	Over land	Near the	Over open ocean	All samples
			coast		
12	STI	814	6782	6525	14121
	JTWC	384	6388	6437	13209
24	STI	547	6272	6220	13003
	JTWC	175	5746	6143	12064
36	STI	_	5592	5931	11891
	JTWC	_	5016	5853	10946
48	STI	_	4920	5639	10814
	JTWC	_	4273	5547	9860
60	STI	_	4257	5340	9785
	JTWC	_	3557	5228	8809
72	STI	_	3622	5041	8803
	JTWC	_	2925	4875	7819

TCs intensity is influenced by two main physical processes (Wang and Wu, 2004; Elsberry et al., 2013), which are synoptic variables such as vertical wind shear, humidity, sea surface temperature, water vapor, divergency (DeMaria, 1996; Ge et al., 2013; Gao et al., 2016; Mercer and Grimes, 2017), and climatological and persistent variables such as latitude, longitude, Julian day, and sea-land ratio (SL ratio) (DeMaria and Kaplan, 1994a; Gao et al., 2016; Li et al., 2018). Statistical-dynamical models were used to predict TCs intensity and rapid intensification probability and outperformed the forecasts of individual physics-based dynamical models (Knaff et al., 2005; Kaplan et al., 2010; Gao and Chiu, 2012). For decades, scientists have dedicated themselves to improving the skill of TCs intensity forecast. Jarvinen and Neumann (1979) proposed a statistical regression equation (Statistical Hurricane Intensity FORecast, SHIFOR) using predictors derived from climatic and persistent variables to predict TCs intensity changes over the North Atlantic Basin for the future 72 h. DeMaria and Kaplan (1994a) proposed a Statistical Hurricane Intensity Prediction Scheme (SHIPS), which considered more synoptic predictors to predict the changes in TCs intensity over the Atlantic Ocean Basin. The



FIGURE 2 | Selection results of the significant variables for TCs intensity prediction models for the STI datasets by the GBRT method. The vertical axis represents the percentage of the feature importance of the variable in the corresponding prediction model, and the horizontal axis refers to the predictor variable. Only the six most important variables are displayed for each lead time forecasting model. (A) for TC samples over land; (B) for TC samples near the coast; (C) for TC samples over the open ocean; (D) for all TC samples.

results showed that the average errors by SHIPS were 10%–15% smaller than the errors by the SHIFOR model that used only climatic and persistent variables. The storm decay over land was further considered by SHIPS (DeMaria et al., 2005). Besides the conventional synoptic and climatological variables, Li et al. (2018) paid close attention to the land effect on TCs intensity change by proposing a new factor involving the ratio of seawater area to land area (SL ratio) in the statistical regression model. TCs intensity changes over the entire TCs life span, including over the ocean basin, near the coast, and after landfall, were considered in

the model (Li et al., 2018). Intensity forecasting accuracy for TCs near the coast and over land was improved with the addition of the SL ratio, compared with that of the models that did not consider the index of SL ratio (Li et al., 2018).

Based on previous observations and studies, in most cases, intensity changes are usually slow and steady over a certain period. However, in some situations, the TCs intensity may vary rapidly. Predicting these rapid intensifications is very challenging (Emanuel and Zhang, 2016). Many factors, such as the complex and chaotic energy exchange process between the sea





and the atmosphere, and the imperfection of real-time data collection, may influence the prediction of TCs intensity change. These rapid intensification processes are challenging to explain, and therefore, neither numerical nor statistical-dynamic models can predict TCs intensity changes accurately (Chen et al., 2020). A vital weakness for numerical methods is insufficient representation of the complicated dynamical process; however, increasing the number of variables or equations would exponentially demand the computation (Gao et al., 2016). On the other hand, statistical methods, usually based on regression and

lower computational costs, may not be effective in capturing nonlinear relationships. Therefore, their forecast results need to be further improved (Lin et al., 2009; Sandery et al., 2010).

To solve these problems, scientists began to use machine learning (ML) to predict the TCs intensity in recent years (Gao et al., 2016; Cloud et al., 2019; Chen et al., 2020; Xu et al., 2021). Some ML algorithms, such as support vector machine, artificial neural network, decision tree, and random forest, have been gradually introduced into meteorology (Breiman, 1996; Mas and Flores, 2008; Behrangi et al., 2009;



Validations in terms of R<sup>-</sup> (the first column), MAE (the second column), and RMSE (the third column) based on the STITCS best track data. (A-C) for TC sal land; (D-F) for TC samples near the coast; (G-I) for TC samples over the open ocean; (J-L) for all TC samples.

Mountrakis et al., 2011; Pan et al., 2019). ML is applicable to highdimensional datasets and can deal with nonlinear relationships between predictors and predictand. It is used to explore satellite, radar, and *in-situ* data to improve the TCs intensity forecast skills (Gao et al., 2016; Zhang et al., 2016; Griffin et al., 2017; Jin et al., 2020). Generally, ML models combine historical storm characteristics, physically motivated information, and stormspecific features to make predictions (Xu et al., 2021).

The gradient boosted regression tree (GBRT) is an ensemble ML algorithm consisting of multiple decision trees, which is robust in model training (Zhang et al., 2020). First, the decision tree algorithm is a classical ML method, and it can deal with inherent nonlinear relationships in variables and missing values (Quinlan, 1987; Fayyad and Stolorz, 1997). In addition, the algorithm can quantify the relative importance of variables and establish decision rules for prediction (Quinlan, 1987). It has been successfully applied in the analyses of the re-curvature, landfall, and intensity change of TCs in the Western North Pacific Ocean (Zhang et al., 2013; Gao et al., 2016). Second, like most ensemble methods, a combination of decision trees can provide more robust and accurate regressions than a single one. The greater-than and less-than structures of the tree module make the GBRT less affected by outliers (Bishop and Nasrabadi, 2006; Ma et al., 2018). Furthermore, GBRT is able to capture the complicated and

nonlinear relationships between TCs intensity change and other related features. Compared with a single decision tree, GBRT pays more attention to the regression errors with less calculation time for high-dimension data (Xie and Coggeshall, 2010; Ding et al., 2016; Yang et al., 2016; Ma et al., 2017; Ma et al., 2018). Therefore, GBRT is used in this study to predict the future change of TCs intensity in the Western North Pacific Ocean.

Most of the studies mainly focused on TCs intensity change over the open ocean. However, TCs that make landfall or approach the coast usually cause most of the loss of life and damage. Forecasting the intensity of TCs near the coast and over land should be, therefore, more critical than forecasting TCs intensity over the open ocean (Li et al., 2018). Li et al. (2018) explored the TCs intensity change over the entire TCs life span by considering a new factor, the "ratio of seawater area to the land area," in the multiple linear regression model (MLR). However, MLR usually performs poorly in handling the nonlinear relationship between predictors and predictand. In this study, the GBRT is proposed to forecast future TCs intensity at 12-, 24-, 36-, 48-, 60-, and 72-h forecasting lead time in the Western North Pacific region for different TCs life spans. GBRT's performance in predicting TCs intensity change over the Western North Pacific is compared with the performance of the MLR model used by Li et al. (2018). The remaining of the paper is as



FIGURE 5] Evaluation of the performance of the GBR1 model in predicting the 1Cs tuture 12, 24, 36, 48, 60, and 72-h intensity changes with four cross validations in terms of R<sup>2</sup> (the first column), MAE (the second column), and RMSE (the third column) based on the JTWC TCs best track data. (A–C) for TC samples over land; (D–F) for TC samples near the coast; (G–I) for TC samples over the open ocean; (J–L) for all TC samples.

follows. The data and methodology are described in **Section 2**. **Section 3** contains the results and discussion. In the final section, summaries are presented.

# **2 DATA AND METHODS**

#### 2.1 Data

At present, there is no standardized TCs intensity estimation method, and the most widely used methods are artificial algorithms based on satellite cloud images; therefore, there is uncertainty about TCs intensity values (Jiang et al., 2019; Lee et al., 2019). To account for the uncertainty, this study used two different TCs datasets: the best track data from the Shanghai Typhoon Institute (STI, 2020) and the best track data from the Joint Typhoon Warning Center (JTWC, 2020) for the period of 2000–2019 in the Western North Pacific region, covering the area with latitude north of 0°N and longitude west of 180°. It is worth noting that for TCs that affect China, STI data might have more advantages than JTWC, as there are more direct observations for those TCs in China (Ren et al., 2011).

The best track datasets include the TCs time (year, month, day, hour), location (latitude and longitude of the TCs center), TCs central pressure, and the maximum sustained wind speed near the TCs center. Figure 1 shows the tracks of the TCs in the Western North Pacific from 2000 to 2019 based on the data from STI. In addition to the TCs data, synoptic variables, climatological and persistent variables are derived from the reanalysis data obtained from the National Centers for Environmental Prediction (NCEP, 2020) at 6-h temporal and  $1^{\circ} \times 1^{\circ}$  spatial resolution. The NCEP FNL (Final) Operational Global analysis data has incorporated the most complete set of observational data and is possibly the best choice for a long-term operational model archive from NCEP (NCEP, 2020). The analyses are available at the surface and 26 mandatory levels from 1,000 to 10 hpa. Furthermore, this study applies global daily means of SST provided by the National Oceanic and Atmospheric Administration (NOAA, 2020), as warm seawater is the energy source for TCs. The data from 2000 to 2019 are divided into four groups according to five consecutive years (i.e., 2000-2004 is a group; 2005–2009 is another group). We use one group as the test set and the other three groups as the training set to build the prediction model and perform cross-validation.





#### 2.2 Method 2.2.1 Gradient Boosted Regression Tree Model

In this study, the Gradient Boosted Regression Tree (GBRT) model is applied to predict the TCs intensity change over the entire TCs life span. The GBRT model is popular for its ability to describe the complicated relationships between input and output data and the explanation of input features (Yang et al., 2020). It enhances the traditional decision tree approach by boosting technology (Friedman, 2001; Friedman, 2002). In boosting, base learners are built sequentially, and each base learner tries to reduce the bias of the previous combined learner (Yang et al., 2020). This approach can combine multiple weak models to make the ensemble model more powerful (Zhou et al., 2021). One of the core ideas of GBRT is to use the value of the negative gradient of the loss function in the current model as an approximation of the residual, which is essentially a first-order Taylor expansion of the loss function to fit a regression tree. Besides, the samples in the training set with the largest residuals are weighted the most heavily in GBRT (Schapire, 2003), encouraging the model to improve its worst predictions (McGovern et al., 2019). In addition, the importance of each input variable can be ranked in GBRT model.

The GBRT is additive models, which can be expressed as the following form:

$$\widehat{\mathbf{y}}_i = F_M(\mathbf{x}_i) = \sum_{m=1}^M h_m(\mathbf{x}_i). \tag{1}$$

where the  $h_m$  are estimators, which are called weak learners in boosting. The GBRT uses fixed-size decision tree models as weak learners. The constant M corresponds to the parameters of the estimator.

Then the GBRT learners are updated continuously, similar to other boosting algorithms.

$$F_m(x) = F_{m-1}(x) + h_m(x),$$
 (2)

where the newly added tree  $h_m$  is fitted to minimize the sum of losses  $L_m$ , based on the previous ensemble  $F_{m-1}$ :

$$h_m(\mathbf{x}) = \underset{h}{arg\min} L_m = \underset{h}{arg\min} \sum_{i=1}^n l(y_i, F(\mathbf{x}_i) + h(\mathbf{x}_i)). \quad (3)$$

#### 2.2.2 Bayesian Optimization Algorithm

Bayesian optimization is a very effective global optimization algorithm whose goal is to find the global optimal solution, which is a kind of black-box function without assuming any function form (Brochu et al., 2010; Snoek et al., 2012). As ML



near the coast, over the open ocean, and all TC samples, respectively.

models are widely used to process large amounts of data, the space and process required for hyperparametric tuning become more and more complex. Nowadays, people prefer Bayesian optimization parameter tuning which has been shown to lead to high model performance after the convenient and efficient tuning (Swersky et al., 2013). The corresponding algorithm for the Bayesian optimization is as follow:

Suppose we have a function  $f: X \to R$  and we need to find the maximum within  $x \in X$ :

$$\boldsymbol{x_{opt}} = \underset{\boldsymbol{x} \in X}{\operatorname{arg\,max}} f(\boldsymbol{x}), \tag{4}$$

where X is a set of hyperparameters, and  $X \in \mathbb{R}^d$ . The computational cost of evaluating model changes is high with a usually small d, and hyperparametric gradients are often not available; therefore, hyperparametric optimization of ML models is important (SIGOPT, 2020). It is worth mentioning that the Bayesian optimization algorithm remains robust for random, non-convex, and even discontinuous fundamental functions f. In this study, we use the Bayesian optimization algorithm to iteratively tune the hyperparameters of the GBRT model, and apply the optimal hyperparameter combination to the prediction model, thereby improving the accuracy of the model.

#### 2.2.3 Potential Predictors

The 26 potential predictors used in the MLR model by Li et al. (2018) are also applied for the GBRT model in this study, which are listed in Table 1. All synoptic variables are generally averaged from the corresponding data within a specific radius of the TCs center, usually within a ring with an outer diameter of 800 km and an inner diameter of 200 km. Note that SST is calculated as the average value of the data with a radius of 800 km. The maximum potential intensity (MPI) and TCs potential future intensity change (POT) are calculated by SST. Following Li et al. (2018), the life spans of TCs are divided into the samples over the open ocean, near the coast, and over land according to the values of SL ratio with the range of SL ratio > 0.99, 0.5 < SL ratio < 0.99, SL ratio < 0.5, respectively, which are computed within the radius of 500 km around the TCs center. The numbers of TC samples over the open ocean, near the coast, and over land, as well as the total TC samples (all TC samples without categorization), for the TCs intensity change prediction models with 12-, 24-, 36-, 48-, 60-, and 72-h forecasting lead time are shown in Table 2.

The performance of the forecasting models for the testing periods are evaluated in terms of the statistical properties of



coefficient of determination (commonly known as  $R^2$ ), mean absolute error (MAE) and root mean square error (RMSE).

# **3 RESULTS AND DISCUSSION**

#### **3.1 Computation and Selection of the Model Predictors**

For the TCs intensity prediction models, the dependent variables are the TCs intensity changes for the future 12-h, 24-h, to 60-h, and 72-h, and the potential predictors are listed in **Table 1**. Because TCs usually decay quickly after their landfall, this study only considers TCs intensity changes within 24 h (including 12-h and 24-h) after landfall. **Figures 2**, **3** show the importance of the variables in the prediction models with different lead times over the open ocean, near the coast, over land, as well as the all TC samples, based on the STI and the JTWC datasets, respectively. As can be seen from **Figures 2**, **3**, the feature importance of variables varies for the STI and JTWC datasets with different prediction periods, and the six most important variables for each prediction period have accounted for more than 80% of the total feature importance for the prediction models.

Figures 2, 3 show that POT plays the most crucial role in TCs intensity change for TCs over the open ocean (Figures 2C, 3C) and near the coast (Figures 2B, 3B), as well as all TC samples (Figures 2D, 3D). The importance of POT is generally increasing with the forecasting lead time. Besides POT, SST is also essential for TCs intensity change for TCs over the open ocean and near the coast, especially based on the JTWC datasets. Dvmax is an important factor affecting TCs intensity change for TCs over the open ocean and near the coast, as well as all TC samples, for short forecasting lead time, i.e., 12 h and 24 h. Dvmax is no longer an important factor for TCs intensity changes when the forecasting lead time is longer than 24 h for TCs near the coast and the forecasting lead time is longer than 48 h for TCs over the ocean. It is reported that vertical wind shear is one of the most critical dynamic parameters influencing TCs intensity change (DeMaria and Kaplan, 1999; DeMaria et al., 2005; Zeng et al., 2010; Wang et al., 2015). In the GBRT model, vertical wind shear variables of LSHRS and LSHRD play important roles in affecting the TCs intensity when TCs are over the open ocean; however, vertical wind shear is not important for TCs intensity change when TCs are near the coast.



It can be seen from **Figures 2B**, **3B**, SL ratio is a crucial variable affecting TCs intensity change when TCs are near the coast. Different from the MLR model used by Li et al. (2018), SL ratio is not an important predictor for the future 12-h TCs' intensity change after TCs landfall. For TCs after landfall, the significant predictors are Vmax and Dvmax, and the importance of Vmax reaches 0.65–0.85 (**Figures 2A**, **3A**).

# **3.2 Evaluation of the Prediction Model for TCs Intensity Change**

As described in **Section 2**, we divide the 2000–2019 data into four groups, i.e., 2000–2004, 2005–2009, 2010–2014, 2015–2019. We use any three groups as the training datasets, and the remaining one as the test datasets to build the prediction model and perform cross validation. We evaluate the GBRT model's performance in predicting the TCs future intensity changes for TCs over land, near the coast, and over the open ocean, as well as all TC samples in terms of  $R^2$ , MAE, and RMSE. **Figures 4**, **5** show the evaluation results of the prediction models' performance with the four cross validations by boxplots based on STI data and JTWC data, respectively.

It can be seen from Figures 4, 5 that the  $R^2$  value will generally increase with the forecasting lead time. For the GBRT model based on all TC samples, the median and mean of  $R^2$ values are 0.58 (0.56) and 0.57 (0.56) by the 12-h lead time forecast and are 0.68 (0.65) and 0.67 (0.64) by the 72-h lead time forecast based on STI (JTWC) datasets. MAE and RMSE also increase with the forecasting lead time. For all sample models of STI (JTWC) data, the MAE median and mean values are approximately 2.39 (3.08) and 2.40 (3.06)  $ms^{-1}$  by the 12-h lead time forecast and are 7.79 (10.35) and 7.93 (10.29) ms<sup>-1</sup> by the 72-h lead time forecast. The RMSE median and mean values are 3.23 (4.19) and 3.24 (4.20) ms<sup>-1</sup> by the 12-h lead time forecast and are approximately 9.75 (12.99) and 9.98  $(12.94) \text{ ms}^{-1}$  by the 72-h lead time forecast. From the comparison of the statistical properties, the quality of the best track data from STI might be slightly better than the one from JTWC.

Specifically, for the GBRT model for TCs over land, the median and mean of  $R^2$  are 0.65 (0.61) and 0.64 (0.62) by the 12-h lead time forecast, and are 0.75 (0.74) and 0.76 (0.76) by the 24-h lead time forecast based on STI (JTWC) datasets. The median and mean of MAE are 1.98 (2.81) and 1.96 (2.88) ms<sup>-1</sup>



by the 12-h lead time forecast, and are 2.4 (3.82) and 2.26 (3.84) ms<sup>-1</sup> by the 24-h lead time forecast based on STI (JTWC) datasets. The median and mean of RMSE are 2.93 (3.83) and 2.77 (3.89) ms<sup>-1</sup> by the 12-h lead time forecast, and are 3.42 (4.74) and 3.08 (4.76)  $ms^{-1}$  by the 24-h lead time forecast based on STI (JTWC) datasets. Again, from the comparison of the statistical properties, the quality of the STI best track data is slightly better than the JTWC best track data. The MAE and RMSE for TC samples over land are less than those values for TCs near the coast and TCs over the open ocean. The reason may be that TCs intensity will usually decay quickly after landfall, and the number of essential variables that affect the TCs intensity change after landfall is less than the number of essential variables when TCs are near the coast and over the open ocean. Therefore, the TCs intensity changing process after landfall is not as complicated as the processes for TCs near the coast and over the open ocean. And consequently, the performance of TCs intensity forecast after landfall is generally better.

# **3.3 Performance Comparison Between the GBRT Model With Default Hyperparameters and Bayesian Optimized GBRT Model**

As described in **Section 2**, We use the Bayesian optimization algorithm to tune the hyperparameters of the GBRT model to improve the accuracy of the prediction model. The prediction results of the GBRT model with default hyperparameters are compared with the performance of the GBRT model with Bayesian optimization. The historical data from 2000 to 2014 are used for model training, and the data from 2015 to 2019 are used for model testing. The performance of the two prediction models is evaluated by the statistical properties of  $R^2$ , MAE, and RMSE. **Figures 6**, 7 show the performance comparison based on STI and JTWC datase-ts, respectively.

In the figures, it can be seen that the performance of the Bayesian optimized GBRT model for the TCs over land, near the coast, over the open ocean, and for all TC samples is generally better than the performance of the model with default hyperparameters. For STI data and JTWC data, after Bayesian optimization for the prediction model,  $R^2$  will increase by an average of 3.37% and 3.26%. MAE (RMSE) will decrease by an average of 3.22% (2.62%) and 3.82% (3.71%).

# 3.4 Performance Comparison Between the GBRT Model and MLR Model

The performance of the GBRT model is compared with that of the MLR model (Li et al., 2018) during the testing period to forecast TCs intensity over the entire TCs life span in the Western North Pacific. Li et al. (2018) used historical data from 2000 to 2011 for model calibration and data from 2012 to 2015 for model validation. To ensure a fair comparison, the same training datasets and testing datasets are selected in the GBRT model. Similarly, the performance of the different prediction models is evaluated in terms of  $R^2$ , MAE, and RMSE.

**Figures 8**, **9** show the performance comparison results based on STI and JTWC datasets, respectively. It can be seen from the two figures that the GBRT model outperforms the MLR model in TCs intensity forecast for both datasets. Compared with the performance of the MLR model,  $R^2$  of the GBRT model increases by an average of 8.47% and 4.45% for STI data and JTWC data, respectively. MAE (RMSE) drops by an average of 26.24% (25.14%) and 10.51% (4.68%) for the two datasets, respectively.

# 3.5 Performance of the GBRT Model for Intensity Prediction of Real TCs Cases

In order to further evaluate the performance of the GBRT prediction model, we apply the model to predict the intensity variation for two TC cases in the Western North Pacific, Soudelor (201513) and Mangkhut (201822), which have long life spans and have impacted China significantly. We use historical data from 2000 to 2014 to train the GBRT model. The trained GBRT model is then used to predict the intensity changes for Soudelor and Mangkhut, respectively. The TCs intensity prediction results are compared with the corresponding values from the best track data. The time series of the TCs intensity of Soudelor and Mangkhut from the two best tack data and the GBRT prediction by the 12-h and 24-h lead time forecast are compared in **Figure 10. Figure 10A** shows the tracks of these two TCs, which have experienced all three processes: over the open ocean, near the coast, and over land. The different colors in **Figure 10A** represent the different TCs intensity levels.

**Figures 10B,C** show the comparison between the time series of Soudelor's intensity prediction by the GBRT model with 12-h and 24-h forecasting lead time, and the corresponding TCs intensity values from the STI and JTWC best track data. The solid lines refer to the TCs intensity values from the best track data, and the dashed lines refer to the intensity forecast by the GBRT model. The red color indicates the comparison based on the STI data, and the black color indicates the comparison based on the JTWC data. **Figures 10D,E** show similar comparisons as **Figures 10B,C**, but for TCs Mangkhut. From the figures, it can be

seen that the prediction model can generally capture the TCs intensity variation by 12-h and 24-h lead time forecast. However, when the TCs experience rapid intensification, the models' performance is not satisfactory.

# **4 CONCLUSION**

Based on the TCs best track datasets from STI and JTWC from 2000 to 2019, this study proposes TCs intensity prediction models by the GBRT method. Using the index of SL ratio, TCs data are grouped into four parts: TC samples over the open ocean, near the coast, over land, and all TC samples. The future TCs intensity changes of 12 h, 24 h, 36 h, 48 h, 60 h, and 72 h are the model predictands, which are obtained from the best track datasets. Synoptic variables, climatological and persistent variables, which are derived from the NCEP reanalysis data and the SST data from NOAA, are used for the model predictors. Unlike the MLR model, the GBRT model can describe well the nonlinear relationship between predictors and predictands. Compared with the MLR model,  $R^2$ of the GBRT model for TCs intensity forecast increases by an average of 8.47% and 4.45% for STI data and JTWC data, respectively. MAE (RMSE) drops by an average of 26.24% (25.14%) and 10.51% (4.68%) for the two datasets. By comparing the statistical properties for the prediction models based on different datasets, the quality of the TCs best track data over the Western North Pacific from STI might be slightly better than that from JTWC.

According to the feature importance of the GBRT model, for TCs after landfall, the significant predictors are Vmax and Dvmax. Dvmax plays the most crucial role in TCs intensity change for TCs over the open ocean and near the coast, as well as for all TC samples when the forecasting lead time is 12 h. And POT is essential for TCs intensity change when the forecasting lead time is more than 24 h. Besides, SL ratio is an important variable affecting TCs intensity change when the forecast and for all TC samples when the forecas

In addition, we use the Bayesian optimization algorithm to tune the hyperparameters of the GBRT model to improve the accuracy of the prediction model. For the prediction performance based on the STI data and JTWC data, the  $R^2$  value increases by approximately 3.37% and 3.26% on average. The MAE (RMSE) decreases by an average of 3.22% (2.62%) and 3.82% (3.71%).

Overall, the GBRT model can describe well the complex nonlinear relationship between predictors and TCs intensity change and improves the performance of the intensity prediction of the TCs throughout their life span, compared with the MLR model. Therefore, the GBRT prediction model is practically valuable, and can be referred to for operational TCs intensity forecast.

# DATA AVAILABILITY STATEMENT

The original contributions presented in the study are included in the article/supplementary material, further inquiries can be directed to the corresponding author.

#### **AUTHOR CONTRIBUTIONS**

Conceptualization: QL and GZ; Data collection and analysis: GZ and QL; Methodology: QL, GZ, and WZ; Writing original draft: GZ and QL. All the co-authors participated the manuscript review and editing.

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