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Multi-band SAR intercomparison study in the Antarctic Peninsula for sea ice and iceberg detection

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Synthetic aperture radar (SAR) systems are one of the best resources to gather information in polar environments, but the detection and monitoring of sea ice types and icebergs using them is still a challenge. Limitations using singlefrequency images in sea ice characterization are well known, and using different SAR bands has been revealed to be useful. In this paper, we present the quantitative results of an intercomparison experiment conducted by the Argentine Naval Hydrographic Service (SHN) using X-, C-, and L-bands from COSMO-SkyMed, Sentinel-1, and SAOCOM satellites, respectively. The aim of the experiment was to evaluate SAOCOM for its use on SHN products. There were 25 images with different SAR parameters that were analyzed, incorporating the diversity in the information that everyday Ice Services attend to. Particularly, iceberg detections, fast first-year ice, and belts and strips were studied in the Antarctic Sound, the surroundings of Marambio Island, and Erebus and Terror Gulf. The results show that the HV polarization channel of the L-band provides useful information for iceberg detection and fast first-year ice surface feature recognition and is a promising frequency for the study of strip identification under windy sea conditions and snow accumulation on first-year ice.

KEYWORDS

SAOCOM, L-band, Sentinel-1, C-band, COSMO-SkyMed, X-band, multifrequency analysis, operational monitoring

1 Introduction

In a climate change context, Ice Services are facing new challenges for operational monitoring, with unusual dynamics and changes in old regimes of sea ice and icebergs. This scenario reveals the significance of having the most updated information available in operational times, in which any source is valuable, either provided by satellite imagery or, less frequently, from *in situ* observations. The capability of synthetic aperture radar (SAR) systems to monitor the polar environments is known; nevertheless, the complexity of sea ice forms, the different properties and surface features of the ice, the variability in environmental conditions, and the great diversity in SAR parameters still make its use a challenge (Onstott and Shuchman, 2004; Dierking, 2013).

The broad scientific literature use SAR for sea ice characterization and iceberg detection (Dierking, 2013; Lyu et al., 2022); however, the incidence angles used are generally bounded to a few degrees of variability or its effect is not considered (Lyu et al., 2022), the image acquisition modes are the ones with the best spatial resolution or only one type of acquisition mode are used (Zakhvatkina et al., 2012; Casey et al., 2016; Johansson et al., 2018; Wang and Li, 2021), full polarization SAR images are considered (Drinkwater et al., 1991; Johansson et al., 2017; Singha et al., 2018), and the few multifrequency analyses that exist are commonly done using the same remote platform, without the limitation of the time gaps between the acquisitions (Matsuoka et al., 2001; Lyu et al., 2022).

Since the launch of the SAR SAOCOM (Satélite Argentino de Observación COn Microondas) in 2018, the Argentine Naval Hydrographic Service (Servicio de Hidrografia Naval, SHN) has incorporated the L-band in its ice charts (Scardilli et al., 2022). Studies with SAOCOM have been done for calibration (Azcueta et al., 2022), surface deformation (Roa et al., 2021; Viotto et al., 2021; De Luca et al., 2022), digital surface model generation (Seppi et al., 2021), soil salinity and moisture estimation (MaChado and Solorza, 2020; Anconitano et al., 2022), forest biomass retrieval (Blomberg et al., 2018), and glacial movement (Ferreyra et al., 2021). However, there were not published results of SAOCOM for sea ice and iceberg conditions (Lyu et al., 2022).

The synergistic use of multifrequency SAR systems has been proven beneficial for ice charting and iceberg detection (Drinkwater et al., 1991; Singha et al., 2018; Dierking, 2021; Dierking et al., 2022). Analysis of multiple SAR frequencies contributes additional information on the observed target due to the different interaction between the ice surface and the signal, with the scattering mechanism and the penetration depth depending on the length scales of the radar waves (Onstott and Shuchman, 2004; Dierking, 2013). The C-band operates in wavelengths of 3.8–7.5 cm, and it is the accepted frequency for all-season sea ice monitoring in the scientific literature and at the Ice Services (Maillard et al., 2005; Arkett et al., 2007; Flett et al., 2008; Dierking, 2013; Scardilli et al., 2022). It was widely studied for sea ice concentration, characterization, drift, thickness, and melt detection, and for iceberg detection (Kwok et al., 2003; Zakhvatkina et al., 2012; Dierking, 2013; Wang and Li, 2021; Lyu et al., 2022).

The X-band, with a wavelength of 2.4–3.8 cm, is less used to Ice Services in daily ice charts because of its similar response to the Cband, although it was suggested that it is more sensitive to the ice surface's properties (Dierking, 2013; Han et al., 2020). In regard to the L-band, a wavelength of 15–30 cm, it was mentioned that it shows superior performance for iceberg detection in rough seas and inside a sea ice field in HV polarization, showing a brighter response (Dierking et al., 2022). Furthermore, it showed a greater distinction between deformed and smooth sea ice, with higher sensitivity to the surface topography such as ridges, hummocks, and rubble fields (Dierking and Busche, 2006; Dierking et al., 2022), and the use of the L-band during the melt season was proposed since it is less sensitive to the first stages of melting (Arkett et al., 2008; Casey et al., 2016).

In 2020, the SHN carried out an Intercomparison Experiment (IE-2020) using different SAR frequencies to test the L-band for its inclusion in daily products and ice advisory. The experiment used the L-band from SAOCOM, the C-band from Sentinel-1, and the X- band from COSMO-SkyMed ("COnstellation of small Satellites for the Mediterranean basin Observation") in the north of Peninsula Antarctica. Ice analysts from the SHN interpreted the images in each frequency making deductions, along with meteorological data and sea ice information from the Argentine Antarctic stations at the time of the SAR image acquisition. Conclusions to the IE-2020 were presented at the 21st meeting of the International Ice Charting Working Group and published in Salvó et al. (2022).

From an operational and maritime point of view, the SHN identified different ice conditions that are relevant for the generation of products of safety at sea. Among them are mentioned the following: (i) the amount of iceberg in an area, which is to date manually identified by ice analysts in the iceberg chart (Scardilli et al., 2022); (ii) fast ice features that can be sensitive to the SAR response but not operationally relevant for the ice chart production, making the information in the images difficult to interpret by the ice analyst; and (iii) the belts and strips that represent a hazard to the vessels that navigate in polar waters and are a challenge to monitor by SAR since they can be formed in a couple of hours from not detectable disperse brash ice and camouflage in windy sea ice condition with only one SAR polarization.

In this paper, we present a quantitative study, a less subjective approach, to the qualitative results found in the SHN IE-2020. The information presented in the following sections is divided into the three typical sea ice and iceberg operational conditions relevant to the ice charting production: (i) icebergs, (ii) fast ice features, and (iii) belts and strips.

2 Materials and methods

2.1 Study area and data

The study area corresponds to the north of the Antarctic peninsula (Figure 1), site where the SHN IE-2020 took place, and zone of strategic relevance in the operational tasks of the SHN. Two main areas of interest were analyzed: the first site was the Antarctic Sound with open water and a sea ice regime arranged in strips and belts, and the second area was in Marambio Island and Erebus and Terror Gulf where different stages of development and forms of sea ice occurred, with a high density of icebergs. All the sea ice and iceberg dynamics and characteristics in the study area were known from the monitoring and information of the SHN.

There were 25 images in the X-, C-, and L-bands that were analyzed corresponding to the COSMO-SkyMed, Sentinel-1, and SAOCOM satellites (Table 1). The center frequency in which the satellites operate is 9.6 GHz, 5.405 GHz, and 1.275 GHz, respectively. The images were acquired between May and September 2020. Sentinel-1 images were downloaded from the Copernicus Open Access Hub from the European Space Agency (ESA) https:// scihub.copernicus.eu/dhus/. SAOCOM images were requested by the SAOCOM catalog from the Comisión Nacional de Actividades Espaciales (CONAE) https://catalog.saocom.conae.gov.ar/catalog/, and the COSMO-SkyMed images were acquired by the Agenzia Spaziale Italiana (ASI) and provided to the SHN by CONAE. The level of processing of the SAR images was in High-Resolution Ground Range Detected for Sentinel-1, Detected Image for



SAOCOM, and Single-look Complex Slant for COSMO-SkyMed. For the intercomparison analysis, the dataset was divided into five groups based on an acceptable time gap between acquisitions concerning the ice regimes under study (Table 1). The use of the different groups is detailed in the next subsections.

The SAR images were processed using the Python Snappy module of ESA's Sentinel Application Platform (SNAP) version 8 http:// step.esa.int. Orbit correction was applied to Sentinel-1 data, and then Sentinel-1 and COSMO-SkyMed images were radiometrically calibrated to sigma naught. The exception with SAOCOM in the calibration step was because the images were obtained in Detected Image Level-1B (data projected to ground range, radiometrically calibrated, and georeferenced). Multi-looking was applied to all the images, and looks in range and azimuth were calculated using the higher range and azimuth spacing of the dataset. Subsequently, an ellipsoid correction was applied and the final projection used was the Antarctic Polar Stereographic (EPSG: 3031). To afford a comparable dataset in space resolution for the intercomparison, each image was resampled to a square pixel, the final pixel size corresponded to the largest of all the images (58.4 meters). At last, the backscatter coefficient values were converted into decibels (dB).

Meteorological data provided by the Argentine Meteorological Service (Servicio Meteorológico Nacional SMN) was used to validate the sea ice conditions. The observations were acquired from three Antarctic stations: Esperanza station (63°23′50″S 56° 59′54″W) situated in the Antarctic Sound, Marambio station (64° 14′50.6″S 56°37′39.3″W) in Marambio Island, and Carlini station (62°14′27.4″S 58°40′01.1″W) located in the north of the Peninsula. These stations provided measurements of air temperature, wind speed, wind direction, and precipitation (liquid or solid state) every 3 h in order to accredit the synoptic atmospheric situation.

2.2 Icebergs

The SAR response to the density of icebergs in a given area was analyzed in different SAR frequencies. All the SAR images of groups 1, 4, and 5 were used (Table 1). Particularly from group 2, the Sentinel-1 and SAOCOM images were selected, and for group 3, only the Sentinel-1 ones. The icebergs were counted in polygons of 1 km² in a fast first-year ice using the Sentinel-2 image of the 30th August 2020 near the Marambio station as a reference (Figure 2). The polygons were distributed in the fast ice covering a wide range of iceberg density, from 0 to 48 in the highest-density areas. The Sentinel-2 image, downloaded from the Copernicus Open Access Hub from the ESA, was the only optical image available in the area of high resolution and cloud-free for the period of study. Since the icebergs were trapped in fast sea ice and part of them are in a grounded area, corroborated by the SHN in the Argentine Antarctic Summer Campaign 2019-2020, the Sentinel-2 optical image was considered as a reliable source of validation.

The mean and standard deviation values of the X-, C-, and Lband backscattering coefficients for HH and HV polarization were extracted from the polygons. Explanatory analysis was performed to examined the correlation among the data using the Pearson's correlation coefficient (Equation 1).

$$r = \frac{\Sigma((x-\bar{x})(y-\bar{y}))}{\sqrt{(\Sigma(x-\bar{x})^2 * \Sigma(y-\bar{y})^2)}}$$
(1)

where r represents the Pearson correlation coefficient, x and y are the number of icebergs and the backscatter coefficient in each polygon, \bar{x} is the mean of the x-values, and \bar{y} is the mean of the yvalues. The standard deviations of the x-values and y-values are represented in the denominator of the equation. The influence of the incidence angle in the samples was also extracted and examined. In every SAR frequency, to each SAR acquisition, a linear leastsquare regression model was applied to obtain the relationship between the number of icebergs and the mean backscatter values (Equation 2).

$$Y = b_0 + b_1 X \tag{2}$$

where *Y* is the dependent variable to predict, represented here as the backscatter coefficient, b_0 and b_1 are the intercept and the slope of the regression line, and *X* is the independent variable representing the number of icebergs per 1 km².

TABLE 1 Description of SAR image acquisitions.

Sensor	Acquisition time (UTC)	Group	Polarization	Acquisition mode	Pixel spacing original Rg x Az (m)	Incidence angle
SENTINEL-1	2020-05-10 07:51	1	HH-HV	EW	25.0 × 25.0	17.99-45.92
SENTINEL-1	2020-05-10 07:52		HH-HV	EW	25.0 × 25.0	17.74-45.74
SAOCOM 1-A	2020-05-10 11:09		HH-HV-VV-VH	TOPSAR Wide	58.4 × 58.4	17.60-35.49
COSMO- SkyMed	2020-05-11 19:10		HH-HV	PINGPONG	6.95 × 2.31	56.00-57.16
SENTINEL-1	2020-05-13 08:16	2	HH-HV	EW	25.0 × 25.0	17.94-45.78
SAOCOM 1-A	2020-05-13 19:57		HH-HV	TOPSAR Wide	31.2 × 31.2	24.9- 48.70
COSMO- SkyMed	2020-05-13 20:09		HH-HV	PINGPONG	2.76 × 2.34	18.94-22.21
SAOCOM 1-A	2020-05-21 11:15	3	HH-HV-VV-VH	TOPSAR Wide	58.4 × 58.4	17.62-35.49
COSMO- SkyMed	2020-05-21 19:45		HH-HV	PINGPONG	5.33 × 2.33	39.74-41.78
SENTINEL-1	2020-05-22 07:51		HH-HV	EW	25.0 × 25.0	17.99-45.92
SENTINEL-1	2020-05-22 07:52		HH-HV	EW	25.0 × 25.0	17.73-45.74
SENTINEL-1	2020-08-21 07:44	- 4	HH-HV	EW	40.0×40.0	17.76-45.73
SAOCOM 1-A	2020-08-23 20:09		HH-HV	TOPSAR Narrow	23.0 × 23.0	24.9-38.29
SAOCOM 1-A	2020-08-25 11:15		HH-HV	TOPSAR Wide	31.2 × 31.2	24.9-48.70
SENTINEL-1	2020-08-26 07:51		HH-HV	EW	40.0 × 40.0	17.99-45.92
SENTINEL-1	2020-08-26 07:52		HH-HV	EW	40.0×40.0	17.73-45.74
SAOCOM 1-A	2020-08-26 11:33		HH-HV	Stripmap	5.2 × 5.2	38.20-41.29
COSMO- SkyMed	2020-09-06 19:22	- 5	HH-HV	PINGPONG	6.53 × 2.31	50.57-52.21
SENTINEL-1	2020-09-06 23:51		HH	EW	25.0 × 25.0	17.94-45.80
SENTINEL-1	2020-09-07 07:51		HH-HV	EW	25.0 × 25.0	17.99-45.92
SENTINEL-1	2020-09-07 07:52		HH-HV	EW	25.0 × 25.0	17.73-45.74
SAOCOM 1-A	2020-09-08 20:09		HH-HV	TOPSAR Narrow	23.0 × 23.0	24.9-32.29
COSMO- SkyMed	2020-09-10 19:22		HH-HV	PINGPONG	6.53 × 2.31	50.57-52.21
SENTINEL-1	2020-09-10 08:16		HH-HV	EW	25.0 × 25.0	17.93-45.78
SAOCOM 1-A	2020-09-13 20:03		HH-HV	TOPSAR Narrow	23.0 × 23.0	24.9-38.29

2.3 Fast ice features

In view of what was seen in SHN IE-2020, where spatial patterns were recognized in the different SAR frequencies, the fast ice's features were studied at each sensor to quantify this variability. First, the hole fast ice area was jointly analyzed considering all the sea ice features together in the platform without differentiation ice characteristics. Second, the presence of snow on the fast ice and its differences observed in the SAR images were analyzed.

The analysis of the hole fast sea ice around Marambio Island was analyzed using all the SAR images of groups 1, 4, and 5 (Table 1). Specifically, from group 2 the Sentinel-1 and SAOCOM images were selected, and for group 3, only the Sentinel-1 ones. The fast ice was delineated to its minimum extent covering all the dates present in the images mentioned. For delineating the polygon, photointerpretation of the SAR images was performed and optical imagery was used as validation when available from the Sentinel-2 or MODIS (Moderate Resolution Imaging Spectroradiometer). Sentinel-2 was obtained as mentioned in Section 2.2, and MODIS was consulted on NASA (National Aeronautics and Space Administration) Worldview application https://worldview.earthdata.nasa.gov, part of NASA's Earth Observing System Data and Information System (EOSDIS). After the delimitation of the fast ice area, values were extracted to all the images and the frequency distribution of the data was analyzed.

The analysis of the presence of the snow cover on the fast ice was done comparing the backscatter of bare and snow-covered firstyear ice near Cockburn Island in the Erebus and Terror Gulf. From the Sentinel-2 image of the 30th August 2020, snow-covered and



bare sea ice were identified and circular samples were collected with a 50 m radius (Figure 3). Mean values were calculated to each sample; afterward, to each SAR image, mean values for the snowcovered areas and bare ice were evaluated, and the difference of the same were performed following Equation 3:

$$\mathrm{Diff}_{\mathrm{s}-\mathrm{b}} = 10^* \mathrm{log}_{10}(\sigma_{\mathrm{s}} - \sigma_{\mathrm{b}}) \tag{3}$$

where Diff_{s-b} is the difference between the snow-covered and bare ice areas as the difference of the power per unit area of the snow-covered ice σ_s and the power per unit area of the bare ice σ_b , converted in decibels. The SAR images used to calculate this difference were from the 21st August to 13th September corresponding to groups 4 and 5 (Table 1).



FIGURE 3

Snow-covered and bare samples on the fast ice north of Cockburn Island. Sentinel-2 MSI 30th August 2020, bands 4-3-2. Copernicus Sentinel data 2020.

Meteorological data were used to validate the use of optical imagery as ground truth. Meteorological observations were taken from the Marambio station with 51 days of accumulated days of precipitation registered, showing 102.1 mm of snow recorded since the formation of the fast ice. Maximum and minimum temperatures were below 0°C with a prevailing wind direction from the south to southwest, accumulating 48% of the daily observations in that direction. Patterns in the snow cover to the north of Cockburn Island were visualized in the Sentinel-2 image, confirming the recorded observations (Figure 3).

2.4 Belts and strips

Strips and belts present in the Antarctic Sound were analyzed, and the images used correspond to groups 2 and 3. In order to compare backscatter values for sea and water around, strip and water samples were selected at similar incidence angles (Figure 4). The samples were determined visually using HH and HV polarization of each band. Boxplots were plotted for each pair of samples (strips and water) and for each band so as to facilitate the comparison between samples (comparison of backscatter values and standard deviation). To ensure a detailed analysis of the backscatter values of strips and water, the wind speed and direction data were taken from the Esperanza station, corresponding to the dates of image acquisition (Table 1).

3 Results and discussion

3.1 lcebergs

The Pearson correlation coefficients were low for all the SAR frequencies and above zero, determining a positive weak relation between the number of icebergs and the backscattering coefficients in the X-, C-, and L-bands for each polarization (Figure 5). The C- and L-bands showed a wider value range in the backscattering coefficient since more image acquisition was analyzed in those frequencies.

To quantify the nature of the relationship between the number of icebergs per km² and the backscatter coefficient in different SAR frequencies, linear regression models were performed. The coefficient of determination of the models (R^2) was low for most of the acquisition in HH polarization in all the frequencies and with p-values that were only significant to 0.05 for two images of three in the X-band, five images of eight in the C-band, and six images of seven in the L-band (Table 2). HV polarization showed higher R^2 values than HH polarization with only one image in the X-band, three in the C-band, and one in the L-band with lower values than 0.7; all p-values were significant (Table 2). The slope for the linear model was higher in the L-band than in the X- and C-bands.

The incidence angle where the polygons were located was contrasted with its response in the HH and HV channels. The difference in the backscatter coefficient for areas with low vs. high amounts of icebergs increases with the incidence angle in both polarizations (Figure 6). Because the availability in C- and L-band imagery was larger than the X-band, a broader range of incidence angles was presented in the dataset for those frequencies, making the analysis of the X-band inconclusive for evaluating the effect of



the incidence angle. The L-band showed a large interval between the minimum and maximum of backscatter values in HV polarization. The standard deviation range obtained for the backscattering coefficient in HH polarization was 1.24–1.6 dB for the X-band, 1.14–2.07 dB for the C-band, and 1.75–3.03 dB for the L-band; meanwhile, for HV, polarization was 1.33–1.78 dB for the X-band, 1.13–2.08 dB for the C-band, and 2.5–6.18 dB for the L-band.

The higher variability of the L-band in HV polarization evidences greater sensitivity of the L-band to the presence of icebergs than the C- and X-bands, with the exception of one TOPSAR Narrow image, in which thermal noise was observed on one of the sub-swaths, the position in which the samples were taken. Additional analysis of the TOPSAR Narrow mode is required to ameliorate the effect or reduce the thermal noise in the sub-swath.

3.2 Fast ice features

The fast ice around Marambio Island contains multiyear ice, north of Marambio station, which was grounded prior to freezing



X-, C-, and L-band backscattering coefficients in dB for HH (top) and HV (bottom) polarization vs. different numbers of icebergs in 1 km². Different colors identify different SAR acquisitions.

		нн		HV		
	Х	С	L	Х	С	L
Slope (dB)	0.04 ± 0.02	0.05 ± 0.02	0.10 ± 0.03	0.08 ± 0.02	0.09 ± 0.02	0.18 ± 0.09
Intercept (dB)	-22.86 ± 0.66	-18.27 ± 3.20	-24.35 ± 3.68	-33.53 ± 1.40	-30.81 ± 1.71	-39.55 ± 7.04
R ²	0.36 ± 0.24	0.39 ± 0.20	0.48 ± 0.15	0.66 ± 0.25	0.68 ± 0.13	0.79 ± 0.13
p-value	0.21 ± 0.29	0.06 ± 0.08	0.02 ± 0.05	0.01 ± 0.02	0.00 ± 0.00	0.00 ± 0.00

TABLE 2 Mean and standard deviation for each linear regression per image acquisition in each SAR frequency.

(Figure 7). Of all the frequencies, the L-band was the one that showed a closer representation of HH polarization to the optical images, accurately following the fast sea ice composition and evolution, highlighting the variety of features in the ice (Figure 7).

The distribution of the backscattering response in dB was analyzed for each intercomparison group in the fast ice area (Figure 8). Conclusions about the range of the backscattering coefficient cannot be done because of the wide range in the incidence angles present in the dataset. In HH polarization, the Lband from SAOCOM showed higher variability in the backscatter values as can be seen in the histograms for all the intercomparison dates (Figure 8). The standard deviation in dB obtained for HH polarization in the first group was 3.58 for the X-band, 3.58 for the C-band, and 3.61 for the L-band. For the second and third groups, it was 3.21–3.28 for the C-band and 4.48 for the L-band. Lastly, in the fourth group, the standard deviation was 3.29–3.47 for the C-band and 4.31–5.6 for the L-band, and for the fifth group it was 3.93–4.13 for the X-band and 3.18–4.36 and 4.1–5.26 for the L-band.

HV polarization presented less variability in the data distribution than HH polarization, with similar curves for each frequency, with the exception of SAOCOM images from the fourth and fifth intercomparison groups, corresponding to two images with incidence angles upper 33°, Stripmap and TOPSAR Narrow mode, and a TOPSAR Wide image with a nominal incidence angle of 25.5° in the fast ice area. SAOCOM images from 23rd August and 8th September showed a bimodal distribution that was caused by



X-, C-, and L-band backscattering coefficients in dB for HH and HV polarization vs. incidence angle in degrees. Each color represents one SAR image, and the saturation corresponds to the increase in the number of icebergs sampled.



FIGURE 7

Fast ice's formation near Marambio Island on MODIS Terra. The same area in the X-, C-, and L-band images, backscattered in dB for HH polarization. Copernicus Sentinel data 2020. SAOCOM® Product – © CONAE – 2020, All Rights Reserved. COSMO-SkyMed Product – [©] ASI 2020 processed under license from ASI - Agenzia Spaziale Italiana, All Rights Reserved, and distributed by CONAE.



Histogram distribution of SAR images in the X-, C-, and L-bands for (from top to bottom) groups 1, 2 and 3, 4, and 5. HH polarization is represented to the left and HV to the right.

the thermal noise in one of the sub-swaths of the TOPSAR Narrow Mode, as it was explained in the previous subsection.

For the analysis of the snow presence on the fast ice, the SAR frequency did not show a distinction between the backscatter coefficient of bare and snow-covered first-year ice, and the variability observed in the results was attributed to differences in the SAR parameters. However, the polarization channel showed an impact on the snow detection with all the differences between the bare and snow-covered positive in HH polarization, 75% of the images above 2 dB of difference (Figure 9). In HV polarization, the differences were lower, with three images from SAOCOM and from COSMO-SkyMed below zero (Figure 9). Patterns in the variability presented in the results considering the SAR frequency, incidence angle, resolution of the image, direction of the antenna in relation to the area sampled, and weather conditions were studied; nonetheless, causes cannot be attributed because the differences in the dataset and more images need to be included to reach conclusions.

In HH polarization, SAOCOM showed an increase in the bare and snow-covered difference between 23rd and 26th August, a period in which snowfalls occurred, and a similarity in the bare and snow-covered difference from 8th to 13th September, a date interval without precipitation (Figure 9). Further research needs to be done on the L-band, since the results revealed a sensitivity to snow precipitation but the dataset is not large enough to conclude.

3.3 Belts and strips

Figure 10 shows the images analyzed on the Antarctic Sound; the backscatter values of strips and water were compared, and the result is shown in Figure 11.

For most of the analyzed cases, the HV polarization showed an advantage for strip detection with respect to HH polarization; this result was noted in the visual and quantitative analyses (Figure 11) where the backscatter values for the water box were below the values for the strip ice box. A clear example was found in the C-band image from 13th May, where it was not possible to distinguish the presence of ice in HH (both samples had similar backscatter values), whereas in HV the presence of strips was observed. This advantage of HV polarization may be due to a higher sensitivity of HH polarization to the wind effect on the ocean surface. This result is consistent with previous studies where the water backscatter values in HV are lower than those of HH even under wind roughened conditions (Shuchman and Flett, 2003; Arkett et al., 2007).

In some cases, the sea surface presented more backscattering than sea ice like X-band image from 13th May, where sea ice had dark tones and the backscatter values for the strip ice box were below the values for the water box. This could be a result of a greater roughness of the sea surface with respect to the ice surface. The dark tones for sea ice in SAR images must be taken into account by the ice analysts when the presence of sea ice is analyzed. This effect was





X-, C-, and L-band backscattering coefficients in dB for HH and HV polarization. Copernicus Sentinel data 2020. SAOCOM® Product – © CONAE – 2020, All Rights Reserved. COSMO-SkyMed Product – © ASI 2020 processed under license from ASI – Agenzia Spaziale Italiana, All Rights Reserved, and distributed by CONAE.

reported previously in the ice edge zone by Onstott and Shuchman (2004) and by Arkett et al. (2007).

In the inter-band analysis, it was observed that in the case of group 2, in the C-band image with HH polarization, it was not possible to identify and differentiate the sea ice from the water around it, unlike the X- and C-band images with HH polarization. According to the wind measurements reported at the closest meteorological station to the study area, there was a decrease of approximately 60% in the intensity of the wind between the acquisition of the C-band and X- and L-band images (Figure 12). An increase in the intensity of the wind over the sea produces a

growth in the roughness of the sea surface, in turn increasing the backscattering of the surface. This results in sea surface roughness, and its response could be limiting the detection of sea ice at HH polarization.

The X- and L-band images of group 2 were acquired within 12 min apart, and no evidence was found that the meteorological conditions were significantly modified in that time period. However, a better distinction was noted for L-band both visually and in the boxplots, with a greater difference between samples. The differences between X- and L-band could be due to the intrinsic characteristics of each band. This would represent an advantage for



the L-band in distinguishing sea ice strips from water, being L-band less sensitive to the wind or the sea state compared with X-band as Dierking et al., 2022 indicated in their study. However, differences in the incidence angles of each acquisition make it difficult to determine a conclusive cause.

For the acquisition dates of group 3, no significant variations in wind intensity (Figure 12) or a particular pattern in wind behavior were found. In all the bands, it was possible to identify and differentiate the water from the strips. The differences between bands may be due to the characteristics of each image acquisition (space resolution and incidence angle).

4 Conclusion

This paper presents the quantitative results obtained from a detailed previous Intercomparison Experiment carried out in 2020 (IE-2020) by the Argentine Naval Hydrographic Service using the

X-, C-, and L-bands from the COSMO-SkyMed, Sentinel-1, and SAOCOM, respectively. The analysis was performed following typical sea ice and iceberg operational conditions: (i) icebergs, (ii) fast ice features, and (iii) belts and strips.

The results found with the HV polarization of SAOCOM suggested the L-band as a convenient frequency for the detection of areas with a high density of icebergs surrounded by first-year ice. The number of icebergs showed positive but weak correlations with the backscatter coefficient in HH polarization for all the sensors; however, significant results were obtained in HV polarization, even under different SAR parameters such as image acquisition and different incidence angles. The linear regressions presented showed the L-band as the frequency with a higher slope and coefficient of determination, being significant in HV polarization, with good results even under spatial resolutions of around 100 m for the TOPSAR Wide Quad polarization mode.

As it was seen from visual interpretation in the IE-2020, the analysis of the fast first-year ice near Marambio Island revealed a



higher sensitivity to the surface features of the L-band, consistent with the evolution and composition of the fast ice. No influence in the backscattering coefficient was observed on the different SAR frequencies in relation to the snow layer on the first-year ice; nevertheless, polarization was important in its detections, with measurable differences in the HH channel.

The identification of strips and belts in seawater depended on the sea state and wind conditions in the area. The HV polarization represented an advantage in the identification in relation to HH polarization because of the lower sensitivity to the surface water roughness. The L-band showed to be less sensitive to the roughness of the sea surface; however, further analysis is needed to determine a benefit in the detection of strips with the L-band. Additionally, future research has to be done to incorporate in the SAOCOM image processing workflow the reduction in sub-swath thermal noise, as was seen in two TOPSAR Narrow images from our dataset.

In addition to the limitation in part of the results found, the dataset used in this intercomparison showed the typical operational situation in the Ice Services, where any SAR acquisition is valuable to approach the last most updated condition of the sea ice and icebergs. Further research still needs to be developed in view of incorporating the broad variability in SAR parameters, such as diversity in acquisition modes, resolution, polarization, incidence angle, different environmental conditions, and season of the year to generate useful outcomes for decision-making.

Data availability statement

The raw data supporting the conclusions of this article will be made available by the authors, without undue reservation.

Author contributions

CS: Writing – original draft, Writing – review & editing. LG: Writing – original draft, Writing – review & editing. JA: Writing – original draft, Writing – review & editing.

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Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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