

In Situ Stress Distribution in Cretaceous Ultra-Deep Gas Field From 1D Mechanical Earth Model and 3D Heterogeneous Geomechanical Model, Kuqa Depression, Tarim Basin, NW China

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The Kuqa Depression boasts rich cretaceous ultra-deep hydrocarbon resources. However, it is in complex geological conditions. At present, sufficient understandings on the in situ stress distribution and influencing factors are lacking, which restricts the process of hydrocarbon exploitation. Therefore, in this study, the Bozi gas field is selected as an example, and a 1D mechanical earth model (1D MEM) is established with the drilling data and logging data through the geomechanical method to clarify the in situ stress distribution of the wellbore. A 3D heterogeneous geomechanical model (3D HGM) is established with the constraint of 1D HEM to clarify the distribution characteristics of the 3D in situ stress field in the Bozi gas field and discuss its influencing factors. The results show that: 1) the Bozi gas field is in an extremely strong in situ stress condition with high stress values. The minimum horizontal principal stress (S_h) of the cretaceous system is 153~180 MPa, and the maximum horizontal principal stress (S_H) is nearly 200 MPa; 2) the in situ stress in the Bozi gas field has obvious vertical stratification characteristics, which can be divided into three stress sequences of "low-high-low", with great differences in interlayer stress; 3) the *in situ* stress distribution of the Bozi gas field is greatly affected by the types of faulted anticline. Different types indicate different stress distribution; 4) within the influence range of overthrusts, the in situ stress in the footwall is lower than that of the hanging wall. The greater the fault offset, the greater the *in situ* stress difference between the hanging wall and footwall. Moreover, the lower the stress in the footwall, the higher is the degree of overthrust, and the larger is the range of footwall stress area; and 5) the means of highly deviated wells is more helpful to the Bozi gas field for hydrocarbon exploitation.

Keywords: in situ stress, geomechanics, numerical simulation, ultra-deep reservoir, Kuqa depression, Tarim basin

1 INTRODUCTION

At present, with the rapid development of the world economy, shallow resources are gradually exhausted, and the resource development is moving toward deep and ultra-deep layers, which have great exploration potential. The deep and ultra-deep hydrocarbon resources are an important replacement field and have become a research hotspot in the global oil and gas field (Bloch et al., 2002; Dutton and Loucks, 2010; Jia and Pang, 2015; Li et al., 2019). In recent years, important progress has been made in global deep and ultra-deep hydrocarbon exploitation. By the end of 2018, 68 oil and gas reservoirs with a depth of more than 8,000 m had been found worldwide (Li et al., 2020), with the new oil and gas reserves and production showing an obvious growth momentum (Bloch et al., 2002; Hu et al., 2013; Bai and Cao, 2014; Cao et al., 2022).

As for the boundary depth of deep and ultra-deep reservoirs in petroliferous basins, different countries, institutions, and scholars have different understandings (Glasmann, 1992; Girard et al., 2002; Hu et al., 2013; Pang et al., 2015; Hu et al., 2019). Cao et al. (2022) proposed a division by considering burial depth, geothermal gradient, burial history, thermal evolution history, and diagenetic history, which is found in cold basins with low geothermal gradient (<2.5°C/100 m); 4,500 ~ 6,000 m is defined as deep layer and more than 6,000 m is defined as ultra-deep layer; in medium warm basins with a geothermal gradient of 2.5 ~ 3.5°C/100 m, 3,500 ~ 4,500 m is defined as the deep layer; and more than 4500 m is defined as the ultra-deep layer; and in the thermal basin with a high geothermal gradient (>3.5°C/100 m), 3,000 ~ 4,000 m is defined as the deep layer, and more than 4,000 m is defined as the ultra-deep layer.

At present, the deepest large-scale clastic rock oil and gas reservoir found abroad is the Jack/St Malo deep-water turbidite oilfield of Lower Tertiary Trend in the Gulf of Mexico, with a maximum burial depth of 8,839 m, and its oil and gas reserves are 6821×10^4 t (oil equivalent) at 7,620 m (25,000 ft) with a water depth of 2,133 m (7,000 ft) of Well Jack-2, and the production is 818 t/d (Meyer et al., 2007; Rains et al., 2007; He et al., 2017).

Currently, the deepest borehole in China is Well Luntan 1 in Tarim Basin (8,882 m in 2019), where high-yield industrial oil and gas flow is obtained in the Cambrian below a burial depth of 8,200 m (Yang et al., 2020). The deepest clastic rock gas field is the Kelasu gas field in Tarim Basin (Yang et al., 2018, 2019; Zeng LB. et al., 2020; Tian et al., 2020). At Well Dabei 4, high-yield natural gas flow is obtained in the year of 2022 at 8,200 m in the cretaceous sandstones.

Theoretically, in such ultra-deep background, the reservoir rocks become extremely dense due to strong diagenetic compaction and cementation. The matrix porosity is generally not more than 10%, and the permeability is even less than 1mD, making it difficult to form an efficient seepage system. However, such ultra-deep reservoirs often develop fractures and are present in strong *in situ* stress environment. The exploration and development practice of many ultra-deep fractured reservoirs shows that the high-quality fractures under the control of *in situ* stress are the optimum point of ultra-deep hydrocarbon exploitation (Zeng et al., 2008; Jiang et al., 2020). Under the control of *in situ* stress field, fractures can increase the reservoir permeability by orders of magnitude (Wang et al., 2018; Zeng QL. et al., 2020; Liu J. et al., 2022), which is an important guarantee for the existence of wells with high and stable production in ultradeep reservoirs. *In situ* stress greatly influences the quality of ultra-deep reservoir and hydrocarbon production capacity by affecting the fracture activity and hydraulic fracturing effect (Zhang et al., 2018; Xu et al., 2022). Therefore, it is necessary to carry out systematic research on the *in situ* stress in ultra-deep reservoirs, clarifying the orientation and magnitude of *in situ* stress, and finding the main factors affecting its distribution. The results are expected to provide suggestions and supports for well location deployment and optimization fracturing scheme.

This study takes the Bozi gas field in Tarim Basin as an example to study the in situ stress. The Bozi gas field is a typical ultra-deep gas field with a burial depth of 6,000 ~ 8,000 m. At present, the existing research of the Bozi gas field mainly focuses on sedimentation, structure, and petroleum geological conditions (Yang et al., 2019; Wang et al., 2022) and understandings on the in situ stress distribution characteristics and influencing factors are lacking, which restricts the hydrocarbon exploitation in this area. Therefore, a one-dimensional mechanical earth model (1D three-dimensional MEM) and а heterogeneous geomechanical model (3D HGM) were conducted and It plays an important role in better analyzed. understanding the geomechanical properties and provides references for the efficient development of natural gas.

2 GEOLOGICAL SETTING

2.1 Location

The Kuqa Depression is located in the north of Tarim Basin. It is connected with the South Tianshan fault fold belt by the thrust fault in the north, Tabei Uplift in the south, Yangxia sag in the east, and Wushi sag in the West. It is a superimposed foreland basin dominated by Mesozoic and Cenozoic sedimentation and can be divided into seven secondary structural units (Figure 1A). Among them, the Kelasu tectonic belt is a typical deep and ultra-deep natural gas resource enrichment area. From west to east, it can be divided into Awate segment, Bozi segment, Dabei segment, and Keshen-Kela segment. The Kelasu tectonic belt develops four first-order faults from north to south, which are divided into Bozi-Kela fault tectonic belt, Keshen fault tectonic belt, Baicheng fault tectonic belt, and Baicheng South fault tectonic belt. Multiple secondary thrust faults are developed between the faults.

The Bozi gas field is located in the west of Kelasu tectonic belt and is characterized by rows of faulted anticlines. The strike of anticline is generally the same, which is in the NE–SW direction. The northern Bozi gas field is a pop-up structure sandwiched by two thrust faults F1 and F2, and the southern part is a series of imbricate structures, with gradually increasing burial depth (**Figures 1B,C**). From east to west, the overlapping degree of the imbricate structure increases gradually.

A104

A105

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Neogene Kuqa Formation (N_2k) , Kangcun Formation $(N_{1-2}k)$, Jidike Formation (N1j), Paleogene Suweiyi Formation (E2-3s), Kumugeliemu Group (E₁₋₂km), Cretaceous Bashijiqike Formation (K_1bs) , and Baxigai Formation (K_1bx) . Among them, The $E_{1-2}km$ is a set of gypsolith rock formation with great thickness distribution and plastic flow characteristics. In the process of compression deformation, layered shrinkage deformation occurs, and the strata above this plastic stratum folds, thrusts, and rises sharply, forming a "roof like" structure. Under the control of "roof" structure, salt rock flow results in the characteristics of abovementioned salt Kelasu tectonic belt, salt body, and subsalt differential structural deformation. The target strata of the study are the Cretaceous K₁bs and

from the top to bottom are: Quaternary Xiyu Formation (Q_1x) ,

 K_1bx . The K_1bs can be divided into three lithologic segments. The first lithologic segment is missing due to the influence of paleosedimentary environment. The second lithologic segment mainly comprises medium thick ~ super thick layered brown and

grayish brown medium sandstone and fine sandstone, which is deposited at the front of braided river delta. The third lithologic segment mainly comprises medium thick ~ thick layered brown fine sandstone, thin ~ medium thick layered siltstone, and argillaceous siltstone, which are deposited at the front of fan delta. The K_1bx comprises thin medium thick brown mudstone and medium thick layered brownish gray, grayish brown fine sandstone, siltstone, and argillaceous siltstone, which are interbedded in an isopachous-slight isopachous manner and are deposited at the front of braided river delta (Figure 2).

2.3 Reservoir

Reservoir pore types in the Bozi gas field are mainly primary intergranular pores, with a small amount of microfractures and micropores. The porosity of Bashijiqike formation mainly ranges between 4% and 10.0%, with an average of 6.30%, and the permeability mainly ranges between 0.1 and 0.5 mD, with an average of 0.256 mD. The porosity of Baxigai formation sandstone ranges between 4 and 9%, with an average porosity of 6.06%, and the permeability mainly ranges between 0.035 and

2.2 Stratigraphy The drilling of the Bozi gas field shows that the drilling strata

D 9 FIGURE 1 | Field location, seismic profile, and top structure of the target layer in the Bozi gas field. (A) The Kuga Depression is the northernmost first level unit in Tarim Basin. The Kuga depression can be further divided into seven secondary tectonic units. The Kelasu tectonic belt is located in the second-row structure from the north, and the Bozi gas field is located in the west of the Kelasu tectonic belt; (B) the Cretaceous top structure map of the Bozi gas field comprises several rows of faulted anticline structures divided by thrust faults. The strike of the anticline structure is basically the same, which is in the NE-SW direction; and (C) the typical thrust imbricate structure and pop-up structure of the Bozi gas field can be seen through the seismic section of PP'.

Α

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С

A901

0.1 mD, with an average permeability of 0.067 mD. The reservoir fractures are relatively developed, mostly being partially filled—unfilled, high angle, and shear fractures. The fracture strike is mainly near E–W, and the fracture density ranges between 0.2/m and 0.4/m.

3 METHODS

3.1 *In Situ* Stress State Interpreted From Wells

At present, there are many methods to measure and test *in situ* stress, including the drill-based borehole collapse method, relaxation strain measurement, stress relief technique, acoustic emission, and differential strain method (Zang and Stephansson, 2010). The Bozi gas field is deep buried and has high bottom hole temperature and high pressure. Hence, the measuring instrument has poor applicability, and there is no constraint of confining pressure after core removal, causing rapid stress relief. In mild cases, it may lead to inaccurate measurement, and in serious cases, it may lead to possible core fracture. Therefore, it is not suitable to use downhole measurement and core testing to obtain the *in situ* stress state.

This article utilizes the drilling and logging data to obtain the *in situ* stress state. Different basins are in different states, and scholars have put forward a variety of *in situ* stress calculation models considering more comprehensive factors. As a result, the accuracy of the calculation has gradually improved, and the applicability has become more extensive. At present, there is little dispute about the calculation of vertical stress (S_v). It is considered that its magnitude is equal to the gravity of the overlying rock mass, which can be obtained by integrating the density curve from the ground to the target depth such as **Eq. 1**:

$$S_V = \int_0^z \rho(z) g dz, \qquad (1)$$

where g is the gravitational acceleration (m/s^2) , $\rho(z)$ is the density as a function of burial depth (kg/m^3) , and z is the burial depth (m).

There are many calculation models for obtaining horizontal principal stress magnitudes, including the uniaxial strain model, Mohr–Coulomb failure model, Coulomb–Navier failure model, Huang's model, combined spring model, porous elastic strain model, and biaxial strain model (Li and Zhang, 1997; Liu and Luo, 1999; Rasouli and Sutherland, 2014; Li et al., 2017). In this article, the calculation model is as follows (Li et al., 1997):

$$\begin{cases} S_{H} = \frac{\mu}{1-\mu} \left(S_{V} - \alpha P_{p} \right) + \frac{E\xi_{H}}{1-\mu^{2}} + \frac{\mu E\xi_{h}}{1-\mu^{2}} + \alpha P_{p} \\ S_{h} = \frac{\mu}{1-\mu} \left(S_{V} - \alpha P_{p} \right) + \frac{E\xi_{h}}{1-\mu^{2}} + \frac{\mu E\xi_{H}}{1-\mu^{2}} + \alpha P_{p} \end{cases}$$
(2)

where S_H is the maximum horizontal principal stress, MPa; S_h is the minimum horizontal principal stress, MPa; S_V is the vertical principal stress, MPa; P_P is the pore pressure, MPa; μ is Poisson's ratio, dimensionless; E is the modulus of elasticity, GPa; α is the Biot coefficient, dimensionless; and $\xi_H \xi_h$ are the maximum and minimum principal stress coefficient, respectively, dimensionless.

The method of calculating the mechanical parameters with the logging data is as follow (Lu et al., 2015):

$$E_{\rm d} = \frac{\rho_{\rm b}}{\Delta t_{\rm s}^2} \cdot \frac{3\Delta t_{\rm s}^2 - 4\Delta t_{\rm p}^2}{\Delta t_{\rm s}^2 - \Delta t_{\rm p}^2},\tag{3}$$

$$\mu_{\rm d} = \frac{\Delta t_{\rm s}^2 - 2\Delta t_{\rm p}^2}{2\left(\Delta t_{\rm s}^2 - \Delta t_{\rm p}^2\right)},\tag{4}$$

where ρ_b is the rock density, kg/m³, and Δt_p and Δt_s are the P-wave time difference and S-wave time difference, respectively, μ s/ft.

Generally, we may determine S_h at a specific location through the hydraulic fracturing construction data, which can be used as the constraint and scale basis to determine the value of ξ_h . However, the maximum horizontal principal stress (S_H) of ultra-deep wells cannot be determined through hydraulic fracturing (Hickman, 1983; Zoback, 2007). In this article, the S_H is determined with the wellbore fracture information. In the process of drilling, as the borehole is drilled, the borehole wall may produce stress concentration under the action of confining pressure. When the stress concentration exceeds the fracture strength of the rock around the borehole, the borehole wall will collapse (Bell and Gough, 1979; Zoback et al., 2003).

Because wells in different structural locations have different stress distributions (stress concentration and tensile range) around the well in different lithologic sections, and with the increase of radial depth, the stress effect will decrease and the degree of borehole wall collapse will also vary. Therefore, it is a more intuitive method to distinguish the direction of *in situ* stress by analyzing the changes of formation characteristics around the borehole from the scanning images of formation micro resistivity (FMI). The collapse width of the borehole wall can be judged from FMI, and the collapse width has the following mathematical calculation relationship with rock uniaxial compressive strength and *in situ* stress state (Zoback, 2007):

$$S_{H} = \frac{(C_{0} + 2Pp + \Delta p + \sigma^{\Delta T}) - Sh \cdot [1 + 2\cos(\pi - W_{BO})]}{1 - 2\cos(\pi - W_{BO})}, \quad (5)$$

where C_0 is the uniaxial compressive strength, P_p is the pore pressure, Δp is the difference between pore pressure and bottom hole pressure, W_{BO} is the collapse width, and $\sigma^{\Delta T}$ is the thermal stress, which can be ignored generally.

Therefore, the gradient of the maximum horizontal principal stress and the minimum horizontal principal stress can be inverted according to the collapse width W_{BO} and the uniaxial compressive strength (C_0) of the borehole wall collapse position.

3.2 3D Geomechanical Modeling of *In Situ* Stress

3.2.1 Principles of the Finite Element Method

The distribution of *in situ* stress field was predicted using a popular finite element method (Liu et al., 2017a; Liu et al., 2017b; Ju et al., 2021). To begin with, the geological mass was discretized

into several finite units that were connected to each other by nodes. Rock mechanical parameters were assigned to the corresponding units. The primary variables of the field functions for the study area included displacement, stress, and strain. The equation set where the node displacement served as the unknown quantity and the total stiffness matrix as the coefficient was solved based on the boundary stress condition and the node balance condition. The displacement at each node was then obtained. Finally, the stress and strain of each unit were calculated.

Generally, based on elastic mechanics, the relationship between strain and displacement follows Eq. 6 (Ju et al., 2021):

$$\begin{cases} \varepsilon_{xx} = \frac{\partial i}{\partial x}, \ \gamma_{xy} = \gamma_{yx} = \frac{\partial i}{\partial y} + \frac{\partial j}{\partial x} \\ \varepsilon_{yy} = \frac{\partial j}{\partial y}, \ \gamma_{yz} = \gamma_{zy} = \frac{\partial j}{\partial z} + \frac{\partial k}{\partial y}, \\ \varepsilon_{zz} = \frac{\partial k}{\partial z}, \ \gamma_{zx} = \gamma_{xz} = \frac{\partial k}{\partial x} + \frac{\partial i}{\partial z} \end{cases}$$
(6)

where ε_{xx} , ε_{yy} , and ε_{zz} are the linear strain components and γ_{xy} , γ_{yz} , and γ_{zx} are the shear strain components.

The matrix form can be derived as follows:

$$[\varepsilon] = [B][\delta], \qquad (7)$$

where $[\delta]$ is the nodal displacement matrix and [B] is the geometric matrix.

From the physical equation, stress and strain can be expressed as follows:

$$[\sigma] = [D][\varepsilon], \qquad (8)$$

where [D] is the elasticity matrix.

3.2.2 Geological Model and 3D Heterogenous Rock Mechanical Parameters

The Bozi gas field is featured by fault developed and complex fault intersection. This article adopts the volume element based (VBM) structural framework modeling method. The grid established is continuously distributed, but it is discontinuous at the fault. Compared with the traditional corner grid modeling, the established complex structure model is more realistic and has higher accuracy. In order to consider the heterogeneity of rock, the "logging + seismic" combination method is used to establish the 3D heterogeneous rock mechanics field (Xu et al., 2018a; Xu et al., 2018b; Ju et al., 2021).

3.2.3 Boundary Conditions

The surface of the Bozi gas field is relatively flat and mostly covered by vegetation and Gobi. However, due to the complex lithology, the shallow Kuqa and Kangcun Formations gravel layer and the plastic Kumugeliemu group gypsolith rock have a certain impact on the distribution of *in situ* stress in the target layer. Therefore, when establishing the model, this article considers the surface fluctuation morphology and different rock mechanical properties of the surrounding rock, rather than just treat the rock body above the target layer as a homogeneous cube. As shown in Figure 3, an outer frame is established outside the Bozi gas field model as the surrounding rock. The surrounding rock is divided into five layers, representing loose rock on the surface, shallow gravel, mudstone layer, gypsum salt rock, and Cretaceous rock layer, which are given by the rock mechanical parameters of different lithology, respectively. The rock mechanical parameters of each layer are obtained according to the mechanical test of field samples in Kuqa Depression and corrected according to the vertical stress calculated by density logging of the target layer (Xu et al., 2020). The target layer model is assigned with 3D heterogeneous rock mechanical parameters, among which the elastic modulus mostly ranges between 20 and 30 GPa, locally as high as 40 GPa; the Poisson's ratio is mainly concentrated in 0.24 \sim 0.27; and the value of rock density in 3D space is mainly between 2.3 and 2.6 g/cm3 (Figure 3). Using the distance scanning method (that is, calculate the distance between a unit and its surrounding rock mechanical parameters one by one and select the rock mechanical parameters with the smallest distance to load on the unit), this article assigns the heterogeneous rock mechanical parameters in the constructed 3D rock mechanical field to each finite element grid of the target layer model.

With the *in situ* stress state of seven wells in the Bozi gas field as the constraints, the iterative optimization process is carried out for the stress boundary loading to optimize and determine the reasonable boundary conditions. Through repeated attempts, the following boundary conditions are finally determined (**Figure 3**): a compression magnitude of 210 MPa is applied in the N–S direction and 163 MPa is applied in the E–W direction of the model.

4 RESULTS

4.1 1D Mechanical Earth Model

In this article, a 1D mechanical earth modeling has been carried out in the Bozi gas field. **Figure 4** shows the 1D MEM result along the well section of A2–A101-2–A102-4–A105, which indicates the distribution and changes of Cretaceous geomechanical parameters in one dimension of the wellbore.

Taking Well A2 as an example, Young's modulus is 20 ~ 30 GPa, the average uniaxial compressive strength is approximately 100 MPa, and Poisson's ratio changes little, at 0.25 ~ 0.30. The minimum horizontal, maximum horizontal, and vertical principal stress gradients are 2.11 ~ 2.18 MPa/100 m, 2.55 ~ 2.72 MPa/100 m, and 2.45 MPa/100 m, respectively, which are in strike–slip stress regime (S_H > S_V > S_h). For a better comparison, the S_h is filled with colors, with warm color representing low value and cold color for high value. It can be seen that the *in situ* stress shows obvious stratification vertically. The stress in the upper part of K_1bs_2 and K_1bs_3 , that is, 7456 ~ 7555 m is low, and the S_h magnitude is approximately 165 MPa. The stress in the lower part of K_1bs_3 , the upper part of K_1bx_1 and K_1bx_2 , that is, 7555 ~ 7700 m is high, and the S_h magnitude is approximately 175 MPa, and local areas can reach 180 MPa. The stress in the lower part of

Formation	Description	Formation	Depth/m	Lithology
Q ₁ x	Medium thick-thick layered fine conglomerate and glutenite, locally intercalated with thin layered mudstone and siltstone	K.I.	7200	••••
N ₂ k	Medium thick-thick layered fine conglomerate, locally intercalated with thin-medium thick layered mudstone and argillaceous siltstone	K ₁ <i>bs</i> ₂	7220 7230 7240	· · · · ·
$N_{1-2}k$	The middle and upper part is dominated by extremely thick layered sandy conglomerate, intercalated with medium thick layered siltstone and mudstone; The bottom is dominated by medium thick layered sandy conglomerate, intercalated with thin-medium thick layered siltstone and mudstone	K, <i>bs</i> ,	7250	
N <i>.j</i>	The upper part is mainly medium thick-thick layered pebbly fine sandstone, and the middle part is brown mudstone mixed with thin layer pebbly fine sandstone and siltstone; The lower part is dominated by mudstone. From the upper part to the lower part, the particle size gradually becomes finer and the mudstone gradually develops		7270	
E ₂₋₃ s	The upper part is mainly medium thick layered mudstone, the lower part is mainly medium thick - thick layered mudstone, intercalated with thin- medium thick layered gypsum mudstone, and the bottom is intercalated with argillaceous siltstone		7290 7300 7310	· · · · · · · · · · · · · · · · · · ·
$\mathbf{E}_{1-2}km$	Medium thick layered mudstone, gypsum containing mudstone and gypsum mudstone are mainly mixed with thin layers of argillaceous gypsum rock, gypsum siltstone and argillaceous siltstone in the middle, and thin to medium thick layered gypsum mudstone, argillaceous gypsum rock and gypsum fine sandstone		7320	
K,bs	The first setion is missing; The second section: medium thick ~ super thick layered fine sandstone, argillaceous fine sandstone, medium sandstone and gravelly sandstone, locally intercalated with thin and medium thick layered mudstone. The third section: medium thick thick layered fine sandstone and medium sandstone, followed by siltstone, intercalated with thin ~ medium thick layered mudstone and silty mudstone. There are more mudstone Intercalations in this section than in the second lithologic section	K ₁ <i>bx</i> ₁	7340 7350 7360 7370 7380	
K ₁ bx	The first section: the upper thin \sim medium thick layered mudstone is interbedded with fine sandstone and siltstone, and the lower lithology is mainly thin \sim medium thick layered mudstone and silty mudstone, mixed with fine sandstone. The second section: fine sandstone, fine sandstone with mud gravel, mudstone and silty mudstone are interbedded with unequal thickness.	K ₁ bx ₂	7390 7400 7410 7420	
K ₁ s	Mainly medium thick layered mudstone		7430	

 K_1bx_2 decreases, and the S_h magnitude is approximately 170 MPa. This shows the distribution characteristics of "low-high-low".

The *in situ* stress of single wells in the Bozi gas field has certain similar characteristics. The stress has the distribution

characteristics of "low-high-low" in the Cretaceous system. That is, the stress in the upper part of K_1bs_2 and K_1bs_3 is low, the stress in the lower part of K_1bs_3 , and the upper part of K_1bx_1 and K_1bx_2 is high, and the stress in the lower part of K_1bx_2 is



reduced. In addition, compared with the Baxigai Formation, the stress characteristics of Bashijiqike Formation are lower, and stress difference between the layers is relatively small. The interlaminar stress difference in the Baxigai Formation is large, resulting in a stronger interlaminar property.

4.2 3D Stress Distribution

The 3D distribution of the *in situ* stress in the Cretaceous layes of Bozi gas field is shown in **Figure 5**. The distributions of three principal stresses are similar, showing a trend of low in the north and high in the south. The S_h is mainly 153 ~ 180 MPa, and S_H is 180 ~ 200 Pa, and S_V is 175 ~ 200 MPa, which is a strike slip stress state. There are great differences in stress distribution among the different fault blocks. The S_h in A104 and A102 fault blocks is 153 ~ 160 MPa and the S_H is approximately 180 ~ 188 MPa, while the S_h in A8 and A9 fault blocks is as high as 175 MPa and the S_H even exceeds 200 MPa.

Figure 6 shows the *in situ* stress distribution of five typical sections in the Bozi gas field. It can be seen that the *in situ* stress distribution is evidently related to the tectonic form. Each fault block in the Bozi gas field is basically a fault anticline structure, and different structural styles vary in terms of stress distribution. **Figure 7** summarizes the structural style of the fault anticline in the Bozi gas field and the corresponding *in situ* stress distribution mode. The structural style of the fault anticline is divided into five

categories: north steep and south gentle type (A8 and A10), north gentle and south steep type (A7), symmetrical imbricate type (A104), complete pop-up structure (A7), and semi complete pop-up structure (A1). Generally, the *in situ* stress distribution of symmetrical imbricate faulted anticline and pop-up structure is regular and symmetrical. The stress is low at the top of the anticline and high at the wings. In north steep and south gentle type, the stress is high in the north wing and low in the south wing. In the north gentle and south steep type, the stress in low in the north wing and high in the south wing.

5 DISCUSSION

5.1 Error Analysis

The *in situ* stress prediction results of seven wells in different fault blocks of the Bozi gas field are compared with the logging calculation results (**Table 1**). The calculation method of error (r) is as follows:

$$r = S_c - S_m, \tag{9}$$

where S_m is the *in situ* stress value actually calculated, MPa and S_c is the *in situ* stress predicted by numerical simulation, MPa. It can



be seen that the average error of the S_H value is 6.5 MPa, and the average error of the S_h value is 6.3 MPa. Both the errors are within 10%, indicating that the *in situ* stress simulation results have high reliability.

5.2 Influence of Thrust Imbricate Structure on *in situ* Stress

The *in situ* stress distribution of the thrust imbricate structure is complex, especially the disturbance effect of the fault on the *in situ* stress (Meyer et al., 2007; Rasouli and Sutherland, 2014). In this study, the *in situ* stress shows a special distribution feature in the thrust imbricate structure (**Figure 6**). In the area around a fault, the *in situ* stress in the footwall is lower than that of the hanging wall, even if the footwall is buried deeper. Moreover, the greater the fault offset, the greater is the stress difference between the hanging wall and footwall, that is, the *in situ* stress value within the control range of the footwall is lower. According to the statistics of stress values from the hanging wall and the footwall of 20 faults (**Figure 8**), when the fault offset is less than 50 m, the stress difference is less than 5 MPa, and while it exceeds 500 m, the stress difference exceeds 15MPa, and there is a good

correlation between the fault distance and stress difference between the hanging wall and footwall. In addition, the footwall of the overthrust generally shows low stress value. The higher the degree of overthrust, the larger is the range of low stress area of the footwall.

5.3 Support for Well Location

Previous studies (Liu et al., 2004; Li et al., 2012; Liu JS. et al., 2022) indicate that the wellbore trajectory is the most stable along the direction of the S_H under the strike–slip stress regime. The practice of Tarim Oilfield also confirms that under the ultra-deep strike–slip stress regime, highly deviated wells have fewer complex drilling accidents and shorter drilling cycle compared with the vertical wells (Xu et al., 2020; Cai et al., 2022). Therefore, in the Bozi gas field, highly deviated wells or horizontal wells have better wellbore stability.

The *in situ* stress field in the Bozi gas field is heterogeneous, and the *in situ* stress distributes in the belt horizontally and stratified vertically. The directional wells (inclined or horizontal wells) are more likely to drill in favorable low-stress zones than the vertical wells, and the directional trajectory or sidetrack







TABLE 1	Analysis o	of numerical	simulation	results	of <i>in</i>	situ s	stress.
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Well	Depth (m)	S _H			S _h			
		Measured (MPa)	Predicted (MPa)	Error (MPa)	Measured (MPa)	Predicted (MPa)	Error (MPa)	
A1	7043	184.8	193.6	8.8	157.5	150.2	-7.3	
A101	6998	179.3	190.5	11.2	148.3	148.8	0.5	
A102	6832	175.5	176.8	1.3	143.2	148.7	5.5	
A104	6774	179.4	171.2	-8.2	144.1	136.7	-7.4	
A2	7500	202.1	191.4	-10.7	158.0	170.5	12.5	
A8	8125	202.2	200.0	-2.2	174.4	176.2	1.8	
A901	7800	202.8	200.0	-2.8	169.0	178.4	9.4	



drilling may be adjusted during the drilling based on *in situ* stress field.

Due to the high degree of overthrust in the Bozi gas field and the favorable low-stress area is often in the footwall, the vertical well method cannot meet the drilling conditions, while the directional well method can avoid the obstacles of nontarget layer and drill into the favorable area of target layer. The existing practice in Tarim Oilfield also proves that the production of directional wells is higher than that of the vertical wells. Therefore, it is considered that the highly deviated wells are more suitable for hydrocarbon exploitation in the Bozi gas field.

6 CONCLUSION

- The Bozi reservoir is in extremely strong *in situ* stress environment, which is featured by high *in situ* stress and horizontal minimum principal stress (S_h) of the Cretaceous reservoir ranging between 153 and 180 MPa, and horizontal maximum principal stress generally exceeds 200 MPa. It is dominant by the strike-slip stress regime.
- 2) The *in situ* stress of the Bozi gas field has obvious stratification characteristics vertically, and it can be divided into three stress sequences of "low-high-low". The upper part of K_1bs_2 and K_1bs_3 is a low-stress segment, the lower part of K_1bs_3 , K_1bx_1 , and K_1bx_2 are high-stress segments, and the lower part of K_1bx_2 is a low-stress segment, resulting in great difference in interlayer stress.
- 3) The *in situ* stress distribution of the Bozi gas field is greatly affected by the structural style of the fault anticline. The symmetrical imbricate faulted anticline and pop-up structure show regular and symmetrical *in situ* stress distribution, with low stress at the top of anticline and high stress at the wings. The stress is low at the top of anticline and high at the wings. In the north steep and south gentle type, the stress is high in the north wing and low in the south wing. In the north gentle and south steep type, the stress is low in the north wing and high in the south wing.

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- 4) The overthrust structure has an obvious influence on the *in situ* stress distribution. In the area around a fault, *in situ* stress in the footwall is lower than that of the hanging wall. The greater the fault offset, the greater is the stress difference between the hanging wall and footwall. The footwall of the overthrust generally shows low stress value. The higher the degree of overthrust, the larger is the range of low stress area of footwall.
- 5) In the Bozi gas field, under the strike-slip stress regime, the directional wells have multiple advantages such as safety and stability, crossing more favorable parts and avoiding nontarget formation obstacles. The highly deviated wells are more suitable for the Bozi gas field to carry out hydrocarbon exploitation.

DATA AVAILABILITY STATEMENT

The original contributions presented in the study are included in the article/Supplementary Material. Further inquiries can be directed to the corresponding author.

AUTHOR CONTRIBUTIONS

KX, Conceptualization, Investigation, Writing—Original Draft; HZ, Supervision, Methodology, Validation; RD, Software, Visualization; WJ, Formal Analysis; YX, Data Curation; DC, Writing—revised manuscript; LF, Formal Analysis; FY, Formal Analysis; SL, Software; JL, Software; WZ, Software; BZ, Software.

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Conflict of Interest: KX, HZ, RD, YX, DC, LF, FY, SL, JL, WZ, and BZ were employed by Tarim Oilfield Company, PetroChina.

The remaining author declares that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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