



Multiple Salinity Equilibria and Resilience of Persian/Arabian Gulf Basin Salinity to Brine Discharge

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Specialty section:

This article was submitted to
Coastal Ocean Processes,
a section of the journal
Frontiers in Marine Science

Received: 08 April 2020

Accepted: 22 June 2020

Published: 10 July 2020

Citation:

Ibrahim HD, Xue P and
Eltahir EAB (2020) Multiple Salinity
Equilibria and Resilience
of Persian/Arabian Gulf Basin Salinity
to Brine Discharge.
Front. Mar. Sci. 7:573.
doi: 10.3389/fmars.2020.00573

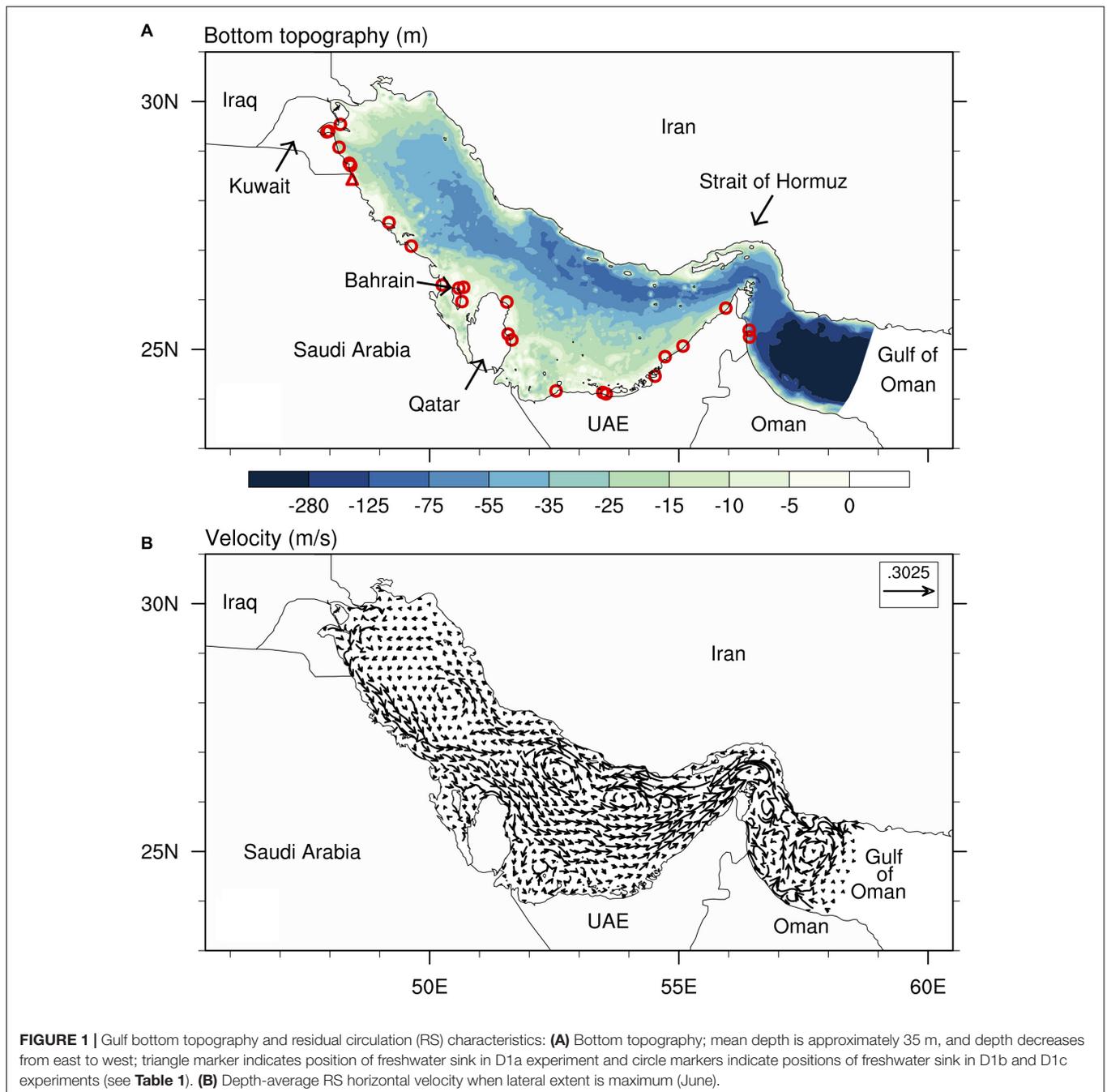
The Persian/Arabian Gulf is the most important region for seawater desalination. Surrounding countries produce about 50% of global desalinated seawater. If Gulf salinity significantly rises because of desalination effluent (brine), marine ecosystems and the water supply for many population centers will be threatened. In order to quantify current and future impacts of seawater desalination on Gulf salinity and avoid costly environmental problems, it is vital to first examine the present Gulf salinity state and its response to salinity perturbation (i.e., determine its stability). Here, using a coupled Gulf-Atmosphere numerical climate model, we test the hypothesis that the Gulf has a single stable equilibrium state under the current climate. Simulations with different initializations under identical external forcing show that the natural coupled Gulf-Atmosphere system may exhibit a mixture of unstable and stable equilibrium salinity states. When continuous salinity perturbation is added to the simulations, results show that the present Gulf equilibrium state, characterized by annual mean basin-average salinity of about 40.5 g/kg, is stable. We conclude that Gulf basin salinity is resilient to present brine discharge activities under the current climate.

Keywords: Persian/Arabian Gulf Salinity, overturning circulation, salinity perturbation, seawater desalination, equilibrium state, stability

INTRODUCTION

The Persian/Arabian Gulf, hereafter Gulf, is a semi-enclosed shallow sea with mean depth of about 35 m and approximate surface area of 2.35×10^{11} m². It is connected to the Gulf of Oman through the Strait of Hormuz, hereafter Strait (**Figure 1A**). Because of the arid climate there, Gulf evaporation (about 1.84 m/year) greatly exceeds freshwater input (about 0.28 m/year) from precipitation and river runoff (Privett, 1959; Johns et al., 2003; Xue and Eltahir, 2015), causing Gulf waters to be hypersaline. Gulf freshwater loss by evaporation is balanced by a net inflow of saline water from the Strait, mainly through upper layers, and the hypersaline return flow at the Strait, which contains the salt left behind after evaporation, is through deep layers. Within the Gulf the main circulation pattern is a basin-wide cyclonic gyre with several cyclonic eddies in the central and northern Gulf, and strong coastal currents extend along the coast of Saudi Arabia and the United Arab Emirates (**Figure 1B**).

The countries surrounding the Gulf are in a region of strategic and ecological significance and have some of the highest rates of economic and population growth; but the Gulf is increasingly



under environmental stress (Roberts et al., 2010; Sheppard et al., 2010; Sale et al., 2011). Because of scarce renewable freshwater resources, many surrounding countries rely on seawater desalination, about 50% of global capacity (Lattemann and Hopner, 2008; Jones et al., 2019). This anthropogenic activity removes about 5 km³/year of freshwater from the Gulf (Ibrahim and Eltahir, 2019), and the associated brine (about 12 km³/year) is disposed into the Gulf (Jones et al., 2019). Brine threatens Gulf marine ecosystems (Dupavillon and Gillanders, 2009; Roberts et al., 2010; Sale et al., 2011). Meanwhile, basin buildup of the salt in brine threatens potable

water supply there because energy required for desalination increases as salinity increases (Elimelech and Phillip, 2011; Voutchkov, 2018). While the impact of seawater desalination on the regional water quality has been studied (Ibrahim and Eltahir, 2019), there are rising concerns about the sustainability of seawater desalination and its impact on Gulf basin-scale salinity and marine ecosystem into the future (Sheppard et al., 2010; Elimelech and Phillip, 2011; Sale et al., 2011). In order to quantify current and future impacts of seawater desalination on Gulf salinity and avoid costly environmental problems, it is therefore vital to understand the stability of Gulf salinity

state and its resilience to salinity perturbations resulting from seawater desalination.

The Gulf is the hot spot for seawater desalination brine discharge (mainly into near-surface waters), accounting for about 50% of the 51.7 billion m³/year global brine production; but ocean and sea brine discharge are also rising in other arid-coastal regions: North Africa (20.3%), East Asia and Pacific (10.5%), North America (3.9%), Western Europe (5.9%), and Latin America and Caribbean (3.9%; Jones et al., 2019). This study therefore develops and tests a methodology that can be applied in these other regions to quantify the salinity response to salinity perturbations and hence to facilitate improved brine management strategies in other seas and oceanic regions.

Most climate systems on earth are characterized by a single stable equilibrium state. But climate systems with multiple equilibrium states (i.e., sensitive to initial conditions) are of practical and theoretical importance (Stommel, 1961; Lorenz, 1968, 1970; Stocker, 1999; Ferreira et al., 2011). The equilibrium state of climatic elements (e.g., temperature, salinity) characterize the past and provide a fixed baseline from which to measure natural or anthropogenic climatic change. Beyond thresholds that separate equilibrium states, natural feedbacks enhance perturbations; whereas within such thresholds, these feedbacks attenuate perturbations. Hence, multiple equilibrium states affect resilience of climate systems to natural or anthropogenic perturbations. Analysis of observations and simulation experiments in numerical models have shown the existence of multiple equilibrium states at the global scale (Budyko, 1969; Sellers, 1969; Broecker et al., 1985; Langen and Alexeev, 2004; Ferreira et al., 2011) and at regional scales (Claussen and Esch, 1994; Wang and Eltahir, 2000; Zeng and Neelin, 2000; Dekker, 2010; Bathiany et al., 2012).

In coastal regions and marginal seas [e.g., the Mediterranean Sea where multiple equilibrium states have been found (Ashkenazy et al., 2012)], a switch from low to high basin-average salinity can have a considerable economic impact. For example, high salinity in coastal waters can adversely affect marine productivity and thus fish catches (Iselin, 1940). This paper aims to test the hypothesis that the Gulf annual mean basin-average salinity (hereafter, Gulf equilibrium salinity) is unique and stable (i.e., unlikely to change) under current climate ocean and atmosphere boundary conditions and current seawater desalination activities. Using a coupled Gulf-Atmosphere regional model (GARM), we performed and analyzed two sets of simulation experiments to examine the uniqueness and stability of Gulf equilibrium salinity. The methodology, results and discussion, and conclusion are given, respectively, in the following sections.

METHODOLOGY

Observational Datasets

Observations of Gulf water properties, especially salinity, are limited in space and time coverage. There are few published basin-wide survey results including a 1948 summer cruise (Emery, 1956), the 1976 wintertime expedition of the Atlantis

from Woods Hole Oceanographic Institution (Brewer and Dryssen, 1985), and the surveys from February to June 1992 on the Mt. Mitchell expedition (Reynolds, 1993). Historical hydrographic datasets, compiled mostly from the Master Oceanographic Observations Data Set (MOODS) maintained by the U.S. Naval Oceanographic Office, was used to describe the seasonal variability of Gulf seawater (Brewer and Dryssen, 1985). However, this dataset has large temporal gaps and a bias toward winter and spring months. Because coverage of *in situ* salinity is limited, we used the Simple Ocean Data Assimilation (SODA v2.2.4) reanalysis data for salinity analysis. SODA is a data assimilation effort to reconstruct the global ocean climate, which combines a first guess of the evolving ocean state from a numerical model, and a correction based on direct observations, including the abovementioned observation efforts. The resulting ocean reanalysis provides an improvement overestimates based solely on either the sparse set of observations or numerical simulation (Carton et al., 2000; Carton and Giese, 2008).

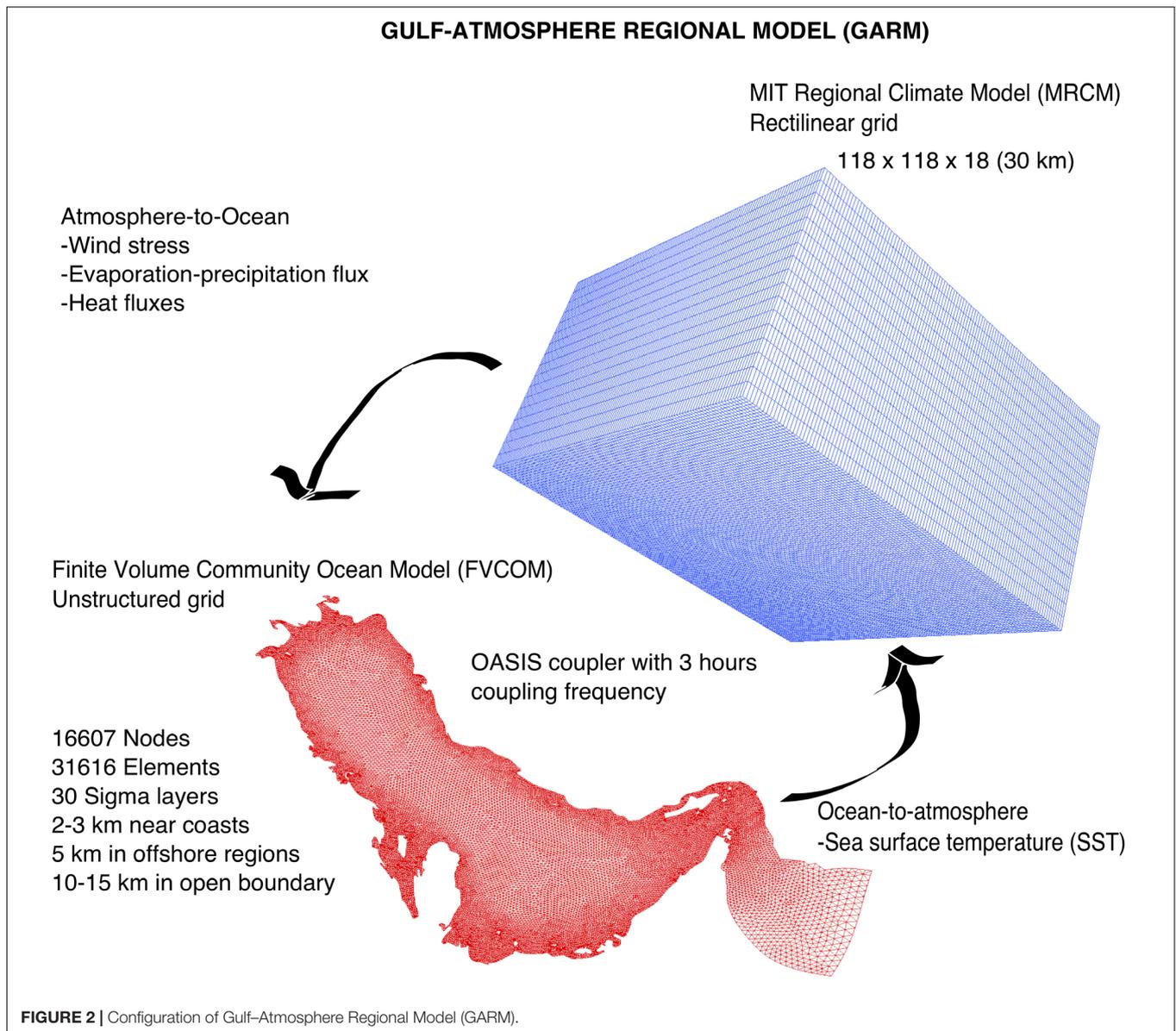
We used two analyses datasets to validate simulated Gulf sea-surface temperature (SST). First, for monthly means, we used NOAA 1/4 degree daily Optimum Interpolation Sea Surface Temperature (Reynolds et al., 2007), OISST v2, because it is available for our simulation experiments period. Second, for climatological monthly means, we used the Jet Propulsion 1 km by 1 km MUR-JPL-L4-GLOB-SST v4.1, which is available only from mid 2002 to present (Chin et al., 2017).

To validate simulated Gulf water surface elevation, we used daily 1/4 degree, merged, absolute dynamic topography (SSALTO/DUACS merged-allsat-phy-14) from Copernicus Marine Environment Monitoring Service (Copernicus, 2019). This dataset combines measurements by 11 satellites from 1993 to 2018. The Gulf is shallow and there are a few other coastal altimetry datasets (e.g., ALES and PISTACH) that apply shallow-region corrections to global altimetry measurements. However, these datasets have only been validated for individual satellites with short periods, which is unsuitable for the climatological analysis here.

Numerical Model

We used a high-resolution, coupled GARM with current climate ocean and atmosphere boundary conditions. GARM includes the two-way interaction between its atmosphere and ocean components with a coupling frequency of 3 h (Figure 2). Synchronous coupling is achieved with OASIS3 coupler (Valcke, 2013), and simulations in the two components are integrated forward simultaneously in a parallelized computation environment for low intrusiveness and high portability. OASIS coupler controls data transfer between GARM components, interpolates the coupling data between the different grids, coordinates the synchronization of the constituent models' simulation, and allows the free air-sea interaction between GARM components.

Gulf-Atmosphere regional model atmospheric component is the MIT Regional Climate Model (MRCM), an advanced version of Regional Climate Model version 3 (RegCM3; Giorgi and



Mearns, 1999; Pal et al., 2007). MRCM has new or improved physical schemes and high skill to simulate different regional climates (Winter and Elfatih, 2009; Marcella and Eltahir, 2012, 2014; Gianotti and Eltahir, 2014a,b; Im and Eltahir, 2017). GARM ocean component that simulates Gulf hydrodynamics is the unstructured-grid Finite Volume Community Ocean Model (FVCOM; Chen et al., 2003, 2006). Because of the different spatial scales of atmospheric and Gulf residual circulation, GARM-MRCM horizontal domain is larger (29–61°E and 12–40.5°N), with 30 km resolution in 120 by 120 grids. GARM-FVCOM has horizontal resolution of 2–3 km near coast, 5 km in offshore regions and 10–15 km at the open ocean boundary. GARM-FVCOM is configured with 30 generalized sigma layers where water depth is <60 m; while in regions with water depth >60 m, the thickness of the top and bottom 5 layers are set to 2 m, and the other 20 intermediate layers are uniformly

divided for the remaining water column. Thus, vertical resolution is <1 m for nearshore Gulf regions and 1–2 m in most offshore Gulf regions.

Gulf-Atmosphere regional model-MRCM boundary condition is from current climate simulations of ECHAM5/Max Planck Institute Ocean Model (MPIOM), now Max Planck Institute Earth System Model (MPI-ESM) in phase 5 of the Coupled Model Intercomparison Project (CMIP5). GARM-FVCOM boundary condition is from climatological monthly mean fields of temperature (Locarnini, 2010) and salinity (Antonov, 2010) from the World Ocean Atlas 2009; mean flow velocity boundary condition is dynamically calculated with wind, temperature and salinity. River runoff into the Gulf from the Shatt Al Arab/Arvand Rud, the main source of freshwater into the Gulf, was specified based on streamflow statistics for the Tigris and Euphrates River Basins (Saleh, 2010).

Gulf-Atmosphere regional model has been used successfully to estimate Gulf heat and water budgets (Xue and Eltahir, 2015) and to characterize the dynamic interplay between brine discharge and Gulf residual circulation (Ibrahim and Eltahir, 2019). MRCM and FVCOM have individually been used in many regional climate numerical studies. Details of GARM configuration and sensitivity to forcing fields are described at length in Xue and Eltahir (2015). For the sake of brevity, additional GARM details and skill assessments are given in the **Supplementary Material**. The methods for computing lateral volume exchange fluxes at the strait, and for converting the volume fluxes into equivalent over-Gulf freshwater fluxes (i.e., m/year) to facilitate comparison, were given in Xue and Eltahir (2015); equation 10) and thus are not repeated here. This is convenient for comparing the magnitudes of the water budget and fluxes, especially since evaporation (~1.8 m/year) plays a vital role in Gulf-Atmosphere dynamics. The average annual lateral fluxes are referenced to Sv. unit (see **Table 2**) to facilitate comparison with previous studies.

Design of Simulation Experiments

We performed two sets of experiments: (1) to examine the influence of initial salinity condition on the establishment of possible equilibrium salinity states, six 10-year-long experiments (Ic1–Ic6) were initialized with different basin-average salinities, and seawater desalination perturbation was excluded. (2) to establish the response of Gulf equilibrium salinity states to current seawater desalination activities (i.e., determine the stability equilibrium salinity states), three 19-year-long experiments were initialized with a low (D1a and D1c), and high (D1b) initial basin-average salinities, and continuous seawater desalination perturbation was added.

Ic1–Ic6 Scenarios: Variation of Initial Basin-Average Salinity

Ic1 started with spatially variable initial salinity and temperature (basin-averages are shown in **Table 1**), representing current

Gulf conditions based on climatological monthly mean fields of temperature (Locarnini, 2010), and salinity (Antonov, 2010) from the World Ocean Atlas 2009. To explore the possible lower limit of Gulf equilibrium salinity, a low initial basin-average salinity was assumed for Ic2; and to explore the possible upper limit, a high initial basin-average salinity was assumed for Ic3 (**Table 1**). Ic4–Ic6 initial basin-average salinities were chosen between those of Ic1 and Ic2 experiments (**Table 1**).

D1a–D1c Scenarios: Addition of Continuous Seawater Desalination Perturbation

The lowest simulated Gulf equilibrium salinity (in Ic2 experiment) was chosen as the initial basin-average salinity of D1a and D1c experiments (**Table 1**). Because the only difference between D1a and D1c experiment is the freshwater sink magnitudes (**Table 1**), comparing the results of these two experiments elucidates the influence of salinity perturbation flux magnitude on the stability of Gulf equilibrium salinity. In D1b experiment, GARM was initialized with a higher-than-present basin-average salinity of 44 PSU (**Table 1**) in order to demonstrate the stability of the present Gulf equilibrium salinity as it restores back from higher-than-present salinity initialization even with desalination activities. Note that the desalination perturbation in D1b and D1c simulation (4.35 km³/year) is the same order of magnitude as current Gulf seawater desalination capacity, about 5 km³/year. There are more than 850 seawater desalination plants in the Gulf, and most of these plants, including the largest 24 (based on production capacity ≥100,000 m³/day) are located near the Arabian coast of the Gulf (Ibrahim and Eltahir, 2019). Hence, for practical reasons, and since small plants are assumed to have negligible impact, we use here only the production capacities of the 24 largest plants as the freshwater sink in D1b and D1c experiments.

RESULTS AND DISCUSSION

Gulf Response to Variation of Initial Basin-Average Salinity

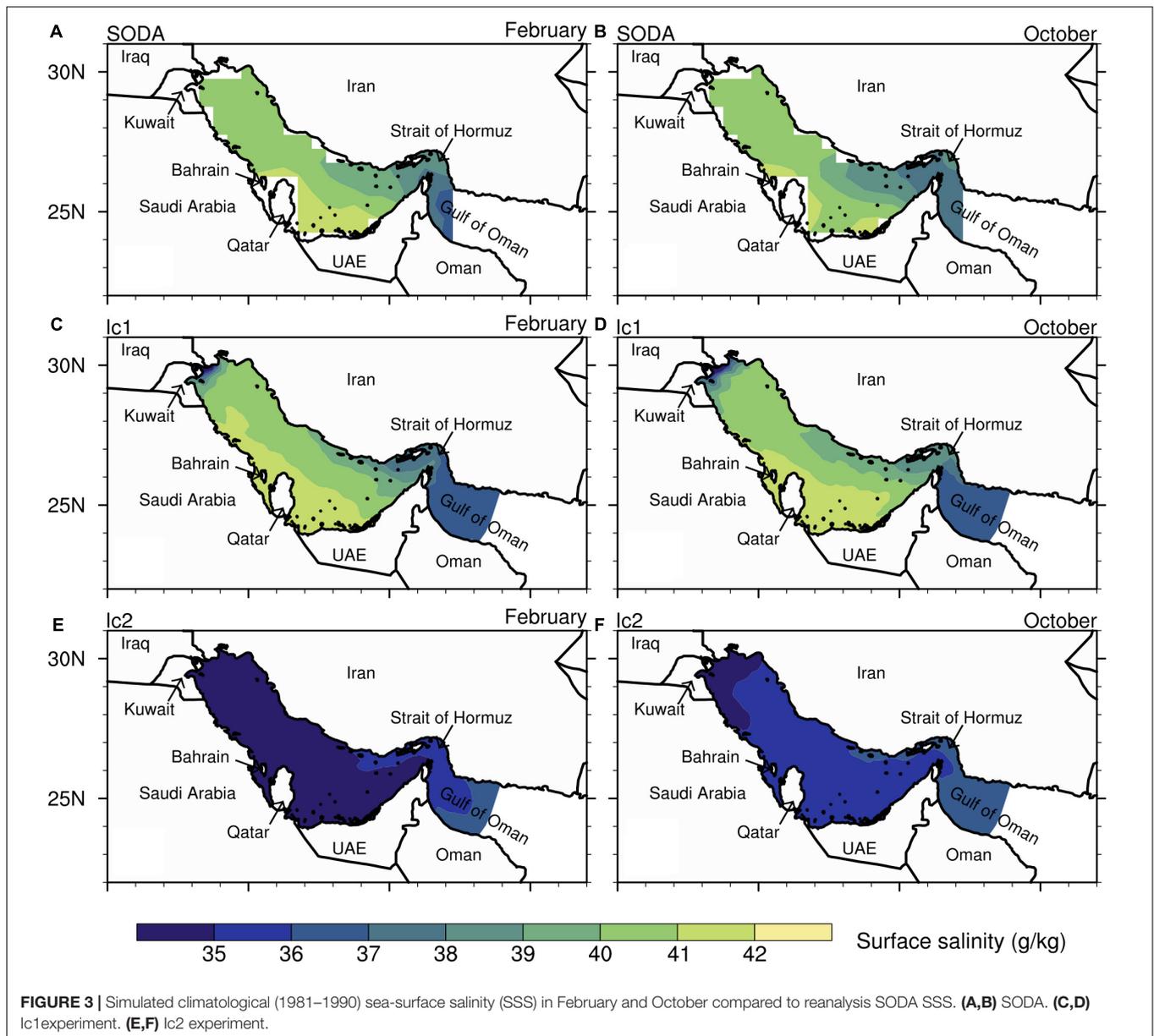
Corresponding to Ic1 and Ic2 experiments, GARM simulates two different equilibrium salinities.

(1) Ic1 simulation Gulf sea-surface salinity (SSS) shows a close agreement, in magnitude and spatial pattern, with the SSS from the SODA dataset (**Figures 3A–D**). Ic1 equilibrium salinity is about 40.5 g/kg (**Figure 5A**). Water with salinity between 40–41 g/kg occupies most of the Gulf region while the saltiest water (salinity > 41 g/kg) is located in the southern Gulf. Salinity gradients around the Strait show that relatively fresher water enters the Gulf and becomes saltier because of excessive evaporation. Although river discharge is not reflected in the SODA dataset because of its coarse resolution, the river plume from the Shatt Al Arab/Arvand Rud traveling southeastward along the coast is well reflected in the GARM simulation (**Figures 3C,D**). (2) Ic2 simulation (initialized with a salinity of 24 g/kg) shows SSS of about 35–36 g/kg with little spatial variation (**Figures 3E,F**). Ic2 equilibrium salinity is about 36.8 g/kg (**Figure 5A**).

TABLE 1 | Overview of simulation experiments: in the first set of 10-year-long experiments GARM is initialized with different Gulf basin-average salinities (Ic1–Ic6); and in the second set of 19-year-long experiments GARM is initialized with a low (D1a and D1c); and a high (D1b) initial Gulf basin-average salinity and continuous seawater desalination perturbation is included for the duration of the experiments.

Experiment ID	Initial basin-average salinity (g/kg)	Initial basin-average temperature (°C)	Freshwater sink/desalination (m/year)
Ic1 ^a	40.10	21.40	No
Ic2	24.80	24.80	No
Ic3	42.60	24.80	No
Ic4	38.00	24.80	No
Ic5	39.00	24.80	No
Ic6	40.00	24.80	No
D1a	36.80	24.80	9.4e-05
D1b	44.00	24.80	0.0185
D1c	36.80	24.80	0.0185

^aSensitivity analysis showed that initial basin temperature of 24.8°C has no impact on Ic1 Gulf equilibrium salinity.

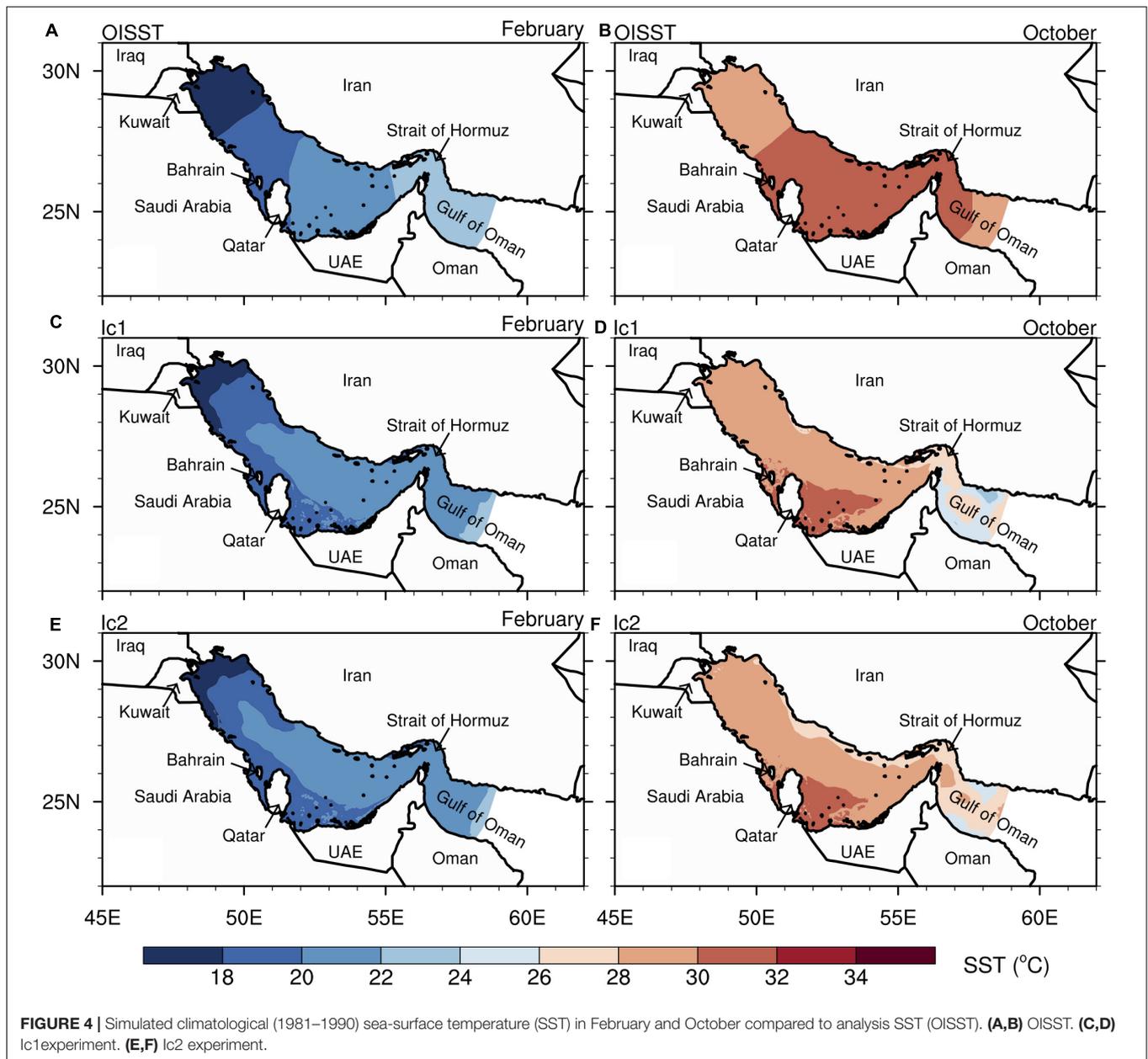


In contrast to the striking difference in SSS between Ic1 and Ic2, SST horizontal structures are similar and basin-average SSTs are identical in the two cases (**Figures 4C–F, 5B**), suggesting that the salinity difference between these two cases has little impact on water temperature. It also suggests that the heat fluxes, which control the water temperature, is mostly unaffected by the difference in equilibrium salinity between Ic1 and Ic2 simulations.

Ic1 and Ic2 water surface elevations (WSE) are different (**Figure 5C**). WSE in Ic1 simulation is similar to the satellite-derived value, which implies that the well-simulated temperature and salinity reproduces the correct water density and therefore WSE. On the other hand, Ic2 WSE is much higher than the satellite-derived value (**Figure 5C**) because of the much lower Ic2 salinity (**Figures 3E,F, 5A**).

Consequently, the overturning circulation between the Gulf and the Indian Ocean that is associated with Ic1 and Ic2 equilibrium states have different temporal and vertical structures at the Strait. In Ic1 equilibrium, lateral inflow to the Gulf and outflow from the Gulf have their respective maximum and minimum in February and October (**Figure 5D**): lateral inflow is in top layers while the outflow is in bottom layers (**Figures 6A,B**), as evidenced *in situ* observations (Johns et al., 2003). The annual structure and direction of Ic1 overturning circulation, year-round counterclockwise inflow in top layers and outflow in bottom layers, thus represents an anti-estuarine circulation.

Unlike Ic1 simulation, lateral fluxes in Ic2 simulation have a much weak seasonality (**Figure 5D**). More importantly, in February, flow into the Gulf is in the northern shallow Strait and flow out of the Gulf is in the deep southern channel

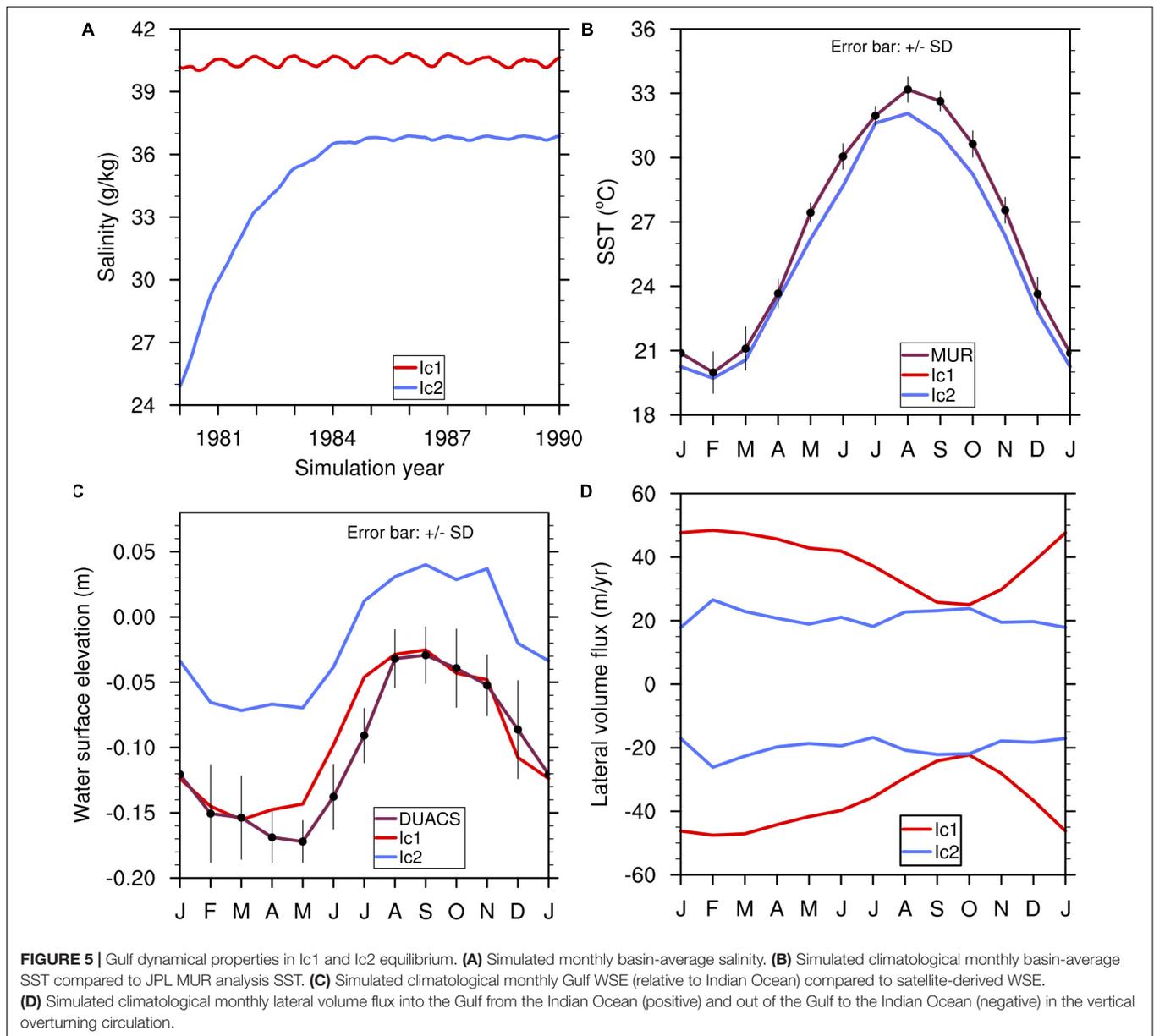


(Figures 6C,D); whereas in October the opposite vertical flow structure occurs. Therefore, the annual structure and direction of Ic2 overturning circulation at the Strait, which switches between counterclockwise and clockwise patterns, represents a mixture of estuarine and anti-estuarine circulations.

In Ic4–Ic6 experiments, GARM was initialized with salinity 38, 39, and 40 g/kg, and simulation results show that monthly mean basin-average salinities and WSEs of Ic4–Ic6 remained between those of Ic1 and Ic2 equilibria (Figures 7A–C); and likewise with Ic4–Ic6 lateral inflow and outflow (Figure 7D). However, in Ic3 experiment where GARM was initialized with a salinity of 42.6 g/kg (which is higher than the present basin-average salinity of around 40.5 g/kg simulated in Ic1 experiment), the Gulf system is restored back to Ic1 equilibrium salinity. Furthermore, whereas

the overturning circulation associated with Ic3 equilibrium is very similar to that of Ic1 equilibrium (anti-estuarine circulation), the overturning circulations associated with Ic4–Ic6 equilibria have vertical structures that are similar to the Ic2 vertical structure: mixture of anti-estuarine and estuarine circulation.

An important follow-up question is the stability of the different dynamical structures associated with Ic1–Ic6 simulations. A system in an equilibrium state is said to be stable if the system returns to its equilibrium state after a small, externally induced disturbance; whereas in an unstable equilibrium, natural feedbacks enhance small perturbations, and move the system away from the original equilibrium state. In the following section, we test the stability of Ic1–Ic6 equilibria to salinity perturbation in order to understand the resilience



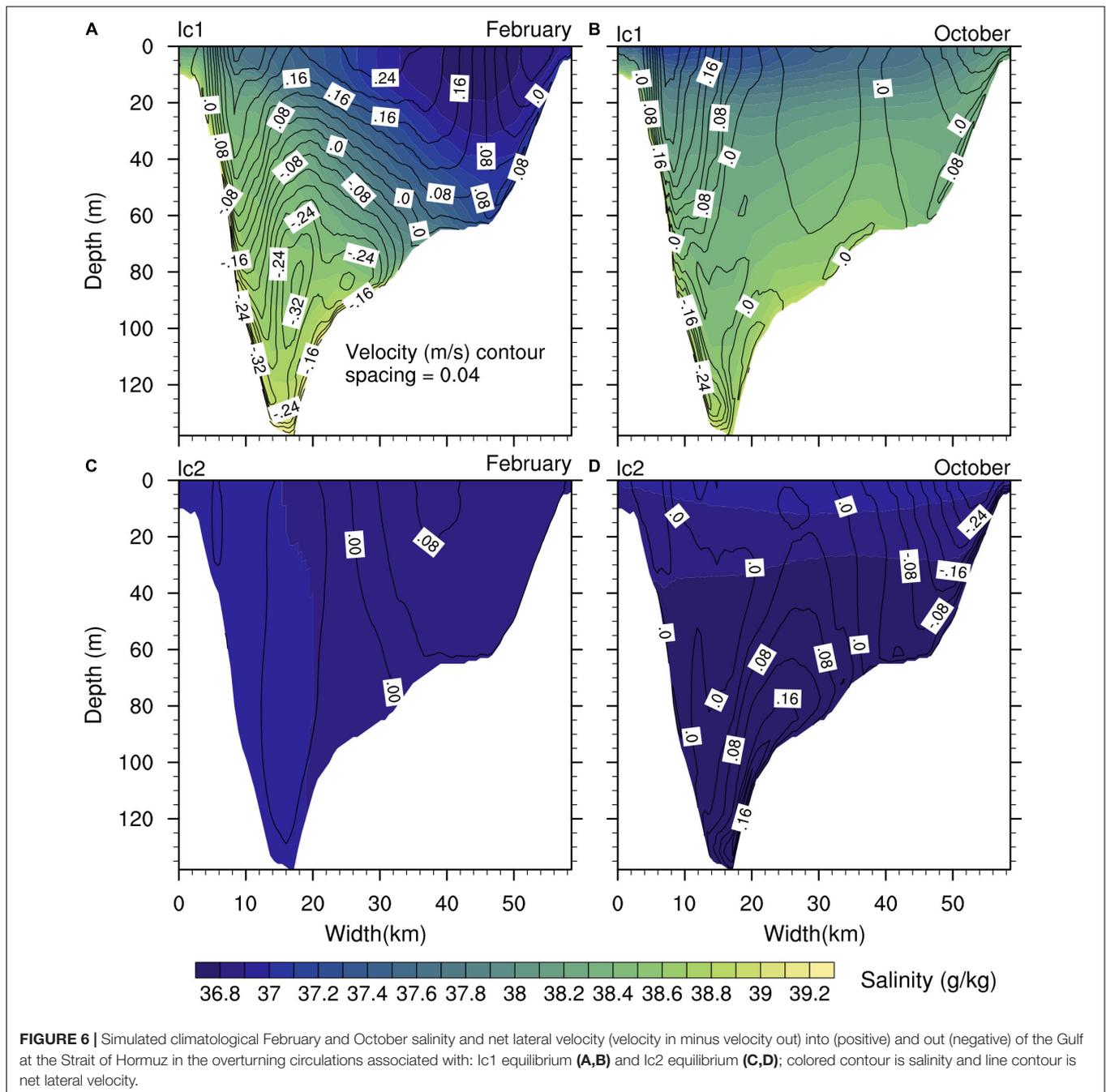
of the Gulf basin salinity to brine discharge from seawater desalination plants.

Stability of Equilibrium States to Seawater Desalination Perturbation

Model results show that the simulated Gulf system, when initialized with different salinities, exhibits both unstable and stable equilibrium salinity states in response to continuous seawater desalination perturbation. Ic1 equilibrium salinity state is stable because the Gulf system returns to this state after a seawater desalination perturbation was introduced starting from either a lower—36.8 g/kg—or a higher—44 g/kg—initial salinity (Figure 8A). Although multiple possible equilibrium salinity states were simulated when the GARM was initialized

with different salinities in Ic1–Ic6 experiments, these states are not all stable. For example, Ic2 simulation was initialized with salinity of 24.8 g/kg and Gulf basin salinity increased to reach an equilibrium state of about 36.8 g/kg (Figures 5A, 7A,B), which turned out to be an unstable equilibrium. This is because when seawater desalination perturbation was introduced (D1a experiment), Gulf basin salinity continued to increase, crossing over several equilibria (Ic4–Ic6), until it reached Ic1 equilibrium salinity state—the current Gulf equilibrium salinity (Figure 8A). Whereas in D1a experiment the lateral fluxes became stronger during the adjustment process to Ic1 equilibrium salinity state, in D1b experiment the lateral fluxes became weaker during the adjustment process (Figure 8B).

At the annual time scale Ic1 equilibrium salinity and net lateral flow are also stable (Table 2). However, although

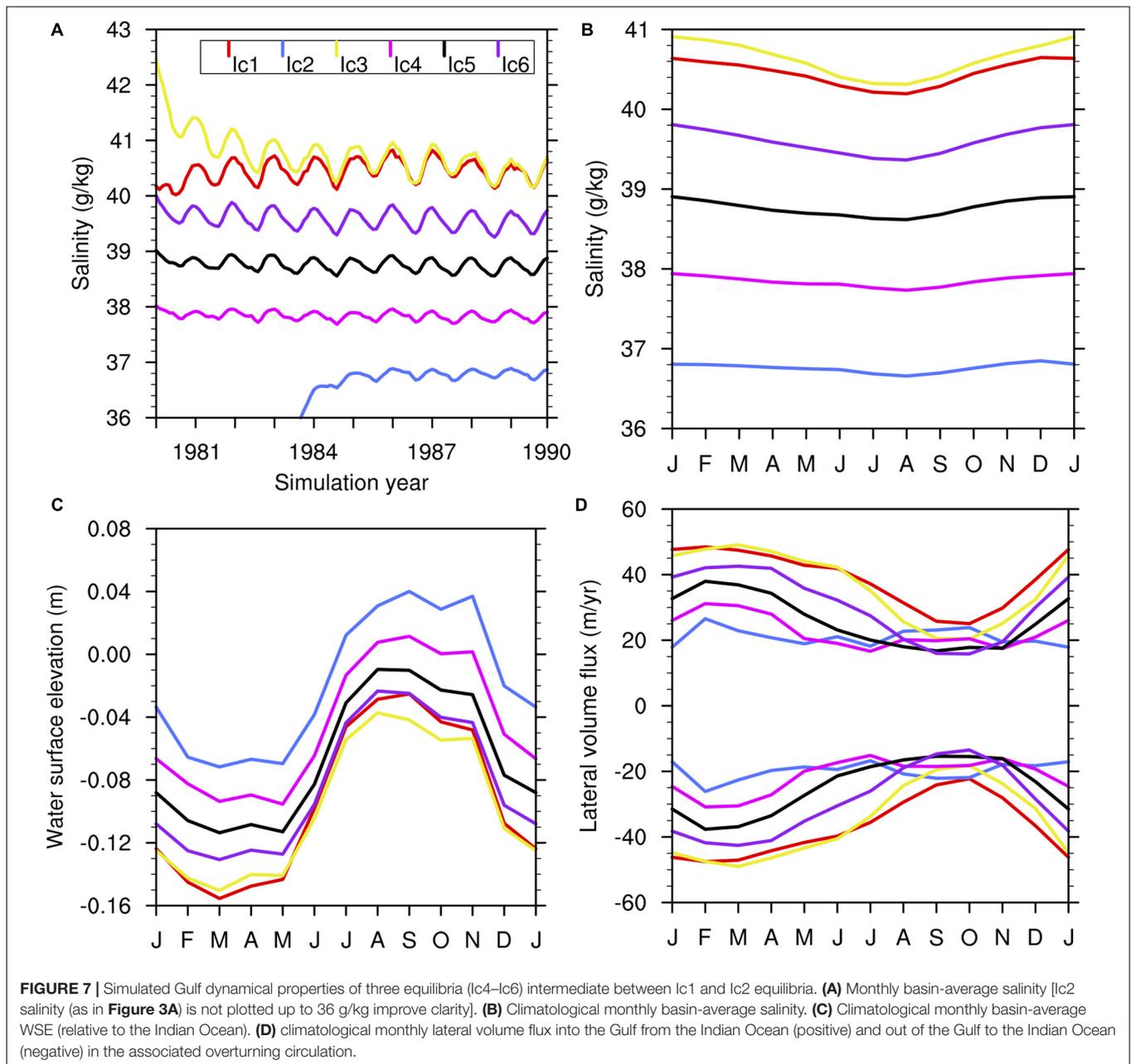


Gulf equilibrium salinity in D1b experiment converges to Ic1 equilibrium salinity, the standard deviation of depth-average annual mean salinity in these two experiments are different (**Figure 9**), indicating that brine discharge causes substantial changes in regional salinity maxima and minima. The largest changes in regional salinities are in Gulf southwest region, especially in nearshore regions of Saudi Arabia, Bahrain, Qatar, and the United Arab Emirates (**Figures 9C,D**). These regional salinity changes, which has been quantified and discussed in Ibrahim and Eltahir (2019), are very likely to have significant impact on local marine ecosystems. Accordingly, D1b experiment

identifies Gulf subdomains that should be monitored for ecosystem stress due to brine discharge.

Statistical Analysis of the Stability of Current Gulf Equilibrium Salinity

We performed non-linear regression analysis, using the exponential model fitting, and linear regression analysis to further examine whether the Gulf has reached (or approaches) a steady equilibrium salinity state in D1a and D1b simulations. As the first step in both regression analyses, the climatological



seasonal cycle is removed from the monthly mean basin-average salinity. For D1b experiment, the robustness of the non-linear regression model is visually evident (**Figure 10B**). Using data from the first 13-years (1981–1993) to build the model and data from the last 6-years (1994–1999) to test the model prediction, the non-linear regression model shows not only good fitting, but also reliable prediction: the 95% confidence bound is narrow (about 0.45), and residual errors of the fitting model are small (± 0.3). The well-fitted regression model ($Y = a + b \cdot \exp(-c \cdot x)$, $a = 40.56$; $b = 2.4318$; $c = 0.074$) result suggests that Gulf annual mean basin-average salinity (i.e., equilibrium salinity) in D1b simulation has reached an equilibrium of ~ 40.56 g/kg (± 0.3), where Y is the monthly mean

basin average salinity (climatological seasonal cycle removed), and x is the time (in month).

Furthermore, we performed a linear trend analysis for D1b simulation using the last 8-years simulation results: that is, we test a linear regression model, $Y = \beta_0 + \beta_1 x + \varepsilon$, where Y is the monthly mean basin-average salinity (climatological seasonal cycle removed), x is time (in month), and ε is the random error. The coefficients of the best-fit linear regression model are $\beta_0 = 40.465$, $\beta_1 = 0.00068943$. This implies that it will take 120 years for Gulf annual mean basin-average salinity to increase by 1 g/kg, and this is so only if the trend is statistically significant. We then conducted a statistical hypothesis test on the linear regression coefficients, and F -test shows that we

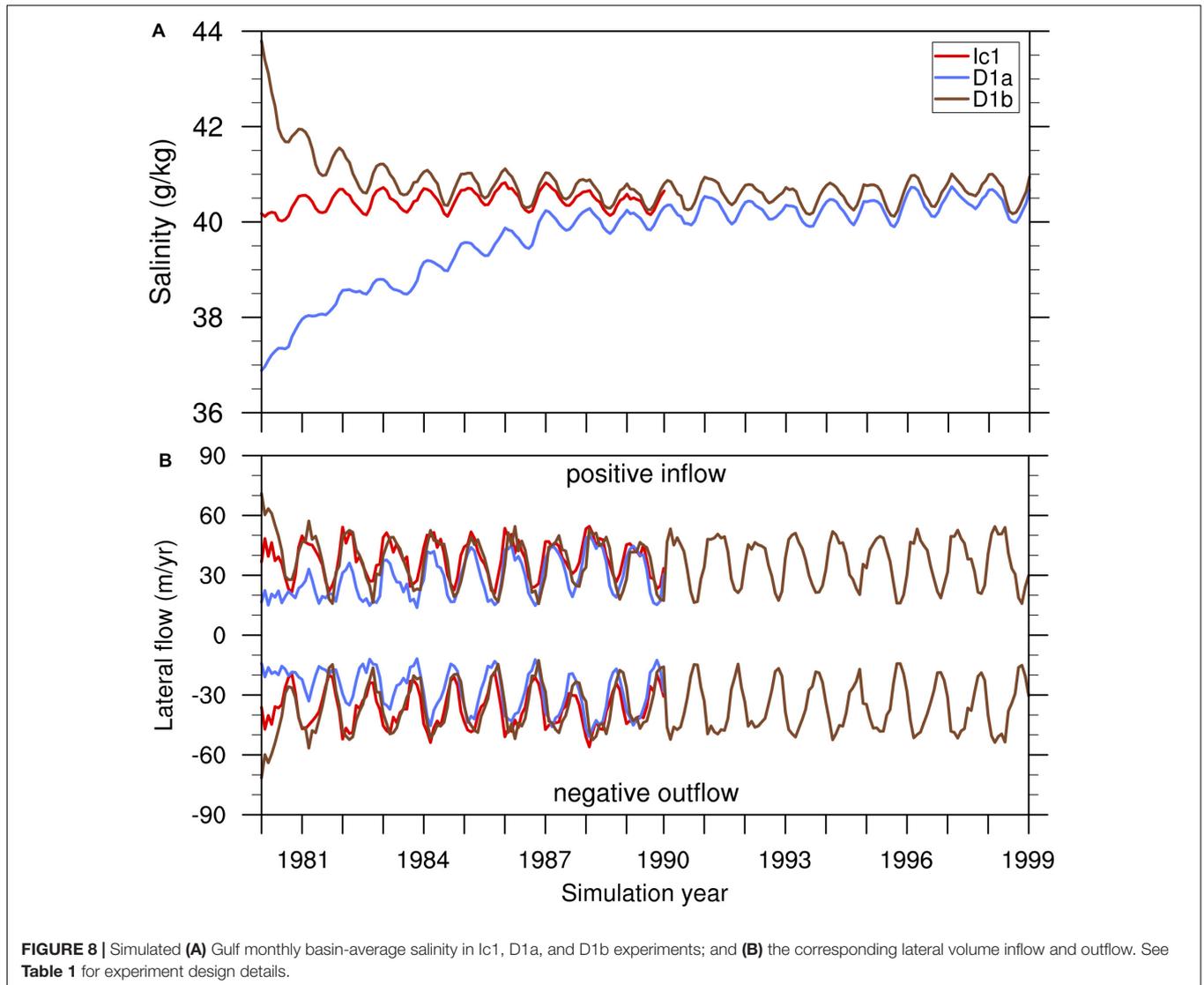
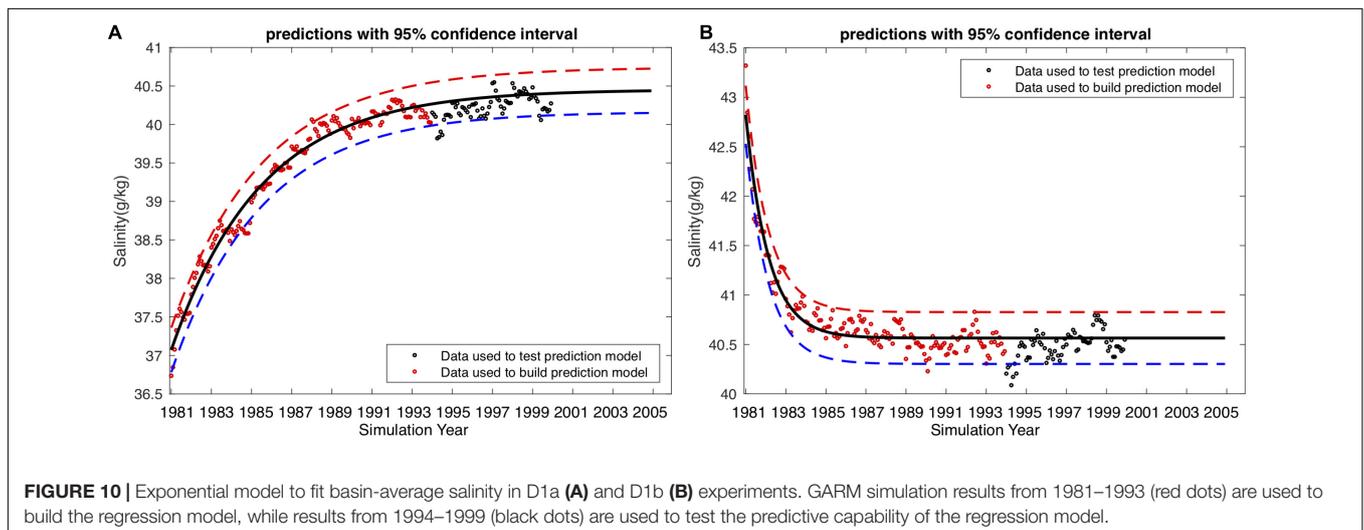
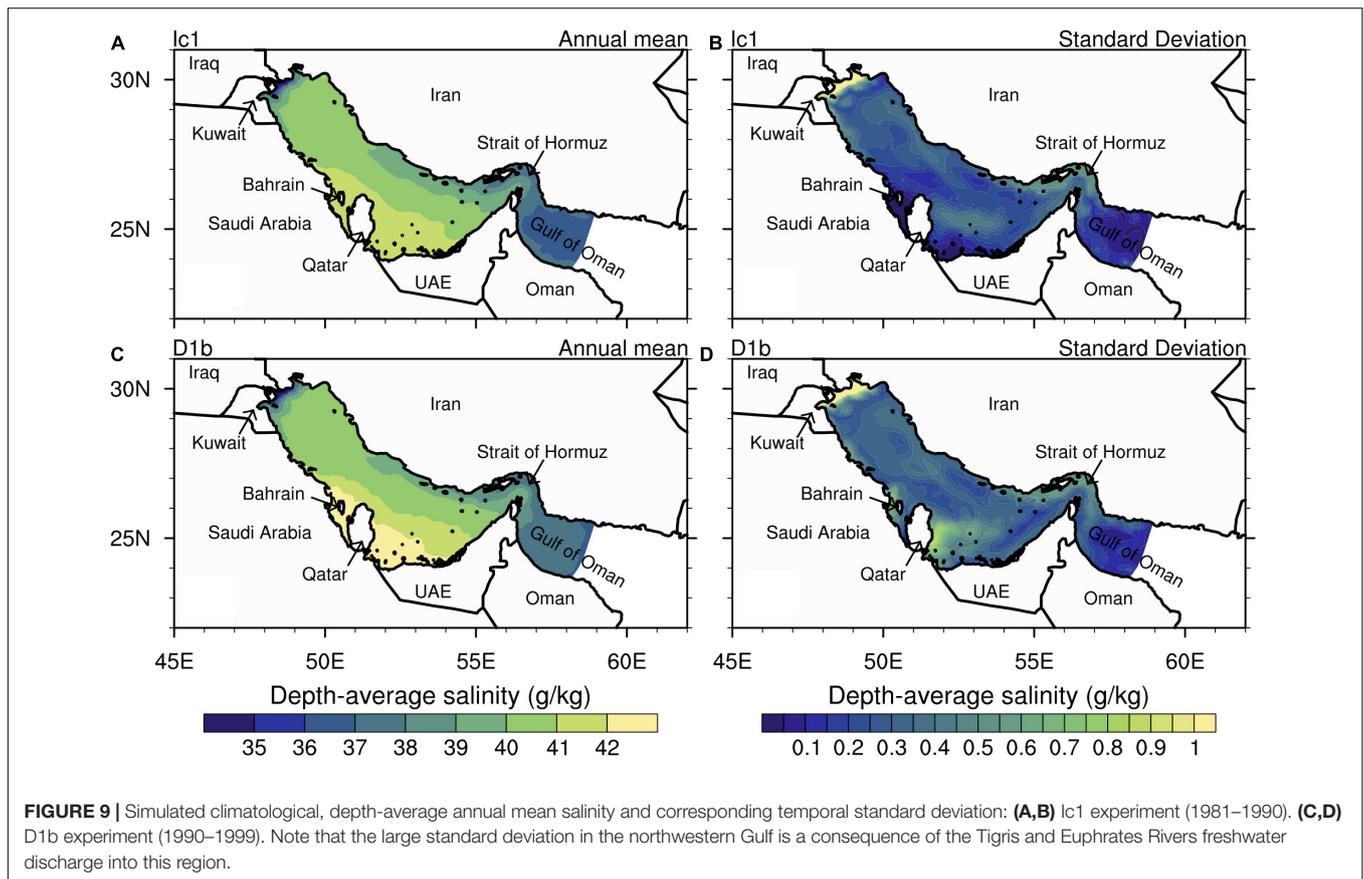


TABLE 2 | Gulf annual mean basin-average salinity and lateral inflow (in), outflow (out), and net exchange fluxes^a in Ic1 experiment.

Simulation year	Salinity (g/kg)	Lateral in (m/year)	Lateral out (m/year)	Lateral net (m/year)
1981	40.20	35.91	-34.37	1.54
1982	40.43	37.32	-35.72	1.60
1983	40.45	38.97	-37.51	1.46
1984	40.44	38.83	-36.82	2.01
1985	40.46	39.65	-38.11	1.54
1986	40.58	37.32	-35.91	1.41
1987	40.51	38.42	-36.54	1.88
1988	40.56	40.17	-38.77	1.40
1989	40.40	41.67	-40.10	1.57
1990	40.42	36.50	-34.90	1.60
Average	40.45 (±0.10)	38.48 (±1.76)	-36.88 (±1.77)	1.60 (±0.2)

For reference with previous studies, 30 m/year over-Gulf flux corresponds to ~ 0.22 Sv exchange flow. ^aNote: unlike in Xue and Eltahir (2015) where these lateral inflow and outflow fluxes were computed with velocity at monthly resolution, here they are computed with velocity at 3-hourly resolution (GARMA coupling frequency); hence the discrepancy there and here in the annual mean inflow and outflow; but the annual net lateral flux there and here are similar.



can't reject the null hypothesis [β_1 (slope) = 0] since the P -value = 0.29 > 0.05. Consequently, no statistically significant trend can be confirmed, reinforcing the conclusion that the Gulf has reached an equilibrium salinity state.

For D1a simulation, non-linear regression model result, using an exponential model ($Y = a + b \cdot \exp(-c \cdot x)$, $a = 40.45$; $b = -3.4472$; $c = 0.0186$), also fit robustly (**Figure 10A**) and suggests that Gulf basin-average salinity will eventually

reach a steady state of ~ 40.45 g/kg (± 0.3). However, linear regression model with the last 8-year of simulation results shows a statistically significant trend ($\beta_0 = 40.11$, $\beta_1 = 0.002312$ with P -value = $2.4e-6 < 0.05$), suggesting that the model has not reached steady state (**Figure 10A**). The reason why Gulf equilibrium salinity in D1a simulation has not reached the steady state is because of the small magnitude of seawater desalination perturbation (**Table 1**). Note that D1a and D1b simulations are

designed to examine the stability of Gulf equilibrium salinity state: D1a simulation shows the transition from an unstable (Ic2) to the current stable (Ic1) equilibrium salinity state with small desalination, and D1b simulation shows that even when the Gulf system is initialized with a higher-than-present salinity and with a large desalination perturbation, Ic1 equilibrium salinity state is restored (**Figure 8**). The third salinity perturbation experiment, D1c simulation, is the same as D1a simulation but a large perturbation flux is applied (as in D1b). Results show that the equilibrium salinity state in D1c and D1b simulations are similar (figure not shown), but the transition period in D1c simulation (3 years) to reach Ic1 equilibrium is shorter compared to D1a simulation (>9 years).

In summary, the statistical analyses of D1a and D1b simulation experiments show the existence of a single stable Gulf equilibrium salinity of about 40.5 g/kg, which is an upper limit in the current climate.

CONCLUSION

Discharge of brine into the Gulf may cause costly environmental problems to marine ecosystems. To ensure the long-term sustainability of seawater desalination, which has grown rapidly in the last five decades, it is necessary and urgent to determine current Gulf equilibrium salinity and its response to seawater desalination perturbation. Model results from Ic1 simulation show a close agreement with multiple observational datasets including reanalysis temperature and salinity, and satellite-derived Gulf WSE, which is also directly influenced by the Gulf basin salinity and temperature. Thus, Ic1 simulation is a high-fidelity representation of the present Gulf system.

The current overturning circulation between the Gulf and the Indian Ocean is an anti-estuarine circulation (lateral inflow and outflow in the low- and high-density layers, respectively) with the year-round counterclockwise overturning circulation. In this anti-estuarine overturning circulation, Gulf WSE controls inflow to the Gulf: lower Gulf WSE, relative to the Indian Ocean, causes the Indian Ocean surface flow to the Gulf to increase, and vice versa. Strong evaporation removes freshwater excessively from the Gulf surface, resulting in high salinity (dense) water that drive more Gulf bottom waters to the Indian Ocean. Consequently, the anti-estuarine overturning circulation in the Gulf depends upon an adjustment of two opposite tendencies: thermal expansion, a function of Gulf heat content, and haline contraction, a function of Gulf salt content. If a positive salinity perturbation is applied to Gulf equilibrium salinity, such as in Ic3 and D1b experiment (**Figures 7A, 8**), then the resulting haline contraction causes flow out of the Gulf to increase until the equilibrium salinity is restored. Therefore, the anti-estuarine overturning circulation between the Gulf and the Indian Ocean is a dynamical upper limit, associated with a stable equilibrium salinity under the current climate.

The hydrodynamic characteristics of this dynamical upper limit (i.e., Ic1 experiment, the present Gulf salinity state) that are described here are consistent with findings from several other observational and modeling studies, including: (1) fresher lateral surface water inflow in northern regions of the Strait

and hypersaline lateral Gulf outflow in southern regions of the Strait near the Omani slope (**Figures 6A,B**; Johns et al., 2003; Pous et al., 2004); (2) seasonal variation and vertical structure of salinity at the Strait (**Figures 6A,B**; Kämpf and Sadrasab, 2006; Yao and Johns, 2010); (3) seasonal variation of lateral inflow and outflow (**Figures 5D, 6A,B**), and (4) interannual variation of Gulf basin-average salinity and volume transport of the inflow and outflow at the Strait (see estimates in **Table 2**; Pous et al., 2015; Lorenz et al., 2020). The salt in the Gulf comes from the Indian Ocean, thus a higher-salinity state will occur only if the anti-estuarine overturning circulation between the Gulf and Indian Ocean brings more salt into the Gulf. This circulation controls Gulf salt content; hence the circulation also controls Gulf salinity stability.

Our results show that the present equilibrium salinity is stable and unlikely to change under current climate ocean and atmosphere boundary conditions. The freshwater sink perturbation in D1b simulation is the same order of magnitude as the current Gulf seawater desalination capacity, about 5 km³/year. Thus, D1b simulation shows that seawater desalination comparable to current capacity does not increase Gulf annual mean basin-average salinity from its current level of 40.5 g/kg (± 0.3 g/kg). However, under a changing climate, if seawater desalination capacity in the Gulf continues to expand to meet the projected freshwater demand in the region (World Bank, 2012), then it is possible that a new equilibrium salinity may be established in the future. However, the impact of changing future climate on Gulf equilibrium salinity is beyond the scope of this work and should be examined in future studies.

DATA AVAILABILITY STATEMENT

The original contributions presented in the study are included in the article/**Supplementary Material**, further inquiries can be directed to the corresponding author/s.

AUTHOR CONTRIBUTIONS

HI and EE conceived and designed the experiments. HI performed the modeling experiments. HI, EE, and PX analyzed the data and co-wrote the manuscript. All authors contributed to the article and approved the submitted version.

ACKNOWLEDGMENTS

The authors are grateful for the support of the MIT & Masdar Institute Cooperative Program (MIT&MICP). This is Contribution No. 75 OF THE GLRC at Michigan Tech. We thank Chenfu Huang for his assistance in the model data processing.

SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/fmars.2020.00573/full#supplementary-material>

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Conflict of Interest: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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