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Limited role of shifting cultivation in soil carbon and nutrients recovery in regenerating tropical secondary forests

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Shifting cultivation is a dominant land-use in the tropical forest-agriculture frontier in Southeast Asia and is blamed for much of the environmental degradation in the region. We examined the distribution and availability of four soil macronutrients-i.e., soil organic carbon (SOC), total nitrogen (N), phosphorus (P) and potassium (K), in secondary forests regenerating after shifting cultivation abandonment. Soil samples were collected along an upland fallow gradient on Leyte Island in the Philippines. The effect of site environmental attributes on the availability of SOC and nutrients was investigated using linear mixed-effect models. We found relatively higher concentrations of SOC and P in the oldest fallows and higher N concentration in the youngest fallow secondary forest. There was no significant difference in SOC and other macronutrients within sites of different fallow categories and soil depths, except in the case of soil K, which was highest in our control old-growth forest. Patch size together with slope of the site and fallow age were the most influential factors in explaining the variability in SOC and nutrients availability in secondary forests recovering after shifting cultivation abandonment. Our study suggests that shifting cultivation may not be detrimental to soil quality, at least on the soil parameters and soil type we studied in the Philippines upland.

KEYWORDS

soil macronutrients, succession, forest restoration, Philippines, Southeast Asia

Introduction

Shifting cultivation, swidden or slash-and-burn agriculture, is a common land-use in the tropical forest-agriculture Frontier and considered as a major source of forest degradation (Ziegler et al., 2011; Curtis et al., 2018). Generally, shifting cultivation involves slashing and burning of forest vegetation before, or at the onset, of the monsoon to release nutrients locked in plant biomass; the sites are then cultivated and harvested over several years before they are left fallow to allow secondary forest development (Fox et al., 2000). Much of the forests in the tropics have been subjected to shifting cultivation, and this is a major contribution to the livelihood and food security of smallholder rural farmers living in upland rural areas (Mertz et al., 2009a; van Vliet et al., 2012; Heinimann et al., 2017).

Although a reliable estimate of the extent of land under shifting cultivation is not available, according to Mertz et al. (2009b) at least 14–34 million people are dependent on shifting cultivation alone in Asia. In the last few decades, a rapid transformation of shifting cultivation landscapes, however, has taken place in much of the Southeast Asia, mainly driven by urbanization, and changing land-use policies and focus of the governments (Fox et al., 2009; van Vliet et al., 2012). Another threat is posed by large-scale plantations of commercial and perennial crops including rubber and oil palm in lands where shifting cultivation was prevalent (Ziegler et al., 2012; Ahrends et al., 2015; Terefe and Kim 2020). Yet shifting cultivation is regarded as one of the major drivers of deforestation and forest degradation in this region and hence is discouraged in local landuse policies (Hett et al., 2012; Mukul et al., 2016a).

Soil quality is one of the key environmental attributes that is largely influenced by shifting cultivation, and other human use and local practices in the humid tropics (Kotto-Same et al., 1998; Hughe et al., 1999; Bruun et al., 2009; Lawrence et al., 2010; Filho et al., 2015; Zhang et al., 2019). Soil quality is the capacity of the soil to support forest growth without causing further degradation of the soil or to the environment (Lal 1997). Soil quality is closely linked to soil resilience which refers to the ability of the soil to restore its functions following major disturbances (Guo and Gifford 2002; Filho et al., 2015). The environmental aspects of shifting cultivation, however, are more complex than is often presented in the literature, and the transformations of shifting cultivation landscapes to other land-uses may have a wide range of environmental consequences, both at local and global levels (Bruun et al., 2009; van Vliet et al., 2012; Mukul and Herbohn 2016). The transition of shifting cultivation landscapes to sedentary agriculture, for instance, will bring a substantial reduction in the above and below ground carbon stocks compared to the transition of shifting cultivation landscapes to secondary forests (Ziegler et al., 2012; van Straaten et al., 2015).

The effect of anthropogenic forest disturbances, including shifting cultivation, on soil carbon and other nutrients are still unclear, and characterized by data scarcity and inconclusiveness (see—Richards et al., 2007; de Neergaard et al., 2008; Bruun et al., 2009; Sang et al., 2013; Paudel et al., 2015; de Blecourt et al., 2017; Teixeira et al., 2020; Mukul et al., 2021). Consequently, we investigated whether shifting cultivation has major impacts on key soil macronutrients and soil quality indicators namely—soil organic carbon (SOC), total nitrogen (N), phosphorus (P) and potassium (K) in secondary forests regenerating after shifting cultivation, locally known as kaingin, represents a dominant land-use (Saurej

and Sajise 2010). After logged forest, post-kaingin forest constitutes the largest group of secondary forests in the country (Chokkalingam and Perera 2001).

The specific objectives of our study were—1) to determine the status of SOC, total N, P, and K in regenerating forests recovering after shifting cultivation abandonment in the Philippines uplands, and 2) to identify the effect of site environmental attributes on soil carbon and nutrient recovery. Understanding soil C and nutrients dynamics after shifting cultivation abandonment also have wider implications for forest management and restoration, not only in the Philippines, but in other parts of the humid tropics where large areas of land are now recovering from anthropogenic and other disturbances (see—Bruun et al., 2009; Sang et al., 2013; Smith et al., 2015; Thomaz 2017; Ota et al., 2020; Mukul et al., 2020, 2022).

Materials and methods

The study area

About 23 percent of the total land area in the Philippines is classified as uplands with elevation ranging between 100 and 500 m asl (Carating et al., 2014). Our study was conducted in an upland area situated on the western side of the island of Leyte (Figure 1). Leyte is the eighth largest island in the country with an area of about 800,000 ha and located between 124°17′ and 125°18′ East longitude and between 9°55′ and 11°48′ North latitude. Forests cover only about 10 percent area in Leyte. The once dipterocarp rich rainforests of the island are now dominated by patches of old-growth forest, secondary forest, coconut plantations, Abaca (Musa textilis, a fiber yielding species from the Musaceae) and fast-growing timber plantations (Asio et al., 1998; Bonner et al., 2019).

According to the Coronas Climate Classification, Leyte has a "type IV" climate with two distinct seasons—monsoon and dry. The annual rainfall is about 4,000 mm with a mean annual temperature of about 28°C that varies little throughout the year (Mukul et al., 2016a). The relative humidity of the island ranges from 75 to 80 percent (Kolb 2003). The soil on Leyte island is mostly Andisol which possesses markedly higher concentrations of organic carbon than other parts of the country (Navarrete et al., 2013).

Site selection and characteristics

We purposively selected Barangay (administrative entity, similar to a village) Gaas (hereafter Gaas only) on Leyte Island because the area has a relatively low population density and forests are relatively undisturbed there, which is one of the prerequisites for the development of second-growth forests (Chazdon 2014). Kaingin in the Philippines can be categorized into three distinct types based on the sites where it is practiced: the tubigan, the katihan and the



dahilig systems (Olofson 1980). The tubigan and katihan systems are practiced in lower elevation areas or on gently sloping land with limited irrigation facilities. In contrast, the dahilig system is widely practiced in heavily forested areas and on steeper slopes (Olofson 1980). For our study, we focused, only on the dahilig system because it represents the shifting cultivation system widespread in most parts of Southeast Asia (Chokkalingam and Perera 2001).

Altogether 25 sites (5 categories x 5 replicates) were established in the area. We took samples from four different fallow age categories, i.e. less than (or equal to) 5 years old fallow (SA0-5, also referred to as new fallows), 6-10 years old fallow (SA6-10, also referred to as young fallows), 11-20 years old fallow (SA11-20, referred to as middle-aged fallows), and 21-30 years old fallow (SA21-30, also referred to as oldest fallows), and from old-growth forests (OF) as our control. Our old-growth forests are similar to the primary forests in terms of structure and floristic composition, never been used for kaingin and/or logging, however, may have experienced limited levels of anthropogenic disturbances, like fuelwood and fodder collection. We took samples only from sites that were 1 ha or more in size. The fallow age was confirmed by asking the kaingin farmer engaged in shifting cultivation at each site. A vegetation survey was conducted at each site and all trees ≥ 5 cm dbh were identified and measured for tree species richness, abundance and basal area (Mukul 2016). To measure site elevation, geographic position, and distance from the nearest oldgrowth control forest, we used a hand-held global positioning system (Model: Garmin eTrex). A digital plant canopy imager (Model: CID Bio-Science, CI-110/120) was used to measure the leaf area index of each site. Table 1 presents the key environmental attributes of our study sites.

Soil sample collection

We established four transects of 50 m \times 5 m in each of the 25 sites. Within each site, transects were spaced at least 5 m apart, parallel to each other and running from the boundary to the center of the site. Soil samples were collected from the beginning, middle, and end of each transect. Samples from the top 30 cm of topsoil were obtained. We collected soil samples from three different depths, i.e., 0-5 cm (also referred to as the top layer), 6–15 cm (the middle layer) and 16–30 cm (the bottom layer) using a standard soil auger (Manufacturer: AMS Inc., United States). The topsoil profile is widely reported to be most affected by land-use/cover change (Bruun et al., 2009; Sang et al., 2013); and is also the standard depth recognized by the Food and Agriculture Organization of the United Nations (FAO) and approved by Intergovernmental Panel on Climate

TABLE 1 Environmental	attributes of	of our	[,] studv	sites	on Le	evte	Island,	the	Philippines.
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Site category	Site environmental attribute ^a
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	EL	SL	PS	DIS	Ν	А	BA	LAI
New fallows	600.8 (±22.19) ^b	33.0 (±5.7)	1.16 (±0.21)	290.0 (±74.16)	6.0 (±7.38)	10.8 (±13.52)	0.89 (±1.32)	1.33 (±0.57)
Young fallows	549.0 (±72.41)	32.4 (±9.4)	1.14 (±0.13)	540.0 (±114.02)	38.4 (±4.83)	142.6 (±20.5)	2.8 (±0.81)	5.7 (±0.96)
Middle-aged fallows	567.2 (±49.24)	32.6 (±9.21)	1.34 (±0.24)	162.0 (±198.17)	39.2 (±9.52)	140.8 (±26.94)	3.85 (±1.86)	5.14 (±1.01)
Oldest fallows	574.8 (±35.35)	38.2 (±7.98)	1.14 (±0.22)	256.0 (±153.88)	45.8 (±5.93)	143.6 (±4.45)	3.92 (±1.39)	5.3 (±0.69)
Old-growth forest	512.4 (±54.77)	36.4 (±9.71)	_	_	45.2 (±4.21)	145.8 (±16.53)	7.81 (±2.23)	6.08 (±0.86)

^aWhere, EL, site elevation (m ASL), SL = slope (°), PS = patch size (ha), DIS = distance (from the nearest control forest site, m), N = tree species richness, A = tree species abundance, BA = tree basal area (m²/site), LAI = leaf area index (m²/site).

^bValues in the parenthesis indicate standard deviation of means.

TABLE 2 Soil physical and chemical characteristics of our study sites on Leyte Island, the Philippines.

Soil characteristic	Soil depth	Site category									
		New fallows	Young fallows	Middle-aged fallows	Oldest fallows	Old-growth forest					
Bulk density (g/cm ³)	Тор	0.85 (±0.09) ^a	0.89 (±0.16)	0.95 (±0.16)	0.91 (±0.11)	0.99 (±0.15)					
	Middle	0.95 (±0.09)	1.07 (±0.16)	1.05 (±0.14)	1.01 (±0.12)	1.08 (±0.12)					
	Bottom	1.02 (±0.08)	1.16 (±0.16)	1.11 (±0.17)	1.05 (±0.18)	1.12 (±0.05)					
Moisture content (%)	Тор	8.90 (±0.27)	8.94 (±0.47)	9.32 (±0.77)	9.11 (±0.18)	9.22 (±0.80					
	Middle	8.91 (±0.24)	9.17 (±0.50)	9.27 (±0.30)	9.19 (±0.23)	9.09 (±0.66)					
	Bottom	8.96 (±0.27)	9.27 (±0.31)	9.39 (±0.93)	9.09 (±0.31)	9.02 (±0.32)					
рН	Тор	5.02 (±0.27)	5.22 (±0.24)	5.12 (±0.51)	5.21 (±0.44)	5.17 (±0.32)					
	Middle	4.87 (±0.39)	5.07 (±0.30)	5.07 (±0.48)	5.16 (±0.43)	4.99 (±0.37)					
	Bottom	4.81 (±0.39))	5.02 (±0.33)	5.01 (±0.44)	5.15 (±0.37)	5.03 (±0.33)					

^aValues in the parenthesis indicate standard deviation of means.

Change (IPCC) for global voluntary carbon markets (FAO 2006). Altogether 900 soil samples were collected from 100 transects distributed across 25 sites. Samples were placed in plastic bags in the field immediately after collection and labeled before further processing in the laboratory.

Soil sample processing

All analyses were performed in the Philippines in the Visayas State University (VSU) Analytical Chemistry Laboratory. Samples were air-dried at room temperature before being passed through a 2 mm sieve to remove rocks, pebbles and plant materials (e.g., roots, coarse woody debris, etc.). We used a homogenized composite sample (n = 3) pooled for each soil depth (i.e., top, middle, and bottom layer) for each transect. Consequently, we had 300 composite soil samples for laboratory analysis to determine SOC, N, P, and K. Subsamples were oven dried (105°C for 48 h), and all calculations are reported on an oven-dry weight basis.

For SOC (% C) we oxidized the composite soil samples for each transect by heating with H_2SO_4 in the presence of dichromate as described in Heanes (1984). Unlike the traditional Walkley-Black method, external heating ensures complete oxidation and a more reliable estimate of SOC (Krishan et al., 2009; Congdon 2012). Soil N and P were determined colorimetrically following the single digestion method (Anderson and Ingram 1989), while soil total K was quantified using an atomic absorption spectrophotometer. Soil carbon and nutrient turnover are expressed as the percentage of respective soil parameters in our control old-growth forest. Soil physical and chemical properties, like bulk density, moisture content and pH were also analyzed following standardized procedures and protocols (Table 2).

Statistical analysis

We performed an analysis of variance (ANOVA) and Tukey's post-hoc analysis to test for significant differences among variables. We developed linear mixed-effect models (hereafter refer to as LMEM) to identify the effect of selected site environmental attributes on the SOC, N, P and K concentration in relation to our control old-growth forests using the package "nlme" (Pinheiro et al., 2011). For LMEM along with fallow age (FA), we first considered eight environmental variables as our explanatory variables (i.e. fixed factors)-elevation (EL), slope (SL), patch size (PS), distance from the nearest control old-growth forest (DIS), tree species richness (N), tree species abundance (A), basal area (BA), and leaf area index (LAI). Tree species richness, abundance, basal area, and leaf area index were expressed per site (0.1 ha area) basis. After performing the Pearson's correlation test only five site environmental variables (i.e., fallow age, elevation, slope, patch size, and leaf area index) were, however, found suitable for our final LMEM (Supplementary Table S1). We used sites nested in fallow categories as the random effect in our LMEM with SOC, N, P, and K at different soil depths as the response variable. All statistical analyses were performed using the 'R' Statistical Package (version 3.0.1; R Development Core Team 2019). We considered the Akaike Information Criterion corrected for small sample sizes (AICc) for the selection of our best models. In our study, only models within four AICc units were considered as competing models (Grueber et al., 2011). We used the package "MuMin" for model selections and to evaluate the contribution that different fixed effects had on explaining the variation in the response variables.

Results

Soil carbon and nutrients in regenerating tropical secondary forests

Soil organic carbon and the other macronutrients were consistently high in the top (i.e., 0-5 cm) soil layer of our fallow sites and control old-growth forest. We found higher SOC concentrations in our oldest fallows followed by young fallows and in new fallows (Table 3; Supplementary Figure S1). The differences in soil C distribution was not significant in the case of the top layer of soil, whilst in the middle (6-15 cm) and bottom (16–30 cm) layers it was significantly ($F_{4, 24} = 2.74, p < 0.05$) different across the sites. Tukey's post-hoc analysis also revealed significantly (p < 0.05) higher SOC concentrations in our middle and bottom soil layers in the oldest fallows followed by young and new fallows. The relative contribution of the SOC at different soil depths to soil C pool was consistently higher in the bottom layer of soil. In the topsoil layer, SOC concentration was relatively higher in our control old-growth forest (26.5%), followed by the middle-aged sites (25.5%) (Figure 2). In the case of the oldest fallows, the relative contribution of SOC in the bottom layer was comparatively higher (46.2%) than all other fallow age categories.

Interestingly soil total N concentrations were highest in the new fallow sites and in the top layer of the soil that we investigated (see Table 3). We found no significant difference in the soil N distribution between different site categories and soil depths as indicated by ANOVA. The relative contribution of soil N present in different soil depths to total N pool in our control old-growth

TABLE 3 Concentrations of soil organic carbon and nutrients across the sites of different fallow categories and old-growth forest on Leyte Island, the Philippines.

Soil parameter	Soil depth	Site category									
		New fallows	Young fallows	Middle-aged fallows	Oldest fallows	Old-growth forest					
SOC (mg C g ⁻¹)	Тор	61.66 (±6.82) ^a	65.36 (±20.53)	52.06 (±6.9)	67.88 (±19.18)	47.47 (±11.09)					
	Middle	36.81 (±6.89)	36.58 (±15.15)	29.23 (±3.6)	46.81 (±17.31)	25.44 (±5.1)					
	Bottom	29.96 (±6.66)	32.97 (±13.78)	24.99 (±3.73)	42.56 (±16.67)	22.68 (±5.38)					
Total N (mg N g ⁻¹)	Тор	4.33 (±0.77)	3.53 (±1.17)	3.46 (±1.14)	3.36 (±1.02)	3.4 (±0.5)					
	Middle	2.67 (±1.17)	1.88 (±0.85)	2.22 (±1.6)	2.21 (±0.89)	1.68 (±0.43)					
	Bottom	2.46 (±1.28)	1.89 (±0.69)	1.51 (±1.15)	1.86 (±1.11)	1.22 (±0.53)					
Total P (mg P g ⁻¹)	Тор	0.51 (±0.09)	0.53 (±0.11)	0.49 (±0.1)	0.61 (±0.14)	0.49 (±0.13)					
	Middle	0.34 (±0.14)	0.4 (±0.1)	0.34 (±0.09)	0.4 (±0.14)	0.28 (±0.14)					
	Bottom	0.30 (±0.13)	0.33 (±0.09)	0.23 (±0.07)	0.34 (±0.18)	0.2 (±0.12)					
K (mg K g ⁻¹)	Тор	1.02 (±0.22)	1.88 (±1.23)	1.52 (±0.99)	1.3 (±0.76)	3.18 (±1.58)					
	Middle	0.86 (±0.26)	1.81 (±1.29)	1.51 (±1.18)	0.9 (±0.48)	2.45 (±1.35)					
	Bottom	0.89 (±0.31)	1.76 (±1.3)	1.54 (±1.24)	0.91 (±0.53)	2.3 (±1.24)					

^aValues in the parentheses indicate the standard deviations of means under respective categories.



forest was comparable (30.4%, 32.7%, and 36.9%). Among all fallow age categories, the relative contribution of soil N in the bottom layer

(47.8%) was higher in our young fallow sites.

Similar to SOC and N there was no significant difference in total P concentrations in our study sites distributed across different fallow age categories and old-growth forests without any kaingin history (Supplementary Figure S1). The concentrations of soil P were, however, higher in the oldest fallows and in the topsoil layer of all fallow age categories although the differences were not statistically significant using ANOVA (Table 3). The relative contribution of soil P at the top and bottom layers varied substantially across our sites. Similar to other soil macronutrients, the contribution of soil P in the bottom layer to total soil P pool was consistently higher in our sites of different fallow age and old-growth forest (Figure 2).

Soil total K concentration was highest in the old-growth forest at all soil depths followed by young fallows and middle-aged fallows (Table 2). We found significantly higher ($F_{4, 24} = 3.16, p < 0.05$) soil K concentrations in the top layer of soil of our old-growth forest using ANOVA and Tukey's post-hoc analysis. The difference was, however, not significant in the middle ($F_{4, 24} = 2.15, p = 0.11$) and bottom ($F_{4, 24} = 1.73, p = 0.18$) soil layers.

Recovery of soil carbon and nutrients in regenerating tropical secondary forests

We found high SOC, soil N and soil P in our fallow sites when compared with control old-growth forest (Figure 3). Interestingly, the recovery of soil macronutrients was higher in our young fallow and oldest fallow secondary forest sites. Among the soil nutrients, soil K recovery was lowest across all fallow age categories. The pattern of soil K recovery was different to soil C and N and P, with relatively higher recovery in middle-aged fallows. Soil K recovery was lowest in the oldest fallows. Despite relatively higher variability within the sites in soil carbon and nutrient turnover, there was no significant difference between sites (of different fallow age) in the recovery of selected soil parameters.

Environmental controls on the recovery of soil carbon and nutrients in regenerating tropical secondary forests.

Patch size (PS) explained the highest amount of variation in soil C and nutrients recovery in regenerating



secondary forests (Table 4). Other site factors important in explaining the variations were slope (SL) and fallow age (FA). There was no influence of the distance from nearby old-growth forest (DIS) to the recovery of the soil C and nutrients across our sites of different fallow age categories. Patch size was found to be consistently important in explaining the variation in recovery for all the soil parameters and soil depths we examined (Table 5). In the case of soil P recovery, we also found that the slope of the sites was equally important in explaining the variation as the patch size.

Discussion

Variability in soil carbon and nutrients in regenerating tropical secondary forests

We did not find a clear or consistent effect of shifting cultivation on SOC, N, and P in our secondary forest sites in the upland Philippines. However, we did find a relatively greater influence of shifting cultivation on soil K concentration in our study sites in upland Philippines. The higher levels of soil C and nutrients (i.e., N and P) in our TABLE 4 Summary of LMEM between soil C and nutrient recovery with environmental attributes obtained using the package MuMin. Where, DF—Degree of Freedom, LL—Log Likelihood, AICc—Akaike Information Criterion corrected for small sample size.

Soil parameter	Soil depth	Explanatory variable ^a				Df	Ll	AICc	\triangle AICc	Weight
		FA	DIS	SL	PS					
SOC	Тор			X	Х	6	-86.12	190.71	0.0	0.47
					Х	5	-88.43	191.15	0.44	0.37
		Х			Х	6	-87.80	194.05	3.34	0.09
		Х		Х	Х	7	-85.54	194.42	3.71	0.07
	Middle			Х	Х	6	-93.6	205.65	0.0	0.44
					Х	5	-96.3	206.88	1.23	0.24
		Х			Х	6	-94.56	207.58	1.93	0.17
		Х		Х	Х	7	-92.25	207.83	2.18	0.15
	Bottom			Х	Х	6	-95.67	209.79	0.0	0.42
					Х	5	-98.48	211.25	1.46	0.20
		Х		Х	Х	7	-94.01	211.36	1.57	0.19
		Х			Х	6	-96.47	211.39	1.6	0.19
Total N	Тор				Х	5	-88.43	191.14	0.0	0.42
				Х	Х	6	-86.93	192.31	1.17	0.24
		Х			Х	6	-87.07	192.6	1.46	0.20
		Х		Х	Х	7	-85.01	193.35	2.21	0.14
	Middle				Х	5	-101.9	218.08	0.0	0.40
				Х	Х	6	-100.1	218.71	0.63	0.29
		Х			Х	6	-100.5	219.53	1.44	0.19
		Х		Х	Х	7	-98.65	220.64	2.56	0.11
	Bottom				Х	5	-106.9	228.03	0.0	0.30
		Х			Х	6	-104.9	228.29	0.27	0.26
				Х	Х	6	-104.9	228.36	0.33	0.25
		Х		Х	Х	7	-102.8	229.0	0.97	0.18
Total P	Тор			Х	Х	6	-80.12	178.7	0.0	0.51
					Х	5	-82.76	179.8	1.1	0.30
		Х			Х	6	-81.67	181.8	3.09	0.11
		Х		Х	Х	7	-79.55	182.42	3.72	0.08
	Middle				Х	5	-93.35	200.98	0.0	0.42
		Х		Х		6	-91.42	201.3	0.32	0.36
		Х			Х	6	-92.33	203.12	2.14	0.14
		Х		Х	Х	7	-90.51	204.35	3.37	0.08
	Bottom			Х	Х	6	-98.21	214.89	0.0	0.45
					Х	5	-100.8	215.81	0.92	0.28
		Х		Х	Х	7	-96.90	217.13	2.24	0.15
		Х			Х	6	-99.47	217.41	2.52	0.13
К	Тор			х	Х	6	-82.95	184.37	0.0	0.76
	-	Х		Х	Х	7	-82.20	187.73	3.36	0.14
					Х	5	-87.04	188.36	3.99	0.10
	Middle			х	Х	6	-88.49	195.44	0.0	0.82
		Х		Х	Х	7	-87.57	198.47	3.03	0.18
	Bottom			Х	Х	6	-89.94	198.35	0.0	0.82
		Х		Х	Х	7	-89.02	201.37	3.02	0.18

^aFA, fallow age, DIS, distance from the old–growth forest, SL, slope, PS, patch size.

Soil parameter	Soil depth	Explanatory	No. of candidate model			
		FA	DIS	SL	PS	
SOC	Тор	0.16 (2) ^b	_	0.54 (2)	1.0 (4)	4
	Middle	0.32 (2)	_	0.59 (2)	1.0 (4)	4
	Bottom	0.38 (2)	—	0.61 (2)	1.0 (4)	4
Total N	Тор	0.34 (2)	_	0.37 (2)	1.0 (4)	4
	Middle	0.31 (2)	_	0.40 (2)	1.0 (4)	4
	Bottom	0.45 (2)	_	0.44 (2)	1.0 (4)	4
Total P	Тор	0.19 (2)	_	0.59 (2)	1.0 (4)	4
	Middle	0.22 (2)	_	0.44 (2)	1.0 (4)	4
	Bottom	0.27 (2)	_	0.59 (2)	1.0 (4)	4
K	Тор	0.14 (1)	_	0.90 (2)	1.0 (3)	3
	Middle	0.18 (1)	_	1.0 (2)	1.0 (2)	2
	Bottom	0.18 (1)	_	1.0 (2)	1.0 (2)	2

TABLE 5 The relative importance of site environmental attributes in the final LMEM.

^aWhere, FA = fallow age, DIS = distance from the old–growth forest, SL = slope, PS = patch size.

^bValues in the parenthesis indicate the number of models containing respective explanatory variable.

young fallows could be due to the initial release of carbon and nutrients from plant materials to soil following the burning (Giardina et al., 2000; Tanaka et al., 2001). Interestingly, we find a high variability of SOC within our sites of different fallow categories, which may be due to the past kaingin history attributed to previous fallow cycles that we could not capture. It may also be due to litterfall dynamics or soil below-ground activity, such as microbial presence or fine root biomass in specific sites (see-Sayer et al., 2011; Langee et al., 2014; Filho et al., 2015; Sarai et al., 2022). Heating of soil during burning acts as an important mechanism of nutrient release and the initial burning associated with shifting cultivation is reported to increase SOC by the addition of organic matter in soil via ash (Tanaka et al., 2001; Thomaz 2017). Burning also results in losses of soil available N due to increased volatilization, although, the losses of soil P to the atmosphere due to volatilization is relatively low (Giardina et al., 2000; Romanyá et al., 2001; Taylor et al., 2017). In shifting cultivation landscapes, burning also results in a decrease in soil microbial activity and associated biomass at the initial stage that may negatively affect soil carbon (Tanaka et al., 2001; Mukul and Herbohn 2016; Hattori et al., 2019). In regenerating forests, N deposition is also coupled with enhanced productivity, thus greater C in soil and living biomass (Thomaz 2017).

The relatively lower concentration of soil C and P in our middle–aged fallows could be attributed to the increasing nutrient uptake by regenerating forests during successional development (Filho et al., 2015). Similarly, relatively greater concentrations of soil C, N, and P in oldest fallows could be

due to higher litterfall, decomposition and microbial activity in regenerating forests recovering after shifting cultivation (see—Sayer et al., 2011; Lange et al., 2015; Paudel et al., 2015; Teegalapalli et al., 2018; Jones et al., 2019). A positive effect of shifting cultivation on soil available P and K in secondary forests has also been reported by Brand and Pfund (1998). Temjen et al. (2022) found a relatively higher concentration of SOC, available N, P, and K with increase in fallow period in Nagaland, India. In contrast to our findings, Garcia-Oliva et al. (1999), however, have found a 32% initial decrease in soil organic C stock due to fire and combustion in Mexico, and Yang et al. (2003) found a reduction in the soil available N and P due to shifting cultivation in China. Soil K was consistently low at all our fallow sites, although Yang et al. (2003) reported an increase in soil K after fire and shifting cultivation. Other than combustion during the burning of vegetation, soil runoff associated with forest clearing during shifting cultivation has also been reported to cause nutrient losses (Brand and Pfund 1998; Rodenburg et al., 2003; Thomaz 2017; Hattori et al., 2019). Such generalizations, however, require careful consideration because in the tropics smallholder farmers usually prefer sites with greater soil fertility for their shifting cultivation practice (Mertz et al., 2008).

Influence of site environmental factors on the recovery of soil carbon and nutrients

We found that site environmental parameters, mainly patch size, have the greatest influence on the recovery of

SOC, total N, and P; and we have found that the same factors influence the aboveground biomass and carbon recovery in regenerating forests after shifting cultivation in Philippines uplands (see Mukul et al., 2016b). As mentioned earlier, the higher levels of soil organic carbon and nutrients in our fallow sites could also be due to the initial release of carbon and nutrients following the burning of biomass and may not portray the actual recovery in young fallows. In the case of soil K, both the patch size and slope were found to be equally important for its recovery. Patch size, however, was found to be the most important factor as models within $\triangle AICc = 4$ are considered as equivalent and explain the same amount of variation as other more complex models incorporating more interactions of variables (Bartoń 2011). Due to the lack of reliable information about past fallow cycles, we did not take account of this in our analysis. It is thus unclear what impact past fallow cycles may have had on the soil C and nutrients in our fallow sites. In an Indonesian study, the number of fallow cycles did not influence soil nutrients, like soil P, although the presence of deep and fine root systems in soil was found to be more important than fire history in the recovery of soil P (Lawrence and Schlesinger 2001). A similar observation was also made by Bruun et al. (2009) where they found higher growth of pioneer species in fallows due to their shallow root systems and the ability to take up nutrients from the top layer. The effect of land-use and site environmental attributes on soil C and nutrients is evident from many other studies (see-Neill et al., 1997; van Noordwijk et al., 1997; Hatter et al., 2010; Parras-Alcantara et al., 2013; Fernandez-Romero 2014). A long fallow period is critical to restoring essential soil nutrients, and it may take as long as 20 years to restore soil nutrients to a level similar to undisturbed forests (Funakawa et al., 2009; Rahamiralala et al., 2010). Forest structure and species composition may also influence the dynamics of soil C and nutrients, and vice versa (Lawrence et al., 2005; Firn et al., 2007; Paoli et al., 2008; Lange et al., 2014; Raich et al., 2014). Furthermore, spatial heterogeneity in soil in terms of texture, clay mineralogy and site topography and climate may vary and can also override the effects of other environmental factors (Bruun et al., 2006; Nottingham et al., 2015).

Conclusion

Our study finds that shifting cultivation may not be as detrimental to soil quality at least on the soil type and climate we studied, and that geographic and site-specific conditions may be important in determining the impact that shifting cultivation has on soil properties. We also found greater availability in the soil C, N and P in regenerating tropical secondary forests compared to old-growth forests after shifting cultivation abandonment. In young fallows, this may be attributed to the initial release of carbon and nutrients from slashing and burning of plant biomass available in the sites. In our old-growth control forests, there might be much more C, N, P, and K present in the biomass (above and below ground). In older fallow sites, repeated clearing, burning and run-off from the cleared landscape presumably led to the loss of much of the stored nutrients. In addition, over time, more of these nutrients are expected to be sequestered in the developing vegetation, leading concentrations of soil macronutrients to change over time due to plant uptake and recycling through litter fall and decomposition. We find, patch size together with the slope and fallow age significant explanatory factors associated with the recovery process of soil macronutrients.

In tropical regions shifting cultivation is still a dominant land-use and likely to contribute to the food security and livelihoods of smallholder farmers for many years ahead. The traditional view of shifting cultivation as environmentally detrimental in terms of its impacts on soil is not well established (Murty et al., 2002). A better understanding of this issue, with systematic and long-term studies, is crucial for the restoration and management of tropical forests recovering after shifting cultivation and other anthropogenic disturbances.

Data availability statement

The raw data supporting the conclusion of this article will be made available by the authors, without undue reservation.

Author contributions

SM: conceptualization. SM and JH: project administration. SM, JH, AF, and RC: methodology. SM, AF, and RC: investigation. SM: data curation, formal analysis, and writing. JH, AF, and RC: review and editing. JH: supervision and funding. All authors contributed to the article and approved the submitted version.

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Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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Supplementary material

The Supplementary Material for this article can be found online at: https://www.frontiersin.org/articles/10.3389/fenvs.2022. 1076506/full#supplementary-material

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