



Response of Subtropical Coastal Sediment Systems of Okinawa, Japan, to Experimental Warming and High pCO_2

Rumana Sultana¹, Beatriz E. Casareto^{1,2*}, Rumi Sohrin^{2,3}, Toshiyuki Suzuki¹, Md. Shafiul Alam¹, Hiroyuki Fujimura⁴ and Yoshimi Suzuki¹

¹ Department of Environment and Energy Systems, Graduate School of Science and Technology, Shizuoka University, Shizuoka, Japan, ² Research Institute of Green Science and Technology, Shizuoka University, Shizuoka, Japan, ³ Department of Geoscience, Shizuoka University, Shizuoka, Japan, ⁴ Department of Chemistry, Biology and Marine Science, University of the Ryukyus, Okinawa, Japan

OPEN ACCESS

Edited by:

Hajime Kayanne, University of Tokyo, Japan

Reviewed by: David I. Kline,

University of California, San Diego, USA Atsushi Watanabe, Tokyo Institute of Technology, Japan

*Correspondence:

Beatriz E. Casareto dcbeatr@ipc.shizuoka.ac.jp; casaretobe@gmail.com

Specialty section:

This article was submitted to Coral Reef Research, a section of the journal Frontiers in Marine Science

Received: 04 March 2016 **Accepted:** 06 June 2016 **Published:** 28 June 2016

Citation:

Sultana R, Casareto BE, Sohrin R, Suzuki T, Alam MS, Fujimura H and Suzuki Y (2016) Response of Subtropical Coastal Sediment Systems of Okinawa, Japan, to Experimental Warming and High pCO₂ Front. Mar. Sci. 3:100. doi: 10.3389/fmars.2016.00100 Increasing seawater temperatures and CO₂ levels associated with climate change affect the shallow marine ecosystem function. In this study, the effects of elevated seawater temperature and partial pressure of CO₂ (pCO₂) on subtropical sediment systems of mangrove, seagrass, and coral reef lagoon habitats of Okinawa, Japan, were examined. Sediment and seawater from each habitat were exposed to four pCO₂-temperature treatments, including ambient pCO2- ambient temperature, ambient pCO2-high temperature (ambient temperature + 4° C), high ρ CO₂ (936 ppm)-ambient temperature, and high pCO₂-high temperature. Parameters including primary production, nutrient flux, pigment content, photosynthetic community composition, and bacterial abundance were examined. Neither high temperature nor high pCO₂ alone impacted mangrove and seagrass sediment primary production significantly (Tukey's test, P > 0.05). However, the combined stress significantly (Tukey's test, P < 0.01) increased primary production in these two habitats. In sediments from the coral reef lagoon, single and combined stress treatments induced a shift from heterotrophy to autotrophy. Significant increases in net primary production (Tukey's test, P < 0.01), and gross primary production (Tukey's test, P < 0.05) under the combined stress suggested that benthic microalgae in mangrove and seagrass sediments were more responsive to high temperature and ρCO_2 conditions than those in coral reef lagoon sediments. Additionally, combined stress significantly increased the sediment chlorophyll a content (Tukey's test, P < 0.05) in all habitats. These increases were associated with increased net primary production, indicating that combined stress stimulates primary production activity by the photosynthetic benthic microalgae in all habitats. Diatom activity increased, as silicate uptake increased in all habitats. Microbial abundance significantly increased under the combined stress treatment (Tukey's test, P < 0.01), but did not exceed autotrophic activity. Respiration did not change significantly in any of the three habitats (Tukey's test, P > 0.05) under the combined stress, suggesting that heterotrophic processes were less

1

affected by the combined stress than autotrophic processes. In summary, mangrove and seagrass sediments minimize the negative impacts of elevated temperature and pCO_2 via increased primary production and carbon storage. Lagoonal sediments also act as a carbon sink under temperature and ocean acidificationstress.

Keywords: primary production, temperature, pCO2, nutrients and bacterial abundance

INTRODUCTION

The of marine ecosystems in mitigating the impacts of climate change has generated considerable interest owing to the capacity of these ecosystems to trap and sequester carbon (Mcleod et al., 2011; Miyajima et al., 2015; Ricart et al., 2015). Mangroves, tidal salt marshes, seagrass beds, and coral reefs are key ecosystems that serve important roles in protecting coastal communities worldwide from storm surges, sea level rise, sediment runoff, and food insecurity (Beck et al., 2001; Mumby, 2006; Barbier et al., 2011; Sifleet et al., 2011; Cullen-Unsworth and Unsworth, 2013; Robertson and Alongi, 2013). These ecosystems help to mitigate the effects of environmental change (e.g., increasing atmospheric and oceanic CO₂ and temperature) via carbon sequestration, and the storage of atmospheric and seawater carbon (Duarte et al., 2005, 2010; Bouillon et al., 2008; Kennedy et al., 2010; Donato et al., 2012; Pendleton et al., 2012; Chmura, 2013). Although, shallow marine ecosystems are currently subjected to a combination of global and regional stressors, few experimental studies have examined shallow benthic areas to understand the consequences of increased temperature and ocean acidification (OA) (Connell et al., 2013; Johnson et al., 2013). Additionally, it is difficult to conclusively determine the impact of different environmental stressors in shallow-water sediment systems (Alsterberg et al., 2011, 2012) since the combinations of environmental stressors may have synergistic (Aswani et al., 2015) or antagonistic effects, resulting in unexpected patterns of variation (Folt et al., 1999). Warming can negatively decouple (i.e., mismatch) the relationship between phytoplankton and zooplankton, and it causes an earlier peak of primary production (Winder and Schindler, 2004; Woodward et al., 2010). While some studies have suggested that many marine autotrophs are insensitive to changes in pCO₂ (Engel et al., 2007), others have revealed that some seagrasses (Zimmerman et al., 1997), macroalgae (Gao et al., 1993), microalgae (Kayanne et al., 2005), diatoms, coccolithophorids (Riebesell et al., 2000; Zondervan et al., 2001; Casareto et al., 2009), and cyanobacteria (Qiu and Gao, 2002) exhibit high rates of photosynthesis under elevated pCO₂ conditions (Delille et al., 2005).

Major ecosystem functions in shallow water sediments are primarily controlled by microorganisms; accordingly, the present study focused on benthic microalgal and microbial communities, and their key ecosystem functions (i.e., primary production and nutrient flux). Primary production, microbial activity, oxygen flux, and carbon and nitrogen at the sediment– seawater interface are indicators of the integrated net responses of sediment communities to environmental and anthropogenic stressors (Alsterberg and Sundbäck, 2013). Mangrove ecosystems are characterized by high primary productivity, a permanent

exchange with terrestrial and marine ecosystems (Jennerjahn and Ittekkot, 2002) and efficient nutrient filtration between land and sea (Bouillon et al., 2008). Seagrass ecosystems are important for their carbon storage capacity (Miyajima et al., 2015) and high primary productivity (Kennedy et al., 2010; Ricart et al., 2015). Benthic microalgae in seagrass sediments, such as benthic diatoms and mats of filamentous cyanobacteria, maintain net autotrophic conditions under increased temperatures (Alsterberg et al., 2011). Some studies have examined the impact of high temperatures on seagrass sediment systems, but the influence of OA is still unclear. Lagoonal sediment responses under stress are not known owing to limited sediment system studies in coral reef lagoon environments. Additionally, many studies have emphasized the effects of warming on single species or trophic levels (Montoya and Raffaelli, 2010), but not on whole communities.

Most sediment system studies have considered the impact of single (Alsterberg et al., 2011; Johnson et al., 2013) or multiple stressors (Alsterberg et al., 2013) in a single habitat. Whether sediment from different subtropical habitats (e.g., mangroves, seagrasses, and coral reef lagoons) show different patterns under elevated pCO_2 , temperature stress, and their combination, and whether such differences can be observed when studying primary production, nutrient fluxes, photosynthetic community composition, pigment content, and bacterial dynamics also need to be investigated. Answering these questions will provide opportunities to assess shallow marine sediment system's responses under elevated temperature and OA.

In the present study, we exposed subtropical sediments from mangrove, seagrass, and coral reef habitats of Okinawa, Japan to four pCO₂-temperature treatments, including ambient pCO₂-ambient temperature, ambient pCO₂-high temperature (ambient temperature + 4° C), high pCO₂ (936) ppm)-ambient temperature, and high pCO₂-high temperature. Parameters including primary production, nutrient fluxes, pigment content, photosynthetic community composition, and bacterial abundance were examined. The following questions were addressed. (1) Do sediments from the three habitats respond differently with respect to primary sediment production under individual and combined stress? (2) What is the effect of stress on nutrient flux? (3) How do sediment and seawater microbial abundances change under elevated temperatures and OA? To address these questions, incubation experiments were performed using sediment and seawater samples from the mangrove, seagrass, and coral reef habitats exposed to elevated temperature and pCO2 treatments.

METHODOLOGY

Site Description

The study was conducted in a subtropical mangrove area (station 1) at Yagachi Island (26°38'N and 128°00'E), a seagrass area (station 2) at Bise (26°42'N and 127°52'E), and a coral reef lagoon area (26°39"N and 127°51'E) at Sesoko reef in Sesoko Island (station 3), Okinawa, Japan (Figure 1). The tidal flat of Yagachi Island was dominated by Bruguiera gymnorrhiza. During high tide, the water level rises to 0.6 m in the mangrove area. Thalassia hemprichii dominated the seagrass area of Bise, which included the intertidal zone covered by a seagrass meadow on the back reef (1.5 m at high tide and 0.5 m at low tide). In Sesoko reef, the lagoonal bed was covered by 10% living corals, 50% coral rubble, 30% sand, and 10% microalgae and turf algae (Nakano and Nakai, 2008). Depths during sampling were 0.6, 0.5, and 0.5 m for mangrove (high tide), seagrass (low tide), and coral reef lagoon (low tide) sediment and seawater. In Yagachi, Bise, and Sesoko, in situ temperature, salinity, pH, and alkalinity were measured during the week of the experiments (Table 1, ambient temperature and ambient pCO₂ [ATAP]).

Incubation Conditions

Incubation experiments were performed with subtropical sediment and seawater samples obtained from each habitat (mangrove, seagrass, and coral reef lagoon) in the laboratory. Mangrove sediments were generally more fine-grained, with a higher porosity, a higher organic carbon content and a higher carbon-nitrogen ratio than seagrass and coral reef lagoon sediments (Tukey's test, P < 0.05; Supplementary Table 1). Sediment porosity and organic carbon content were measured following the methods of Buchanan (1984) and Miyajima (2015), respectively. Sediment samples were collected using sediment collection cores (4.5 cm radius). Samples were treated following Cornwell et al. (2014) to adjust for sampling disturbances. Intact sediment was transferred to incubation chambers, maintaining their natural heterogeneity and complexity. In each incubation chamber (chamber radius, 5 cm), the upper 3 cm of sediment was kept in the bottom and 1 L of seawater sampled at the same location was added. Levels of pCO₂ were adjusted by adding CO₂-saturated seawater (100%) to the incubation chambers at different volumes until the desired pH values equivalent to the desired pCO₂ levels were obtained (Supplementary Table 2). CO₂ saturated seawater was prepared following the methods of Casareto et al. (2009); briefly, CO₂ gas was bubbled into natural filtered $(0.2 \,\mu\text{m})$ seawater for about 1 h. The computer program CO_2 SYS was used to estimate the CO_3^{2-} , HCO_3^{-} , CO₂(aq.), TIC, and pCO₂ from the pH, temperature, total alkalinity, and salinity (Lewis and Wallace, 1998). Table 1 shows the estimates of pCO₂. Four pCO₂-temperature treatments, including ambient pCO₂-ambient temperature, ambient pCO₂high temperature (ambient temperature $+ 4^{\circ}$ C), high *p*CO₂ (936 ppm)-ambient temperature, and high pCO₂-high temperature (Table 1, Supplementary Figure 1), with three replicates per treatment were used. pCO2 conditions of 400 ppm (in-situ pCO₂) and 936 ppm (Intergovernmental Panel on Climate Change (IPCC), 2013), were chosen as present day and nearfuture scenarios, respectively. 27.5, 28.1, and 27.1°C were chosen as ambient temperatures (AT), which were similar to in-situ temperatures in Yagachi, Bise, and Sesoko, respectively, during the incubation experiments (Table 1), and ambient (in-situ temperature) + 4°C (Intergovernmental Panel on Climate Change (IPCC), 2013) was chosen as the stress temperature. Around Okinawa, the average temperature for the month preceding the experiment was also 27°C (unpublished data). Incubations were performed in triplicate for each condition and habitat. Incubation vessels were maintained in water baths under four different incubation conditions (Table 1, Supplementary Figure 1) and run for 24 h following natural illumination (276- $0\,\mu$ mol m⁻² s⁻¹ during 12:12 h light: dark cycle) during summer (August-September, 2014). Heating and cooling systems were used to keep the temperature conditions in the water baths stable.

The pH (NBS scale) was measured with an Orion 4 Star pH sensor and an 8156BNUWP pH electrode (Thermo Scientific, Inc., Waltham, MA, USA) calibrated with NIST (NBS)-scaled buffer solutions (Mettler pH 10.012, 6.865, and 4.008 buffers) at 25°C. The accuracy of pH measurements was ± 0.005 pH units (1 σ , n = 10). Seawater samples for total alkalinity (TA) measurements were collected from the field and each incubation chamber before starting the incubation. Collected seawater was filtered with 0.45-µm cellulose acetate filters. TA was determined via titration with 0.1 mol L^{-1} HCL using a TA titrator (ATT-05; Kimoto, Japan), and computed using the Gran plot method (Stumm and Morgan, 1981). Titrations were performed on 30.000 g of filtered seawater that was weighed on a high-precision balance (GH-202, A&D, Tokyo, Japan, 0.00001 accuracy). The precision of the TA measurements was $\pm 2 \,\mu$ mol kg⁻¹ (1 σ , n = 10). The accuracy was within 10 μ mol kg⁻¹, as verified using certified reference material (CRM Batch # 50) distributed by A. Dickson of the Marine Physical Laboratory, University of California (San Diego, CA, USA). The carbonate system in seawater was calculated using the apparent equilibrium constants K'₀ from Weiss (1974), and K'₁ and K'₂ from Mehrbach et al. (1973), as described by Millero (1979). The constants are among the suitable combinations for the calculation of carbonate systems with the NBS pH scale. Recently, the Tris buffer pH scale with an acid dissociation constant for the total hydrogen ion concentration has been recommended for the precise calculation for seawater carbonate systems (Dickson et al., 2007). However, it has been suggested that the pCO_2 value from the NBS scale is very close to that of the Tris scale when an appropriate combination of equilibrium constants is used (Millero, 1979).

Measurements

Nutrients $(NO_3^- + NO_2^-, NH_4^+, PO_4^{3-}, and Si(OH)_4)$, tracer pigments (chlorophyll *a*, chlorophyll *b*, chlorophyll c_2 , fucoxanthin, zeaxanthin, and β -carotene), net primary production (NPP), gross primary production (GPP), respiration (R), and bacterial abundance were analyzed. Samples for pigment analyses and bacterial abundances were obtained at the initial (0 h) and final (24 h) stages. Dissolved oxygen (DO) and nutrients were measured at the initial (0 h), middle (after 12 h of light), and



	рН	Alkalinity (mmol kg ⁻¹)	Temperature (°C)	Salinity PSU	[CO ₃ ^{2–}] (mmol kg ⁻¹)	[HCO ₃ -] (mmol kg ⁻¹)	[CO ₂ (aq.)] (mmol kg ⁻¹)	TIC (mmol kg ⁻¹)	<i>р</i> СО ₂ (ppm)
MANG	ROVE								
ATAP	8.23 ± 0.07	2.31 ± 0.01	27.5 ± 0.2	28.0 ± 0.0	0.2 ± 0.0	1.80 ± 0.01	0.01 ± 0.00	2.01 ± 0.01	408 ± 7
ATHP	7.97 ± 0.01	2.50 ± 0.01	27.5 ± 0.2	28.0 ± 0.0	0.2 ± 0.0	2.11 ± 0.01	0.02 ± 0.00	2.30 ± 0.01	973 ± 1
HTAP	8.22 ± 0.04	2.31 ± 0.01	31.5 ± 0.4	28.0 ± 0.0	0.2 ± 0.0	1.70 ± 0.00	0.01 ± 0.00	2.01 ± 0.00	423 ± 8
HTHP	7.97 ± 0.02	2.50 ± 0.00	31.5 ± 0.4	28.0 ± 0.0	0.2 ± 0.0	2.10 ± 0.00	0.02 ± 0.00	2.30 ± 0.02	973 ± 1
SEAGR	ASS								
ATAP	8.21 ± 0.03	2.32 ± 0.01	28.1 ± 0.4	33.0 ± 0.0	0.2 ± 0.0	1.81 ± 0.00	0.01 ± 0.00	2.01 ± 0.01	400 ± 6
ATHP	7.91 ± 0.05	2.30 ± 0.00	28.1 ± 0.4	33.0 ± 0.0	0.1 ± 0.0	2.01 ± 0.01	0.03 ± 0.00	2.20 ± 0.02	953 ± 9
HTAP	8.22 ± 0.06	2.32 ± 0.00	32.2 ± 0.4	33.0 ± 0.0	0.2 ± 0.0	1.71 ± 0.00	0.01 ± 0.00	2.01 ± 0.03	407 ± 8
HTHP	7.91 ± 0.05	2.30 ± 0.01	32.2 ± 0.4	33.0 ± 0.0	0.1 ± 0.0	2.01 ± 0.03	0.02 ± 0.00	2.10 ± 0.02	953 ± 13
CORAL	REEF LAGOON								
ATAP	8.18 ± 0.03	2.21 ± 0.00	27.1 ± 0.6	33.5 ± 0.0	0.2 ± 0.0	1.71 ± 0.01	0.01 ± 0.00	1.91 ± 0.03	424 ± 1^{-1}
ATHP	7.91 ± 0.04	2.30 ± 0.01	27.1 ± 0.8	33.5 ± 0.0	0.1 ± 0.0	2.01 ± 0.01	0.02 ± 0.00	2.20 ± 0.02	950 ± 12
HTAP	8.22 ± 0.05	2.21 ± 0.01	31.5 ± 0.5	33.5 ± 0.0	0.2 ± 0.0	1.80 ± 0.01	0.01 ± 0.00	2.01 ± 0.01	404 ± 8
HTHP	7.91 ± 0.07	2.21 ± 0.01	31.5 ± 0.5	33.5 ± 0.0	0.1 ± 0.0	1.91 ± 0.03	0.02 ± 0.00	2.11 ± 0.01	950 ± 1

TABLE 1 | Incubation Conditions.

ATAP, Ambient temperature and ambient pCO2; ATHP, Ambient temperature and high pCO2; HTAP, High temperature and ambient pCO2; and HTHP, High temperature and high pCO2.

final (24 h) stages. Parameter changes (Δ values) were calculated by subtracting the initial values from the final values.

Nutrients

To measure nutrients, seawater samples were collected (from the field and each replicate incubation chamber) in previously acid-washed 100-mL polyethylene bottles and were preserved at -20°C until measurement. Pore waters were collected following the methods of Brito et al. (2010). The concentrations of nitrate $(NO_3^-) + nitrite (NO_2^-)$, ammonium (NH_4^+) , phosphate (PO_4^{3-}) , and silicate $(Si(OH)_4)$ were analyzed using an auto analyzer (TRAACS 2000; Bran and Luebbe, Norderstedt, Germany) according to the methods of Hansen and Koroleff (2007). The detection limits were $0.043 \,\mu\text{M}$ for $\text{NO}_3^- + \text{NO}_2^-$, $0.01 \,\mu\text{M}$, for NH_4^+ , $0.020 \,\mu\text{M}$ for PO_4^{3-} , and $0.310 \,\mu\text{M}$ for $\text{Si}(\text{OH})_4$. Reproducibility (precision) of nutrient analyses were $\pm 0.2\%$ for $\text{NO}_3^- + \text{NO}_2^-$, $\pm 1.2\%$ for NH_4^+ , $\pm 0.7\%$ for PO_4^{3-} , and $\pm 0.5\%$ for $\text{Si}(\text{OH})_4$.

Pigments and Benthic Microalgal Observations

To assess the pigment concentration, three sediment cores were collected from each replicate incubation chamber using a 1-mL plastic syringe with a cut-off tip. Additionally, to quantify changes in the chlorophyll a concentration at the sediment surface inside the incubation chamber between initial conditions and combined stress conditions, 1 g of sediment was collected

at radii of 0, 2.5, and 5 cm (total radius, 5 cm) for angles (θ) of 30, 60, 90, 120, 150, 180, 210, 240, 270, 300, 330, and 360°. Samples were kept in separate pre-combusted GF/F filters and preserved at -20° C until used for measurements. Pigment analyses were carried out following the methods of Zapata et al. (2000) using HPLC (high-performance liquid chromatography) (LC-30AD; Shimadzu, Kyoto, Japan). The HPLC system was equipped with a ZORBAX Eclipse Plus C8 column (2.1 \times 50 mm; Agilent Technologies, Santa Clara, CA, USA). Pigments were extracted using a combination (1:3) of dry sediment and acetone (90% final concentration; Alsterberg et al., 2011), sonicated for 15 min, and maintained overnight in the dark at -30° C. Extracts were filtered using a syringe filter (0.2 µm, Millex-LG; Millipore, Bedford, MA, USA; Suzuki et al., 2015) and 1.0 mL of each extract was mixed with 0.2 mL of deionized water (Milli-Q water, Millipore) immediately before injections (Zapata et al., 2000) in the HPLC system. Separated pigments were detected spectrophotometrically using a photodiode array detector (SPD-M30AD; Shimadzu, Kyoto, Japan) with an optical resolution of 1.2 nm across the 320-720 nm bandwidth. For micro-algal observations, sediment samples were obtained from each incubation chamber using a 2.5-mL plastic syringe with a cut-off tip and kept frozen $(-20^{\circ}C)$ until analysis. After thawing, the samples were diluted with filtered seawater following the methods of Sundback et al. (2010). Epifluorescence microscopy (Eclipse/E6000; Nikon, Tokyo, Japan) and an inverted microscope (Nikon Bk-201) were used to identify the dominant photosynthetic community following Chihara and Murano (1997).

Bacterial Abundance

Bacterial abundance was measured for both seawater and sediment. For seawater, 50 mL of sample was collected from each of the replicate incubation chambers and preserved with 25% glutaraldehyde (1% final concentration). For sediment, three sediment cores were collected from each replicate incubation chamber using a 1-mL plastic syringe with a cut-off tip. To estimate bacterial abundance within the whole incubation chamber at initial conditions and after treatment, 1 g of sediment was collected at the center and radii of 2.5 and 5 cm (total radius, 5 cm) in lines separated by a 30° (θ) angle. Sediments were preserved with 25% glutaraldehyde (4% final concentration) until analysis. Sediments were diluted with artificial seawater (1:99 combinations) and pre-filtered with 0.8- μ m polycarbonate filters to remove large particles. Heterotrophic bacteria from seawater and sediment samples were collected on 0.2-µm black polycarbonate filters by filtering 10-mL aliquots of seawater that were previously stained with 4',6-diamidino-2-phenylindole (DAPI) (Porter and Feig, 1980). Filters were mounted on slides and cells were counted under UV excitation using epifluorescence microscopy (Eclipse/E6000; Nikon). Growth rates and doubling were calculated following the methods of Stanier et al. (1979).

Primary Production

NPP and R were measured based on oxygen flux during day (light) and night (dark) conditions, respectively. Initial (0 h),

middle (after 12 h of light), and final (24 h) seawater samples were obtained using sterile plastic syringes. Dissolved oxygen (DO) was measured with an Orion 4 Star sensor and a polarographic DO Probe on a 5-ft Cable (Thermo Scientific, Inc., Waltham, MA, USA). NPP was estimated as DO production during 12 h of illumination. R was estimated as the night-time changes in DO. It is considered that R in the light equals R in the dark (Del Giorgio and Williams, 2005; Alsterberg et al., 2011), such that R = R (12 h of dark) × 2. GPP was calculated as follows: GPP = NPP + R. All estimates of GPP, NPP, and R were converted to carbon (C) using the Redfield–Richards empirical formula (Libes, 2009).

Statistical Analysis

Two-way analyses of variance (ANOVA) were used for comparisons, and incubation conditions (ambient temperature and ambient pCO_2 [ATAP], ambient temperature and high pCO_2 [ATHP], high temperature and ambient pCO_2 [HTAP], and high temperature and high pCO_2 [HTHP]) and habitats (mangrove, seagrass, and coral reef lagoon) were used as fixed factors. Significant differences between means were assessed using *post-hoc* Tukey's tests. A regression analysis was used to evaluate relationships between variables. The homogeneity of the variances among replicates was assessed by applying Levene tests. Statistical analyses were performed using SPSS (version 19).

RESULTS

Abiotic Parameters

In Yagachi, *in situ* temperature, salinity, pH, and alkalinity during the week of the experiments were $27.5 \pm 0.2^{\circ}$ C, 28 PSU, 8.23 ± 0.07 , and 2.38 ± 0.01 mmol kg⁻¹, respectively. Similarly, in Bise, *in situ* temperature, salinity, pH, and alkalinity were $28.1 \pm 0.4^{\circ}$ C, 33 PSU, 8.21 ± 0.03 , and 2.37 ± 0.02 mmol kg⁻¹, respectively. In Sesoko, *in situ* temperature, salinity, pH and alkalinity were $27.1 \pm 0.2^{\circ}$ C, 33.5 PSU, 8.18 ± 0.03 , and 2.28 ± 0.01 mmol kg⁻¹, respectively.

Nutrient Fluxes

Nutrients were generally taken up during incubation in both light and dark periods, with the exception of ammonium and phosphate (Figure 2), and the rate of uptake differed significantly depending on habitat (ANOVA, P < 0.01; Table 2). The seagrass and mangrove samples had comparatively high amounts of silicate initially (Tukey's test, P < 0.05; Supplementary Table 3). Though it was not statistically significant (Tukey's test, P > 0.05), nitrate + nitrite and silicate uptake increased under the HTHP stress during both light and dark periods, especially in mangrove and seagrass samples (Figure 2). The highest nitrate concentration was observed in the seagrass habitat (Tukey's test, P < 0.05; Supplementary Table 3), but the rate of uptake in seagrass did not differ significantly among incubation conditions (Tukey's test, P > 0.05). In lagoonal sediment, the rate of nitrate + nitrite uptake seemed to increase under HTAP and HTHP conditions during light. Additionally, this sediment seemed to release ammonium under ATHP and HTHP conditions during both light and dark periods.





Photosynthetic Community

Based on tracer pigments and microscopic observations, a well-developed benthic microalgal community was detected in all habitats. The photosynthetic community was dominated by pennate diatoms and filamentous cyanobacteria. Centric diatoms (in mangrove and seagrass sediments), euglenoids (in seagrass and coral reef lagoon sediments), and flagellates (in all sediments) were occasionally observed. Nano-sized (5–20 μ m) benthic diatoms (Supplementary Table 4), especially naviculoids, were frequently found in all sediment types. Additionally, Bacillariophyceae (in all sediments) and Coscinodiscophyceae

(in seagrass and coral reef sediments) were identified. Pico-sized $(<2 \,\mu\text{m})$ diatoms were abundant in mangrove and seagrass sediments. A higher abundance of photosynthetic community taxa was evidenced by a higher concentration of chlorophyll *a*, fucoxanthin, and zeaxanthin in mangrove sediment (Tukey's test, P < 0.05; Figure 3, Table 2, Supplementary Figure 2) than in seagrass and lagoonal sediments. Under HTHP stress, chlorophyll *a* increased in the surface sediment, representing the bulk of the photosynthetic community in mangrove and seagrass incubations (Figure 4). Changes in the concentrations of other pigments under HTHP stress also indicated an increase

TABLE 2 The effects of different habitats (H) and incubation conditions (IC) on environmental parameters based on two-	way ANOVA.
--	------------

Variables	Significant factor	MS	df	F	P-value
$\Delta NO_3^- + NO_2^-$	Н	126.85	2	293.48	<0.001
ΔNH_4^+	Н	2.52	2	38.77	< 0.001
ΔPO_4^{3-}	Н	0.55	2	7.371	<0.01
$\Delta Si(OH)_4$	Н	172.12	2	214.35	<0.01
	IC	10.02	3	12.45	< 0.001
	$H \times IC$	2.53	6	3.15	<0.05
Δ Chlorophyll <i>a</i>	Н	17995.30×10^3	2	703	<0.001
	IC	31936.12×10^3	3	1248	< 0.001
	$H \times IC$	6990.26 × 10 ³	6	273.09	<0.01
Δ Fucoxanthin	Н	3289.32×10^3	2	8004	< 0.001
	IC	741.80 × 10 ³	3	1804	< 0.001
	$H \times IC$	790.40 × 10 ³	6	1923.01	< 0.001
Δ Zeaxanthin	Н	197.02 × 10 ³	2	3107	< 0.001
	IC	76.56×10^{3}	3	1202	<0.01
	$H \times IC$	24.32×10^{3}	4	383.01	< 0.001
Δ β -Carotene	Н	93.12 × 10 ³	2	974	<0.001
	IC	36.43×10^{3}	3	382	<0.001
	$H \times IC$	17.74 × 10 ³	6	183.12	<0.01
Δ Chlorophyll b	Н	1195.01× 10 ³	2	18838	< 0.001
	IC	241.61× 10 ³	3	3812	< 0.001
	$H \times IC$	151.91× 10 ³	6	2381.02	<0.01
Δ Chlorophyll c_2	Н	7300.21× 10 ³	2	54472	< 0.001
	IC	2278.00 × 10 ³	3	17000	<0.01
	$H \times IC$	1802.01 × 10 ³	6	13448.01	< 0.001
NPP	Н	18.53	2	310	<0.01
	IC	14.64	4	245	<0.01
	$H \times IC$	4.49	4	75.29	<0.01
GPP	Н	18.52	2	5.527	<0.05
R	Н	1.98	2	8.53	<0.01
	IC	5.75	4	24.56	<0.01
	H × IC	1.94	3	8.361	<0.01
Δ Bacterial abundance in sediment	Н	62.78×10^{6}	2	127.82	<0.01
	IC	165.18 × 10 ⁶	3	340.81	<0.001
	H × IC	14.60×10^{6}	6	28.96	< 0.001
Δ Bacterial abundance in seawater	Н	10.27 × 10 ⁵	2	7.95	< 0.01
	IC	282.69×10^5	3	221.66	< 0.001
	H × IC	16.27×10^5	6	12.667	<0.001

MS, mean square; df, degrees of freedom; F, F-ratio. H represents mangrove, seagrass and coral reef lagoon habitats and IC represents four incubation conditions; ATAP, Ambient temperature and ambient pCO₂; ATHP, Ambient temperature and high pCO₂; HTAP, High temperature and ambient pCO₂; and HTHP, High temperature and high pCO₂.

in the photosynthetic community; however, in the reef lagoon, some degradation of pigments was observed (**Figure 3**). The highest concentration of chlorophyll *a* was found in incubated mangrove sediment (Tukey's test, P < 0.05; **Figure 3**). The chlorophyll *a* concentration in the surface sediments increased after HTHP treatment in all habitats, and this change was greater in mangrove sediments and lower in coral reef lagoon sediments. Under HTHP treatment chlorophyll *a* concentration increased 1550–2400, 750–1455, and 620–860 ng g⁻¹ in mangrove, seagrass, and coral reef lagoon sediments, respectively (**Figure 4**).

Sediment Primary Production and Respiration

Oxygen fluxes increased significantly under HTHP stress during light periods in mangrove and seagrass habitats (Tukey's test, P < 0.001; **Figure 5**, Supplementary Table 5). In coral reef lagoons, stress increased oxygen fluxes. NPP and R varied depending on incubation conditions and habitat (ANOVA, P < 0.01; **Table 2**). NPP was comparatively higher under HTHP stress in mangrove (Tukey's test, P < 0.001) and seagrass (Tukey's test, P < 0.01) incubations than in the coral reef lagoon (**Figure 6**). Additionally, GPP and NPP in seagrass incubations were lower



differences at P < 0.05 and P < 0.01 determined using least significant differences at P < 0.05 and P < 0.01 determined using least significant difference tests (*post-hoc* Tukey's test) depending on incubation conditions (ATAP, ATHP, HTAP, and HTHP) and habitat (mangrove, seagrass, and coral reef lagoon). Lack of significance is indicated by ns. ATAP, ambient temperature and ambient pCO_2 ; ATHP, ambient temperature and high pCO_2 ; HTAP, high temperature and ambient pCO_2 ; and HTHP, high temperature and high pCO_2 .

than in mangroves (Tukey's test, P < 0.01), but higher than in coral reef lagoon incubations. Neither high temperature nor high pCO_2 alone had a significant impact on primary production (Tukey's test, P > 0.05; **Figure 6**), but temperature had a slightly greater influence. Respiration did not differ significantly under ATHP, HTAP, or HTHP treatment (Tukey's test, P > 0.05), with the exception of the coral reef lagoon (Tukey's test, P < 0.05; **Table 2, Figure 6**). Lagoonal sediment showed higher respiration (Tukey's test, P < 0.05) in the ATAP treatment, but it decreased under single or combined stresses, and this was reflected in the increased primary production.

Bacterial Abundance

Initial bacterial abundance in sediment was highest (Tukey's test, P < 0.05) in seagrass (19 ± 0.12 × 106 cells cm⁻² in mangrove

samples, $21 \pm 1.1 \times 106$ cells cm⁻² in seagrass samples, and 7 \pm 1.0 \times 106 cells cm⁻² in coral reef sediments). Additionally, bacterial abundance in mangrove seawater was higher than other seawater samples $(22 \pm 2.0 \times 10^5 \text{ cells mL}^{-1} \text{ in the mangrove}, 21$ \pm 1.1 \times 10⁵ cells mL⁻² in the seagrass, and 5 \pm 0.5 \times 10⁵ cells mL^{-1} in the coral reef lagoon). Bacterial abundance in sediment and seawater increased significantly by HTHP stress(Tukey's test, P < 0.01, Figures 7, 8), except in seagrass sediments. The impacts of ATHP and HTAP conditions were not significant in the cases of seagrass sediment and seawater, and coral reef lagoon seawater (Tukey's test, P > 0.05, Figure 8). In mangrove seawater and sediments, the impacts of HTAP stress were significantly greater than the impact of ATHP stress (Tukey's test, P < 0.01). Bacterial abundance in the surficial sediments increased after HTHP stress in all habitats, and this change was greater in mangrove sediments and lower in coral reef lagoon sediments. Under HTHP stress bacterial abundance increased 14×10^5 - 23×10^5 cells g^{-1} , $8\times10^5\text{--}16\times10^5$ cells g^{-1} , and $5\times10^5\text{--}8\times10^5$ cells g^{-1} in mangrove, seagrass, and coral reef lagoon sediments, respectively (Figure 7). In mangrove sediment and seawater, growth rate and doublings of bacteria increased significantly under HTHP stress (Tukey's test, P > 0.05, Supplementary Table 6). Similarly, in seagrass and coral reef lagoon seawater, bacterial growth rate and doublings were significantly impacted by HTHP stress (Tukey's test, P > 0.05, Supplementary Table 6).

DISCUSSION

The responses of subtropical shallow marine sediment systems to elevated temperature and pCO_2 varied depending on the habitats from which the sediments were collected. In this study, we used sediments collected with cores to test the responses of relatively undisturbed sediments. Changes in primary production in response to a temperature increase of 4° C and 936 ppm pCO₂ were detected. The following general responses were observed in this study. (1) Under HTHP stress, sediment primary production rates were higher in mangrove and seagrass environments than in the coral reef lagoon. (2) Nutrients flux was lower in the coral reef lagoon. Silicate and nitrate + nitrite were noticeably consumed in the mangrove and seagrass. (3) Bacterial abundance in sediment and seawater generally increased under ATHP, HTAP, and HTHP. Growth rate and doublings indicated increased the activity of bacteria in seawater (mangrove, seagrass, and coral reef lagoon) and sediment (mangrove) under HTHP conditions.

Effect of Stressors on Primary Production and Bacterial Abundance

Previous studies have suggested that the combination of warming and OA positively stimulate benthic microalgal primary production in seagrass habitats, and benthic microalgae appear to be resilient to future environmental changes (Alsterberg et al., 2013). Our incubation experiments indicated increased NPP and GPP at the sediment–water interface under HTHP stress, with the exception of the coral reef lagoon. OA may cause stimulated diatom activity in benthic assemblages (Johnson et al., 2013). The impacts of temperature (HTAP) on NPP, GPP, and



FIGURE 4 | **Changes in surficial chlorophyll a concentration in sediment.** Initial (**A**) and after incubation in HTHP (high temperature and high pCO_2) condition (**B**). The graph was derived from r (radius of the incubation chamber), θ (angles of sediment samples collection), and z (chlorophyll *a* concentration) data. Chlorophyll *a* concentrations in the incubation chambers' sediment surface are shown with colors. Corresponding chlorophyll *a* value for each color is presented on the color scale. Colors from purple to red represent lower to higher chlorophyll *a* concentrations.



FIGURE 5 | Sediment-water fluxes of oxygen during light (A) and dark (B) periods. *, ** Indicate significant differences at P < 0.01 and P < 0.001 based on the least significant difference test (*post-hoc* Tukey's test) depending on habitats (mangrove, seagrass, and coral reef lagoon). Lack of significance is indicated by ns. ATAP, ambient temperature and ambient pCO_2 ; ATHP, ambient temperature and high pCO_2 ; HTAP, high temperature and ambient pCO_2 ; and HTHP, high temperature and high pCO_2 .

R were slightly, but not significantly higher than those of pCO_2 (ATHP). Significant differences in NPP (Tukey's test, P > 0.01) and GPP (Tukey's test, P > 0.05) under the HTHP stress

suggested that benthic microalgae in mangrove and seagrass sediments are more responsive to HTHP stress than those in coral reef lagoon sediments. Previously, it has been suggested



(NPP), and respiration (R) in mangrove, seagrass, and coral reef sediment. *, **, and *** Indicate significant differences at P < 0.05, P < 0.01, and P < 0.001 determined using least significant difference tests (*post-hoc* Tukey's test) depending on incubation conditions (ATAP, ATHP, HTAP, and HTHP). Lack of significance is indicated by ns. ATAP, ambient temperature and ambient ρ CO₂, ATHP, ambient temperature and high ρ CO₂, HTAP, high temperature and ambient ρ CO₂, and HTHP, high temperature and high ρ CO₂.

that microphytobenthic assemblages, including cyanobacteria, phototrophic bacteria, and microalgal mats, increase rapidly in shallow marine environments, such as intertidal flats, coastal embayments, and lagoons (MacIntyre et al., 1996). GPP in tropical mangrove sediments was much higher during the warm season owing to the greater seasonal temperature range (Alongi, 1994). Moreover, phototrophic microphytobenthos are important primary producers, contributing almost half of the primary production in estuarine ecosystems (Underwood and Kromkamp, 1999). Our results suggest that only HTHP stress significantly stimulated sediment primary production in mangrove and seagrass sediments. In contrast, sediment primary production in the coral reef lagoon was stimulated by stress, but the result was not statistically significant; however, decreased R and dark oxygen flux under ATHP, HTAP, and HTHP treatments in lagoonal incubations caused a shift from heterotrophy to autotrophy. Nonsignificant changes in respiration in different habitats under HTHP stress suggest that heterotrophic processes were less affected by HTHP stress than autotrophic processes. Similar results were observed by Alsterberg et al. (2012) for oxygen flux in seagrass sediment under elevated temperature. Furthermore, Johnson et al. (2013) suggested that future changes in CO_2 and bicarbonate (HCO_3^-) availability could stimulate photosynthesis in marine autotrophs.

Previously, it has been suggested that microphytobenthic productivity is correlated with chlorophyll a content (MacIntyre et al., 1996). Based on our primary production data, photosynthetic community activity was stimulated by high temperatures and high pCO_2 . However, not only high temperature (Alsterberg et al., 2012), but also high pCO_2 significantly increased the chlorophyll a content at the sediment surface. These increases are associated with increased NPP, indicating that HTHP conditions stimulate primary production activity by the photosynthetic community in all habitats. Tracer pigment composition and microscopic observations indicated dominance of diatoms in all sediment types. This result was consistent with those of Hendrarto and Nitisuparjo (2011). CO₂ enrichment can cause significant increases in chlorophyll a concentrations and in diatom abundance (Johnson et al., 2013). In our incubation experiments, significant increases in pigment concentrations (chlorophyll a, fucoxanthin, and zeaxanthin) were observed, except in the coral reef lagoon. These results revealed that photosynthetic communities inhabiting mangroves and seagrasses more strongly fuel sediment primary production compared to those of coral reef lagoons. These patterns indicate that new organic matter was produced during incubation. The highest nutrient consumption (e.g., ammonium, phosphate, and silicate) and the greatest increase in bacterial abundance were observed in mangrove sediments (Figure 8); these results are supported by the increased nutrient uptake during dark periods in these sediments. Increased bacterial abundance under HTHP conditions may be a direct effect of warming (Sander and Kalff, 1993). However, increased bacterial abundance could be partially dependent on bottom-up effects, since bacteria are considered to use dissolved organic matter produced by benthic microalgae (Evrard et al., 2008; Casareto et al., 2012).

Nutrient Dynamics

Microphytobenthos and mat-forming macroalgae have different direct and indirect effects on nitrogen-associated processes via consumption and the inhibition or stimulation of microbial activity (Bartoli et al., 2012). The rate of nutrient uptake, particularly of silicate, ammonium (not in the coral reef lagoon), and nitrate + nitrite, seemed to be released under HTHP conditions. Higher uptake of silicate in seagrass and mangrove incubations (**Figure 2**) was associated with a higher initial concentration (Supplementary Table 3). A previous study suggested that a relatively higher abundance of silicate compared to nitrate and phosphate occurs in fresh



graph was derived from r (radius of the incubation chamber), θ (angles of sediment samples collection), and z (bacterial abundance) data. Bacterial abundances in the incubation chambers' sediment surface are shown. Corresponding bacterial abundance value for each color is presented in color scale. Purple to orange colors represent relatively low to high bacterial abundance.





water-influenced coastal zones (Cloern et al., 1985). However, the consumption of silicate reflected the predominant activity of benthic diatoms. This finding is consistent with those of Alsterberg et al. (2011) and Alsterberg and Sundbäck (2013). Diatoms benefit from increasing CO₂ via a reduction in the energetic costs of their carbon-concentrating mechanism (Hopkinson et al., 2011). The tracer pigment data indicated an increased abundance of fucoxanthin (Figure 3) and zeaxanthin, indicating stimulated activity of diatoms and cyanobacteria in mangrove and seagrass incubations. In coral reef incubations, the activity of cyanobacteria increased as evidenced by an increase in the zeaxanthin concentration; slightly higher uptake of silicate under ATHP, HTAP, and HTHP treatments indicated stimulated activity of diatoms in lagoonal incubations. The release of ammonium in lagoonal incubations under ATHP and HTHP treatments indicated increased nitrogen mineralization, consistent with the results of Alsterberg et al. (2012). Initially higher concentrations of nutrients in pore water than that of overlying seawater implied that biological processes could be more important than physical diffusion in determining nutrient fluxes. Therefore, microbes may play an important role consuming nutrients (Arístegui et al., 2009; Miyajima, 2015) in all habitats, and an increase in bacterial abundance could suggest increased nutrient uptake (Ye et al., 2012).

Tentative Biogeochemical Implications

Our results suggest that HTHP stress has a major impact on primary production, especially on carbon cycling. Sediments from mangrove and seagrass habitats are more productive than those from coral reef lagoons, suggesting that they act as potential carbon sinks. Significantly increased uptake of nutrients (particularly silicate) in mangrove and seagrass sediments may enhance the role of sediments as nutrient sinks. Oxygen fluxes and increased bacterial abundance suggest that microbial community activity increased in these sediments, but did not exceed that of autotrophic compartments. In mangroves, community respiration may increase slightly, but not enough to increase heterotrophy owing to the capacity of primary producers to photosynthesize under stressful conditions. Significantly enhanced sediment primary production in shallow marine mangroves and seagrass sediments compared to coral reef lagoon sediments provided evidence for positive responses to HTHP stress with concomitant carbon storage.

CONCLUSIONS

The combined impact of increased temperature and pCO_2 associated with climate change differs significantly from the impact of individual stressors, and can result in nonlinear effects and unexpected ecological patterns (Segner et al., 2014). To our knowledge, this is the first demonstration of differences in subtropical sediment systems between mangrove, seagrass, and coral reef lagoon habitats of Okinawa, Japan under high temperature and OA. A synergistic increase in sediment primary production was observed under the combination of stressors

(HTHP), except in the coral reef lagoon. Elevated temperature and pCO_2 stimulated primary production, autotrophic activity, bacterial abundance, and nutrient uptake. Greater carbon fixation in mangrove and seagrass sediments suggests that they have increased resilience compared to lagoonal sediment. However, lagoonal sediment also responded positively via a shift from heterotrophy to autotrophy under single (ATHP, HTAP) or combined stressors (HTHP). These results provide a basis for further studies to understand the responses of different shallow marine habitats to environmental stress. However, further efforts are needed to assess whether the roles proposed for mangroves, seagrass, and coral reef lagoon habitats within this study can be applied to a broader spatiotemporal scale.

AUTHOR CONTRIBUTIONS

Conceived and designed the experiments: RS, BC, HF, and YS. Performed the experiments: RS, BC, TS, and YS. Analyzed the data: RS, BC, RS, TS and MA. Contributed reagents/materials/analysis tools: RS, BC, RS, TS, MA, HF, and YS. Wrote the paper: RS, BC, RS, TS, MA, HF, and YS.

FUNDING

This study was supported by the Environmental Leaders Program of Shizuoka University (ELSU), a Grant-in-aid for Scientific Research on Innovative Areas "Coral reef science for symbiosis and coexistence of humans and ecosystems under combined stresses" (No. 20121003) of the Ministry of Education. Culture, Sports, Science and Technology (MEXT), Japan and the Global Coral Reef Conservation Project (GCRCP) of the Mitsubishi Corporation.

ACKNOWLEDGMENTS

The authors are grateful for the support of the staff at Sesoko Station, Tropical Biosphere Research Center, University of the Ryukyus, Japan for providing necessary facilities during this experiment. The authors are also thankful to the members of Marine Ecosystem and Biogeochemistry Laboratory, Shizuoka University, Japan. We are grateful for the help of Dr. Tomihiko Higuchi, The University of Tokyo, Japan during sample analysis. We thank Dr. Mohan Prasad Niraula, Tribhuvan University, Nepal for providing constructive comments on earlier versions of this manuscript. The authors would like to thank Mr. S. Uehara for his help with fieldwork and sample collection. Special thanks to Mr. K. Toyoda who helped with benthic microalgal identification. We thank two anonymous reviewers for constructive comments.

SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: http://journal.frontiersin.org/article/10.3389/fmars. 2016.00100

REFERENCES

- Alongi, D. M. (1994). Zonation and seasonality of benthic primary production and community respiration in tropical mangrove forests. *Oecologia* 98, 320–327. doi: 10.1007/BF00324220
- Alsterberg, C., Eklöf, J. S., Gamfeldt, L., Havenhand, J. N., and Sundbäck, K. (2013). Consumers mediate the effects of experimental ocean acidification and warming on primary producers. *Proc. Natl. Acad. Sci. U.S.A.* 110, 8603–8608. doi: 10.1073/pnas.1303797110
- Alsterberg, C., Hulth, S., and Sundback, K. (2011). Response of a shallowwater sediment system to warming. *Limnol. Oceanogr.* 56, 2147–2160. doi: 10.4319/lo.2011.56.6.2147
- Alsterberg, C., and Sundbäck, K. (2013). Experimental warming and toxicant exposure can result in antagonistic effects in a shallow-water sediment system. *Mar. Ecol. Prog. Ser.* 488, 89–101. doi: 10.3354/meps10357
- Alsterberg, C., Sundbäck, K., and Hulth, S. (2012). Functioning of a shallow-water sediment system during experimental warming and nutrient enrichment. *PLoS ONE* 7:e51503. doi: 10.1371/journal.pone.0051503
- Arístegui, J., Mendonc, A., Vilas, J. C., Espino, M., Polo, I., Montero, M. F., et al. (2009). Plankton metabolic balance at two North Atlantic seamounts. *Deep Sea Res. II* 56, 2646–2655. doi: 10.1016/j.dsr2.2008. 12.025
- Aswani, S., Mumby, P., Baker, A. C., Christie, P., McCook, L. J., Steneck, R. S., et al. (2015). Scientific frontiers in the management of coral reefs. *Front. Mar. Sci.* 2:50. doi: 10.3389/fmars.2015.00050
- Barbier, E. B., Hacker, S. D., Kennedy, C., Koch, E. W., Stier, A. C., and Silliman, B. R. (2011). The value of estuarine and coastal ecosystem services. *Ecol. Monogr.* 81, 169–193. doi: 10.1890/10-1510.1
- Bartoli, M., Castaldelli, G., Nizzoli, D., and Viaroli, P. (2012). Benthic primary production and bacterial denitrification in a Mediterranean eutrophic coastal lagoon. J. Exp. Mar. Biol. Ecol. 438, 41–51. doi: 10.1016/j.jembe.2012. 09.011
- Beck, M. W., Heck, K. L., Able, K. W., Childers, D. L., Eggleston, D. B., Gillanders,
 B. M., et al. (2001). The identification, conservation, and management of estuarine and marine nurseries for fish and invertebrates. *Bioscience* 51, 633. doi: 10.1641/0006-3568(2001)051[0633:TICAMO]2.0.CO;2
- Bouillon, S., Borges, A. V., Castaneda-Moya, E., Diele, K., Dittmar, T., Duke, N. C., et al. (2008). Mangrove production and carbon sinks: a revision of global budget estimates. *Glob. Biogeochem. Cycles* 22, 12. doi: 10.1029/2007GB003052
- Brito, A., Newton, A., Tett, P., and Fernandes, T. F. (2010). Sediment and water nutrients and microalgae in a coastal shallow lagoon, Ria Formosa (Portugal): implications for the Water Framework Directive. J. Environ. Monit. 12, 318–328. doi: 10.1039/B909429F
- Buchanan J. B. (1984). "Sediment Analysis," in *Methods for the Study of Marine Benthos*, eds. N. A. Holme, and A. D. McIntyre (London: Blackwell Scientific Publications), 41–65.
- Casareto, B. E., Niraula, M. P., Fujimura, H., and Suzuki, Y. (2009). Effects of carbon dioxide on the coccolithophorid *Pleurochrysis carterae* in incubation experiments. *Aquat. Biol.* 7, 59–70. doi: 10.3354/ab00182
- Casareto, B. E., Niraula, M. P., and Suzuki, Y. (2012). Dynamics of organic carbon under different inorganic nitrogen levels and phytoplankton composition. *Estuar. Coast. Shelf Sci.* 102-103, 84–94. doi: 10.1016/j.ecss.2012.03.019
- Chihara, M., and Murano, M. (1997). An Illustrated Guide to Marine Plankton in Japan. Tokyo: Tokai University Press.
- Chmura, G. L. (2013). What do we need to assess the sustainability of the tidal salt marsh carbon sink? *Ocean Coast. Manage.* 83, 25–31. doi: 10.1016/j.ocecoaman.2011.09.006
- Cloern, J. E., Cole, B. E., Wong, R. L. J., and Alpine, A. E. (1985). Temporal dynamics of estuarine phytoplankton: a case study of San Francisco Bay. *Hydrobiologia* 129, 153–176. doi: 10.1007/BF00048693
- Connell, S. D., Kroeker, K. J., Fabricius, K. E., Kline, D. I., and Russell, B. D. (2013). The other ocean acidification problem: CO₂ as a resource among competitors for ecosystem dominance. *Proc. R. Soc. B Biol. Sci.* 368, 1–9. doi: 10.1098/rstb.2012.0442
- Cornwell, J. C., Glibert, P. M., and Owens, M. S. (2014). Nutrient fluxes from sediments in the San Francisco Bay delta. *Estuar. Coast* 37, 1120–1133.doi: 10.1007/s12237-013-9755-4

- Cullen-Unsworth, L., and Unsworth, R. (2013). Seagrass meadows, ecosystem services, and sustainability. *Environ. Sci. Policy Sus. Dev.* 55, 14–28. doi: 10.1080/00139157.2013.785864
- Del Giorgio, P., and Williams, P. (2005). *Respiration in Aquatic Ecosystems*. New York, NY: Oxford University Press.
- Delille, B., Harlay, J., Zondervan, I., Jacquet, S., Chou, L., Wollast, R., et al. (2005). Response of primary production and calcification to changes of pCO_2 during experimental blooms of the coccolithophorid *Emiliania huxleyi. Glob. Biogeochem. Cycles* 19, GB2023. doi: 10.1029/2004GB002318
- Dickson, A. G., Sabine, C. L., and Christian, J. R. (2007). Guide to Best Practices for Ocean CO₂ Measurements. PICES special publication 3, IOCCP Report No.8.
- Donato, D. C., Kauffman, J. B., Mackenzie, R. A., Ainsworth, A., and Pfleeger, A. Z. (2012). Whole-island carbon stocks in the tropical Pacific: implications for mangrove conservation and upland restoration. *J. Env. Manage*. 97, 89–96. doi: 10.1016/j.jenvman.2011.12.004
- Duarte, C. M., Marbà, N., Gacia, E., Fourqurean, J. W., Beggins, J., Barrón, C., et al. (2010). Seagrass community metabolism: assessing the carbon sink capacity of seagrass meadows. *Glob. Biogeochem. Cycles* 24, 1–8. doi: 10.1029/2010GB003793
- Duarte, C. M., Middelburg, J. J., and Caraco, N. (2005). Major role of marine vegetation on the oceanic carbon cycle. *Biogeosciences* 2, 1–8. doi: 10.5194/bg-2-1-2005
- Engel, A., Schulz, K., Riebesell, U., Bellerby, R., Delille, B., and Schartau, M. (2007). Effects of CO₂ on particle size distribution and phytoplankton abundance during a mesocosm bloom experiment (PeECE II). *Biogeosciences* 5, 509–521. doi: 10.5194/bg-5-509-2008
- Evrard, V., Cook, P. L. M., Veuger, B., Huettel, M., and Middelburg, J. J. (2008). Tracing carbon and nitrogen incorporation and pathways in the microbial community of a photic subtidal sand. *Aquat. Microb. Ecol.* 53, 257–269. doi: 10.3354/ame01248
- Folt, C. L., Chen, C. Y., Moore, M. V., and Burnaford, J. (1999). Synergism and antagonism among multiple stressors. *Limnol. Oceanogr.* 44, 864–877. doi: 10.4319/lo.1999.44.3_part_2.0864
- Gao, K., Aruga, Y., Asada, K., and Kiyohara, M. (1993). Influence of enhanced CO₂ on growth and photosynthesis of the red algae *Gracilaria* sp. and G. *chilensis. J. Appl. Phycol.* 5, 563–571. doi: 10.1007/BF02184635
- Hansen, H. P., and Koroleff, F. (2007). "Determination of nutrients," in *Methods of Seawater Analysis*, eds K. Grasshoff, K. Kremling, and M. Ehrhardt (Weinheim: Wiley-VCH Verlag GmbH), 159–228.
- Hendrarto, I. B., and Nitisuparjo, M. (2011). Biodiversity of benthic diatom and primary productivity of benthic micro-flora in mangrove forests on central Java. J. Coast. Dev. 14, 131–140. Available online at: http://www.omicsonline. com/open-access/biodiversity-of-benthic-diatom-and-primary-productivityof-benthic-microflora-in-mangrove-forests-on-central-java-1410-5217-14-309.pdf
- Hopkinson, B. M., Dupont, C. L., Allen, A. E., and Morel, F. M. M. (2011). Efficiency of the CO₂-concentrating mechanism of diatoms. *Proc. Nat. Acad. Sci. U.S.A.* 108, 3830–3837. doi: 10.1073/pnas.1018062108
- Intergovernmental Panel on Climate Change (IPCC) (2013). Climate Change 2013 The Physical Science Basis Working Group I Contribution to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, 2013.
- Jennerjahn, T. C., and Ittekkot, V. (2002). Relevance of mangroves for the production and deposition of organic matter along tropical continental margins. *Naturwissenschaften* 89, 23–30. doi: 10.1007/s00114-001-0283-x
- Johnson, V. R., Brownlee, C., Rickaby, R. E. M., Graziano, M., and Milazzo, M. (2013). Responses of marine benthic microalgae to elevated CO₂. *Mar. Biogl.* 160, 1813–1824. doi: 10.1007/s00227-011-1840-2
- Kayanne, H., Hata, H., Kudo, S., Yamano, H., Watanabe, A., Ikeda, Y., et al. (2005). Seasonal and bleaching-induced changes in coral reef metabolism and CO₂ flux. *Glob. Biogeochem. Cycles* 19, GB3015. doi: 10.1029/2004GB002400
- Kennedy, H., Beggins, J., Duarte, C. M., Fourqurean, J. W., Holmer, M., Marbá, N., et al. (2010). Seagrass sediments as a global carbon sink: isotopic constraints. *Glob. Biogeochem. Cycles* 24, 1–8. doi: 10.1029/2010GB003848
- Lewis, E., and Wallace, D. W. R. (1998). "Program developed for CO₂ system calculations. ORNL/CDIAC-105," in *Carbon Dioxide Information Analysis Center* (Oak Ridge, TN: Oak Ridge National Laboratory, U.S. Department of Energy).

- Libes, S. (2009). "Organic matter: Production and destruction," in *Introduction to Marine Biogeochemistry*, ed S. Libes (Elsevier Inc.), 207–220.
- MacIntyre, H. L., Geider, R. J., and Miller, D. C. (1996). Microphytobenthos: the ecological role of the "secret garden" of unvegetated, shallow-water marine habitats. I. Distribution, abundance and primary production. *Estuaries* 19, 186–201. doi: 10.2307/1352224
- Mcleod, E., Chmura, G. L., Bouillon, S., Salm, R., Björk, M., Duarte, C. M., et al. (2011). A blueprint for blue carbon: toward an improved understanding of the role of vegetated coastal habitats in sequestering CO₂. *Front. Ecol. Environ.* 9:4. doi: 10.1890/110004
- Mehrbach, C., Culberso, C. H., Hawley, J. E., and Pytkowic, R. M. (1973). Measurement of apparent dissociation constants of carbonic acid in seawater at atmospheric pressure. *Limnol. Oceanogr.* 18, 897–907. doi: 10.4319/lo.1973.18.6.0897
- Millero, F. J. (1979). The thermodynamics of the carbonate system in seawater. Geochim. Cosmochim. Acta 43, 1651–1661. doi: 10.1016/0016-7037(79)90184-4
- Miyajima, T. (2015). Abiotic versus biotic immobilization of inorganic nitrogen in sediment as a potential pathway of nitrogen sequestration from coastal marine ecosystems. *Geochem. J.* 49, 453–468. doi: 10.2343/geochemj.2.0370
- Miyajima, T., Hori, M., Hamaguchi, M., Shimabukuro, H., Adachi, H., Yamano, H., et al. (2015). Geographic variability in organic carbon stock and accumulation rate in sediments of East and Southeast Asian seagrass meadows. *Glob. Biogeochem. Cycles* 29, 397–415. doi: 10.1002/2014GB004979
- Montoya, J. M., and Raffaelli, D. (2010). Climate change, biotic interactions and ecosystem services. *Phil. Trans. R. Soc. Lond. B* 365, 2013–2018. doi: 10.1098/rstb.2010.0114
- Mumby, P. J. (2006). Connectivity of reef fish between mangroves and coral reefs: algorithms for the design of marine reserves at seascape scales. *Biol. Cons.* 128, 215–222. doi: 10.1016/j.biocon.2005.09.042
- Nakano, Y., and Nakai, T. (2008). Succeeding to consider reef and humans: a comprehensive view of the enlightenment for coral reef conservation. *J. Jpn. Coral Reef Soc.* 10, 105–115. doi: 10.3755/jcrs.10.105
- Pendleton, L., Donato, D. C., Murray, B. C., Crooks, S., Jenkins, W. A., Sifleet, S., et al. (2012). Estimating global "blue carbon" emissions from conversion and degradation of vegetated coastal ecosystems. *PLoS ONE* 7:e43542. doi: 10.1371/journal.pone.0043542
- Porter, K. G., and Feig, Y. S. (1980). The use of DAPI for identifying and counting aquatic microflora. *Limnol. Oceanogr.* 25, 943–948. doi: 10.4319/lo.1980.25.5.0943
- Qiu, B., and Gao, K. (2002). Effects of CO₂ enrichment on the bloomforming cyanobacterium *Microcystis aeruginosa* (Cyanophyceae): physiological responses and relationships with the availability of dissolved inorganic carbon. *J. Phycol.* 38, 721–729. doi: 10.1046/j.1529-8817.2002.01180.x
- Ricart, A. M., York, P. H., Rasheed, M. A., Pérez, M., Romero, J., Bryant, C. V., et al. (2015). Variability of sedimentary organic carbon in patchy seagrass landscapes. *Marine Poll. Bull.* 100, 476–482. doi: 10.1016/j.marpolbul.2015.09.032
- Riebesell, U., Zondervan, I., Rost, B., Tortell, P. D., Zeebe, R. E., and Morel, F. M. (2000). Reduced calcification of marine plankton in response to increased atmospheric CO₂. *Nature* 407, 364–367. doi: 10.1038/35030078
- Robertson, A. I., and Alongi, D. M. (2013). *Tropical Mangrove Ecosystems*. Washington, DC: American Geophysical Union.
- Sander, B. B., and Kalff, J. (1993). Factors controlling bacterial production in marine and freshwater sediments. *Microb. Ecol.* 26, 79-99. doi: 10.1007/BF00177045
- Segner, H., Schmitt-Jansen, M., and Sabater, S. (2014). Assessing the impact of multiple stressors on aquatic biota: the receptor's side matters. *Environ. Sci. Technol.* 48, 7690–7696. doi: 10.1021/es405082t

- Sifleet, S., Pendleton, L., and Murray, B. C. (2011). *State of the Science on Coastal Blue Carbon a Summary for Policy Makers*. (Durham, NC: Nicholas Institute for Environmental Policy Solutions).
- Stanier, R. Y., Adelverg, E. A., Ingraham, J. L., and Wheelis, M. L. (1979). "Microbial growth," in *Introduction to the Microbial World*, (Englewood Cliffs, NJ: Prentice-Hall).
- Stumm, W., and Morgan, J. (1981). Aquatic Chemistry: an Introduction Emphasizing Chemical Equilibria in Natural Waters, 2nd Edn. New York, NY: Wiley Interscience.
- Sundback, K., Alsterberg, C., and Larson, F. (2010). Effects of multiple stressors on marine shallow-water sediments: response of microalgae and meiofauna to nutrient-toxicant exposure. J. Exp. Mar. Biol. Ecol. 388, 39–50. doi: 10.1016/j.jembe.2010.03.007
- Suzuki, T., Casareto, B. E., Shioi, Y., Ishikawa, Y., and Suzuki1, Y. (2015). Finding of 13², 17³-cyclopheophorbide a enol as a degradation product of chlorophyll in shrunk zooxanthellae of the coral *Montipora digitata. J. Phycol.* 51, 37–45. doi: 10.1111/jpy.12253
- Underwood, G. J. C., and Kromkamp, J. (1999). Primary production by phytoplankton and microphytobenthos in estuaries. *Adv. Ecol. Res.* 29, 93–153. doi: 10.1016/S0065-2504(08)60192-0
- Weiss, R. (1974). Carbon dioxide in water and seawater: the solubility of a non-ideal gas. *Mar. Chem.* 2, 203–215. doi: 10.1016/0304-4203(74) 90015-2
- Winder, M., and Schindler, D. E. (2004). Climate change uncouples trophic interactions in an aquatic ecosystem. *Ecology* 85, 2100–2106. doi: 10.1890/ 04-0151
- Woodward, G., Perkins, D. M., and Brown, L. E. (2010). Climate change and freshwater ecosystems: impacts across multiple levels of organization. *Phil. Trans. R. Soc. Lond. B* 365, 2093–2106. doi: 10.1098/rstb.20 10.0055
- Ye, L., Shi, X., Wu, X., and Kong, F. (2012). Nitrate limitation and accumulation of dissolved organic carbon during a spring-summer cyanobacterial bloom in Lake Taihu (China). J. Limnol. 71, 67–71. doi: 10.4081/jlimnol. 2012.e6
- Zapata, M., Rodríguez, F., and Garrido, J. L. (2000). Separation of chlorophylls and carotenoids from marine phytoplankton: a new HPLC method using a reversed phase C8 column and pyridine-containing mobile phases. *Mar. Ecol. Prog. Ser.* 195, 29–45. doi: 10.3354/meps195029
- Zimmerman, R. C., Kohrs, D. G., Steller, D. L., and Alberte, R. S. (1997). Impacts of CO₂ enrichment on productivity and light requirements of eelgrass. *Plant Physiol.* 115, 599–607.
- Zondervan, I., Zeebe, R. E., Rost, B., and Riebesell, U. (2001). Decreasing marine biogenic calcification: a negative feedback on rising atmospheric pCO₂. Glob. Biogeochem. Cycles 15, 507–516. doi: 10.1029/2000GB 001321

Conflict of Interest Statement: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

Copyright © 2016 Sultana, Casareto, Sohrin, Suzuki, Alam, Fujimura and Suzuki. This is an open-access article distributed under the terms of the Creative Commons Attribution License (CC BY). The use, distribution or reproduction in other forums is permitted, provided the original author(s) or licensor are credited and that the original publication in this journal is cited, in accordance with accepted academic practice. No use, distribution or reproduction is permitted which does not comply with these terms.